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# An assessment of wind and wave climate as potential sources of renewable energy in the nearshore Shenzhen coastal zone of the South China Sea

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#### Abstract

In this study, nearshore wind and wave climates and their potential as renewable energy sources are evaluated by means of buoy observational data for the Shenzhen coastal region. Six buoys were originally deployed in the region by the city local government of China in 2014, and are located in different areas of the study region, including Dapeng Bay, Daya Bay, Shenzhen Bay. The waters in these areas are relatively shallow, ranging in depth between about 3-22 m. The results show that during 2014-2016, annual mean wind speeds (at 2.5 m above the sea surface) in the region varied between  $3.1-4.1 \text{ m s}^{-1}$ , leading to wind powers between  $37-94 \text{ W m}^{-2}$ ; significant wave

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heights were mostly less than 1 m, while wave energy periods were mostly in the range 3-7s. As a result, wave power was mostly less than  $1.0 \text{ kW m}^{-1}$ . It is concluded that the potential of wave energy as a renewable resource at the buoy locations was very small. This may be due to the fact that, first, water depth is very shallow, and, secondly, the buoys are located in bays where the sea is somewhat semi-enclosed, all of which are not favourable for the development of wind waves.

*Keywords:* Wave climate, Wave energy, Shenzhen, Wind climate, Wind energy, the South China Sea

#### 1 1. Introduction

Energy has become one of the hottest words in China. On the one hand, 2 the demand for energy in China is rapidly increasing, since China has become 3 the largest energy consumer and producer in the world (US Energy Infor-4 mation Administration, EIA; http://www.eia.gov/), whereas on the other, 5 energy is closely associated with the environment. For example, coal and oil 6 consumption in the country has resulted in very seriously bad air pollution impacts and environmental problems, which have been reported on in count-8 less situations. Air pollution and smog in China have become one of the most 9 contentious issue for the international media. Hundreds of millions of people 10 in the world's most populous country are suffering the effects of this pollu-11 tion, which is putting a lot of pressure on environmental and public health 12 conditions in China. To mitigate these problems, China has to accelerate 13 the adjustment of energy structure and to increase the share of clean sources 14 in its energy mix. According to China's Action Plan for the Prevention and 15

Control of Air Pollution [1], China desires to reduce coal consumption to less
than 65% in terms of total energy consumption by 2017.

China has a very long coastline, possessing rich ocean resources, which 18 attach great importance to marine development and exploitation of renewable 19 energy. As important renewable types, ocean wind and wave energy not only 20 provide China with energy sources, but also resources with which to address 21 and relieve the challenge of energy demands with resepect to the environment 22 while implementing a sustained development strategy. As [2] mentioned, 23 renewables can also provide tools to address many pressing needs, including 24 improving energy security, reducing human health problems, and mitigating 25 against greenhouse gas emissions. 26

In recent times (decades), previous researchers have made great contribu-27 tions toward the assessment of wind/wave energy potential for various seas 28 in many regions and countries, based on the analysis of wind/wave data col-29 lected from buoys, remote sensing, numerical hindcasts, and combinations of 30 these sources. Included among these are the following studies: for the UK 31 [3, 4], Portugal [5–7], Sweden [8], Belgium [9], Spain [10–13], Ireland [14, 15], 32 Europe [16], the North Sea [9], the Baltic Sea [17, 18], the Red Sea [19], the 33 Caribbean [20], Australia [21, 22], Canada [23], Iran [24], India [25–29], Korea 34 [30], Singapore [31], Chile [32], the Hawaiian islands [33, 34], Southern New 35 England [35], California [36, 37], the Atlantic coast of the southeastern USA 36 [36, 38], the US Pacific Northwest [39], as well as for the global ocean (e.g., 37 [40-44]). In addition, wave/wind energy resource assessment has also been 38 conducted for China [e.g., 45–54], and also including for Hongkong [55–58]. 39

The region of interest in the present study is near the coast of Shenzhen,

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located at the northern extreme of the South China Sea. Shenzhen shares a 41 border with Hong Kong to the south, is 160 km south of the provincial capital 42 of Guangzhou, and 70 km south of the industrial city of Dongguan. To the 43 west, the resort city of Zhuhai is a 60 km away, (see Fig. 1). Shenzhen was 44 the earliest of the five special economic zones in China, originally established 45 in 1979, and was given the right of provincial-level economic administration. 46 Since then, it has been one of the fastest growing cities in the world, and 47 eventually became one of the largest cities in the Pearl River Delta region 48 from the one-time small fishing village, and one of the economic powerhouses 49 of China, as well as of the largest manufacturing bases in the world. As a 50 result of this tremenduous economic growth, the demand for energy is no 51 doubt correspondingly rapidly increasing. The first nuclear power plant in 52 China was built in the coast of Daya Bay (see Fig. 1), a coastal region of 53 Shenzhen. However, relative to the economic growth, the marine develop-54 ment of Shenzhen has fallen far behind. It wasn't until 2014 that the city 55 local government put buoys in the surrounding waters to observe and moni-56 tor the atmospheric and hydrodynamic climate of the region, and for use in 57 marine and meteorology environmental studies and forecasting, and so on. 58 The locations of the buoys have been shown to be reasonably representative 59 of the hydrodynamic climate of different areas in this region, in a limited ex-60 penditure, and, from these, 6 buoy locations were selected (see section 2 for 61 details). This was a great progression and good start for marine observations 62 and monitoring for this region. 63

<sup>64</sup> By means of data collected from six buoys located in the nearshore Shen-<sup>65</sup> zhen zone, this study aims to evaluate wind and wave climate for the region,

in terms of wind speed and direction, wave height, wave period and wave di-66 rection. The potential of wind and wave energy as resources is assessed based 67 on buoy observations for the period 2014-2016. Location of the area of in-68 terest and accumulated data at each of the six buoys, and also bathymetry 69 of the region are described in detail in section 2, where the methods for es-70 timating wind and wave power are also presented. Wind and wave resource 71 variability in the region are investigated and discussed in section 3. Finally, 72 conclusions are presented in section 4. 73

#### 74 2. Data and methods

#### 75 2.1. Study area and buoy data

The region of wind and wave energy resource under investigation extends 76 from  $22.0^{\circ}N-23.0^{\circ}N$  and  $113.5^{\circ}E-114.5^{\circ}E$  (Fig. 1), and includes the entire 77 nearshore region of Shenzhen adjacent to Hong Kong. Six buoys, acquired 78 from Shenzhen Marine Montoring Forecasting Center, are available in the 79 study region, and their locations and corresponding mean water depths are 80 listed in Table 1. Data from these buoys represent the wind and wave cli-81 mate for different areas of the Shenzhen coastal region, mainly to monitor 82 atmospheric and hydrodynamic changes in this region and also to be used for 83 marine environmental investigating, forecasting, etc. The study area consists 84 of the waters extending from Shenzhen Bay in the northwest, southward and 85 eastward containing waters surrounding Hong Kong to Dapeng Bay, east of 86 Hong Kong and further east to include Daya Bay. The six buoys are located 87 very close to shore in very shallow water across the three bays: B1 is located 88 in the northwest end of Dapeng Bay in water depth of about 11 m; B2 is also 89

located in Dapeng Bay, close to the western side of Dapeng Peninsula; B3
is located in Aozaixia Bay in inner Daya Bay in water depth of only about
3m; B5 is situated in southern Daya Bay (water depth: 12m); B4 is located
in the relatively open area off the tip of Dapeng peninsula in water depth of
about 22m; and B6 is located in inner Shenzhen Bay, in a relatively narrow
and closed area with water depth only 3m.

The six buoys have been in operation since April 2014. Wave data were 96 provided hourly, consisting of wave parameters, including wave height, period 97 and direction. However, some wave data were not recorded over a span 98 of several days, and data gaps for each buoy can be inferred from Fig. 7. 99 Significant wave height  $(H_s)$ , which is identical to the average of the highest 100 one-third of all wave heights recorded during each wave acquisition, is utilized 101 and analyzed for wave energy assessment in this study, while the measured 102 wave period acquired from buoy measurements refers to mean wave period, 103  $T_m$ . However, only buoys B1-B4 and B6 provided records of wind data 104 (at 2.5 m above sea surface), including wind speed and direction at quarter-105 hourly intervals, except for B4 (mainly half-hourly). Wind direction at B6 106 was also missed. 107

#### 108 2.2. Analytical methods and approach

Wind power is defined as the power per unit section perpendicular to wind flow, and is computed in this study by the following equation [52]:

$$W = \frac{1}{2}\rho_a V^3,\tag{1}$$

in which W is wind power in units of W m<sup>-2</sup>, V is wind speed (unit: m s<sup>-1</sup>), and  $\rho_a$  is air density taken as  $1.292 \,\mathrm{kg m^{-3}}$  and corresponding to that close to the sea surface in the region of interest in this study.

It was noticed that, for wave power calculations and assessments, the 114 widely used wave period is the so-called wave energy period,  $T_e$ , instead of 115  $T_m$ .  $T_e$  can be defined as  $T_e \equiv T_{-10} = \frac{m_{-1}}{m_0}$ , in which  $m_n$  is the  $n^{th}$  moment 116 of spectral density, i.e.,  $m_n = \int_0^{2\pi} \int_0^\infty f^n S(f,\theta) df d\theta$ , and here f is the wave 117 frequency,  $\theta$  is the wave direction, and  $S(f, \theta)$  is the 2D wave spectrum. In 118 general, the observed wave period measured by buoys for real sea states is 119 rarely specified by  $T_e$ , but is specified in terms of the mean wave period  $T_m$ , or 120 in terms of the peak period  $T_p$ .  $T_e$  is often estimated by means of its relation 121 to other observational wave periods, such as  $T_m$  and  $T_p$ , when the spectral 122 density is unknown [40]. Therefore, the relationship between  $T_e$  and  $T_p$  can 123 be estimated by the formula  $T_e = \alpha T_p$ , in which  $\alpha$  depends on the shape of 124 the wave spectrum used to define the sea state. The relationship between 125  $T_p$  and  $T_e$  used in this study is computed by a conservative approximation 126 that  $T_e = 0.9T_p$ , according to the study of [40]. This relationship has been 127 widely adopted in assessing the wave energy resource such as off the coast of 128 Canada [23], in the North Sea [9], and for the global ocean [40] as well as for 129 the offshore wave power in the East China Sea [54]. 130

Based on the study of [8, 10-12, 28, 41, 54], the wave power, P, known as wave energy flux as well, is calculated by the following expression

$$P = \frac{\rho_w g^2}{64\pi} H_s^2 T_e,\tag{2}$$

where  $\rho_w$  represents sea water density taken as 1025 kg m<sup>-3</sup>, the average sea water density in the study area. Since  $T_p$  was not provided in the buoy records acquired in this study,  $T_m$  was used to estimate the wave power in the region of interest. Following the study of [54, 59], the relationship between <sup>137</sup>  $T_m$  and  $T_p$  is adopted as  $T_p = 1.2T_m$ , and as a result,  $T_e$  investigated in the <sup>138</sup> present study is computed by means of  $T_m$  as  $T_p = 0.9 \times 1.2T_m = 1.08T_m$ .

#### 139 3. Results and Discussion

#### 140 3.1. Wind climate and assessment of wind energy potential

To assess the wind climate in the Shenzhen coastal region, the time series 141 of wind speeds based on the buoy measurement data (B1-B4 and B6) for the 142 period 2014-2016 is plotted by Fig. 2. It can be observed that wind speeds 143 in the study area were mostly less than  $8 \,\mathrm{m \, s^{-1}}$  for the 5 buoys except B4 144 buoy, where the wind speed was generally less than  $10 \,\mathrm{m\,s^{-1}}$ ; relatively high 145 wind speeds of greater than  $15 \,\mathrm{m\,s^{-1}}$  occurred occasionally. Based on Eq. 2, 146 Fig. 3 displays the calculated wind power for the 5 buoys. It shows that the 147 wind power for the 5 buoys (except B4) was mostly less than  $300 \,\mathrm{W m^{-2}}$ , 148 while B4 shows relatively larger wind speeds, with values mostly smaller 149 than  $500 \text{ W} \text{m}^{-2}$ . Large wind power values of more than  $3500 \text{ W} \text{m}^{-2}$  can 150 occasionally be found in the observed time. 151

The fundamental characteristics of wind energy resources, in terms of 152 the annual mean wind speed with its standard deviation  $((V_{mean} \pm \text{std.dev.}))$ , 153 maximum wind speed  $(V_{max})$ , annual mean and maximum wind power (i.e., 154  $W_{mean}$  and  $W_{max}$ ), is summarized in Table 1. In Dapeng bay, represented 155 by buoys B1 and B2, the annual mean wind speed during 2014-2016 was 156  $3.1 \,\mathrm{m\,s^{-1}}$ , and in Daya Bay, represented by B3,  $V_{mean}$  was  $3.4 \,\mathrm{m\,s^{-1}}$ ; simi-157 larly, in Shenzhen bay, as B6 shows,  $V_{mean}$  was 3.6 m s<sup>-1</sup>; relatively stronger 158 wind speed was found at B4, with mean wind speed of  $4.1 \,\mathrm{m \, s^{-1}}$ . In ac-159 cordance with the mean wind speed, the annual mean wind power,  $W_{mean}$ , 160

at B1 and B2 was, respectively,  $58 \text{ W m}^{-2}$  and  $37 \text{ W m}^{-2}$ , and  $W_{mean}$  was around 50 W m<sup>-2</sup> in Daya bay and Shenzhen bay, highest  $W_{mean}$  among the buoys was found at B4 with value of 94 W m<sup>-2</sup>.  $V_{mean}$  averaged at the buoys was about  $3.5 \text{ m s}^{-1}$ , leading to an average  $W_{mean}$  of around  $58 \text{ W m}^{-2}$  for 2014-2016.

For the period 2014–2016, maximum wind speed,  $V_{max}$ , at B1, B2, and B4 was, respectively,  $17.5 \text{ m s}^{-1}$ ,  $17.6 \text{ m s}^{-1}$ ,  $17.1 \text{ m s}^{-1}$ , leading to wind power of more than  $3000 \text{ W m}^{-2}$ , while  $V_{max}$  was relatively smaller at B3 and B6, with values of  $15.66 \text{ m s}^{-1}$  and  $15.7 \text{ m s}^{-1}$ , respectively, giving a corresponding  $W_{max}$  of over  $2300 \text{ W m}^{-2}$ .

It is noted that the study region is influenced by tropical cyclones (TCs; 171 normally called typhoons in China) relatively frequently, and they have been 172 reported many times. Therefore, the variance of maximum wind speeds can 173 be quite large for different time periods. Maximum wind speed depends 174 greatly on the extent and degree of the TC effects. For example, for the 175 period 2014–2016, 3 TCs passed through the study region that had much 176 influence it, see Fig. 4. The periods of these TCs are plotted in Fig. 2, and 177 it is observed that wind speeds in these time were relatively high. TCS are 178 normally classified into different categories. In China, in accordance with 179 the World Meteorological Organization's recommendation, the classification 180 is divided into 6 categories by the classification of TCs standardization of 181 China  $(GB/T19201-2006 \ [60])$ , in terms of wind speed averaged over a pe-182 riod of 10 minutes near the center of the TC. The six classifications are 183 as follows: Tropical Depression (TD;  $10.8-17.1 \,\mathrm{m \, s^{-1}}$ ); Tropical Storm (TS; 184  $17.2-24.4 \text{ m s}^{-1}$ ; Severe Tropical Storm (STS;  $24.5-32.6 \text{ m s}^{-1}$ ); Typhoon 185

 $(TY; 32.7-41.4 \text{ m s}^{-1})$ ; Severe Typhoon (STY; 41.5-50.9 m s<sup>-1</sup>); and Super 186 Typhoon (SuperTY;  $\geq 51.0 \text{ m s}^{-1}$ ). Fig. 4 shows that TCs Nida (No. 1604) 187 and Haima (No. 1622) were in the classification of STY when they landed 188 (around Aug. 2 and Oct. 21 in 2016, respectively), and correspondingly, the 189 wind speeds could reach up to  $15 \,\mathrm{m\,s^{-1}}$  at B1, B2 and B4 locations, while 190 they were relatively small (about  $10 \text{ m s}^{-1}$ ) at locations B3 and B6. Linfa 191 (No. 1510) was much weaker when it came to the Shenzhen region, and 192 its influence on wind speed was much less. The other 3 TCs displayed in 193 Fig. 4 were relatively far from the study region, but their influence is still 194 seen: wind speed at the 4 buoys were all relatively large with values reaching 195  $15 \,\mathrm{m\,s^{-1}}$ , in the influence of the Kalmaegi (No. 1415); the other two TCs 196 were also evident for the wind speeds in the study area. 197

The seasonal and monthly mean wind speed variations at the wind buoys 198 in the study area are presented by Fig. 5. For the present study, the four 199 boreal seasons are winter (December-February), spring (March-May), sum-200 mer (June-August), and Autumn (September-November). The monthly 201 wind speeds over the period 2014-2016 varied from about  $2-6 \text{ m s}^{-1}$ , with 202 different variability found at each of the buoys. Highest monthly mean wind 203 speeds occurred at station B4, with values of about  $5.6 \,\mathrm{m \, s^{-1}}$  in November 204 2015, followed by station B1 ( $5.2 \,\mathrm{m \, s^{-1}}$  in December 2014). However, at lo-205 cation B6 in Shenzhen Bay, the largest monthly wind speed was not found 206 in winter, but in June 2015  $(4.3 \,\mathrm{m \, s^{-1}})$ , followed by June 2016  $(4.2 \,\mathrm{m \, s^{-1}})$ , 207 while the smallest monthly value was reported in October 2016. 208

Corresponding to wind speeds, monthly mean wind powers at the 5 buoy locations varied from  $20 \text{ Wm}^{-2}$  to  $200 \text{ Wm}^{-2}$ : relatively large wind powers can be found in autumn and winter months at B1 and B4; while at B6 wind
powers in June and July were larger than those in winter months.

Fig. 5 also shows that the variation of monthly mean wind speeds and 213 powers within a year was largest at B1, followed by B4, and was smallest 214 at B3. Furthermore, no coincident seasonal variability was found in the 215 winds and the wind powers at any of the buoy locations for 2014-2016. Of 216 these, stations B1 and B4 show similar seasonality: winds and wind powers 217 were relatively large in autumn and winter, and were smaller in the spring 218 and summer months; however, B2 and B3 locations did not show evident 219 seasonality. Note that the data in some months for B2 was lacking, which 220 may influence the accuracy of its seasonality. On the contrary, B6 location 221 shows opposite seasonality when compared with B1 and B4 during the study 222 period: winds and wind powers were relatively large in spring and summer, 223 and were smaller in autumn and winter months. 224

Besides wind speed, wind direction is another important parameter in 225 wind energy assessment. For the data collected at the 6 buoys, only that at 226 buoys B1–B4 include wind direction data. Wind direction and wind speed 227 for different seasons at the 4 buoys is presented in terms of a wind rose figure. 228 Fig. 6. It can be seen that variability in wind direction is different for the 229 4 buoy locations. At location B1, in winter, easterly winds prevailed (about 230 31%), followed by winds directed from the ESE (about 28%), and occasionally 231 from the ENE and N (about 7%), whereas westerly winds are least frequent 232 (less than 5%); in spring, prevailing direction was from the II quadrant, 233 i.e., from E to S, with least occurrence also from the west; in autumn and 234 winter, winds were mostly southerly, followed by SSE and SSW directions, 235

respectively, and winds from land, i.e., from the IV and I quadrants, provided 236 the smallest contribution to the wind energy at B1. For B2, the prevailing 237 wind direction was easterly in all seasons (over 45%), followed by ENE, and 238 lowest occurrence occurred from SE to NW. At location B3 in Daya Bay, for 239 all seasons the largest contribution to wind energy resources was provided 240 by southerly winds, and SSE and SSW winds also provided considerable 241 contributions. At the B4 location, the prevailing wind direction varied for 242 all seasons, with almost all occurrence from each direction being less than 243 15%: in winter, most winds were from the the I quadrant followed by the 244 II quadrant; W (about 16%) prevailed in spring, followed by WSW (14%) 245 and WNW (13%); southerly winds prevailed in autumn, and winds from E 246 to W accounted for more than 75% of occurrence; in autumn, most winds 247 were from the III quadrant. 248

The occurrence of wind speed at the 4 buoy locations during 2014–2016 249 can also be observed from Fig. 6. For B1, the occurrence of wind with speed 250 less than  $5 \,\mathrm{m \, s^{-1}}$  was, respectively, about 85%, 88%, 90%, and 70% in the 4 251 seasons (winter, spring, summer and autumn); relatively large wind speeds 252 of over  $8 \text{ m s}^{-1}$  are mostly found in autumn (about 10%), followed by winter 253 (about 3%). For B2, around 86% of wind speed was less than  $5 \,\mathrm{m\,s^{-1}}$  in all 254 seasons except spring, when it was around 80%; the occurrence of wind speed 255 greater than  $8 \text{ m s}^{-1}$  was always less than 1% of the time in all seasons. At B3 256 wind speed was less than  $5 \,\mathrm{m \, s^{-1}} \, 80\%$  of the time, while wind speed greater 257 than  $8 \,\mathrm{m \, s^{-1}}$  also occurred relatively rarely (less than 2% for all seasons). 258 Among the 4 buoys wind speed was greatest at B4, where wind speeds were 259 over  $5 \,\mathrm{m \, s^{-1}}$  in winter and autumn about 45% of the time, and 20% of the 260

time in summer and spring; the occurrence of wind speed over  $8 \text{ m s}^{-1}$  was, respectively, about 10%, 3.6%, 3.8%, and 1.5% in winter, spring, summer and autumn.

#### 264 3.2. Wave climate and assessment of wave energy potential

In this section, wave state in terms of observed wave height and calculated 265 wave energy period is analyzed. Time series of significant wave height,  $H_s$ , 266 and wave energy period,  $T_e$ , at the six buoys considered in this study are 267 presented in Fig. 7. It can be seen from this figure that  $H_s$  is relatively low 268 for most of the period 2014–2016. At the 6 buoy stations,  $H_s$  was less than 269 1 m, most of the time, and was even less than 0.5 m at buoy stations B2, B3, 270 and B6. It is not surprising that  $H_s$  is small in this region, since, first of all, 271 water depth is very small, and, secondly, the buoys are located in bays that 272 are somewhat semi-closed, all of which are not favorable for the development 273 of wind waves. Among these six buoys,  $H_s$  at B4 was relatively higher than 274 that at the others, which is to be expected since the water is deepest here 275 while being closer to the open South China Sea. Relatively higher values 276 exceeding 2 m are found only in some rare cases due to relatively strong winds 277 caused by tropical cyclones passing through the region, see discussion above. 278 For the period 2014–2016, no extreme waves (e.g.,  $H_s > 10 \text{ m}$ ) were observed, 279 even during periods of influence from tropical cyclones passing through. 280

 $T_e$  at the buoys is also presented in Fig. 7. In inner Dapeng Bay, represented by B1 and B2,  $T_e$  mostly varies between 3s and 7s, with occasional values of more than 10s. Similarly, at B5, in Daya bay,  $T_e$  was also mostly in the range of 3–7s. Relatively higher  $T_e$  was found at B4, with most values between 4–7s. In Aozaixia Bay (inner Daya Bay) represented by B3,  $T_e$  was relatively small and mostly in the range 3-5 s, Similar values of  $T_e$  are found in Shenzhen Bay represented by B6. Moreover,  $T_e$  values greater than 10 s occasionally occurred at the 6 buoys, mostly due to the influence of tropical cyclones.

Based on  $H_s$  and  $T_e$ , the wave power, P, was estimated at the 6 buoy 290 stations, Fig. 8. From comparison with time series of  $H_s$  shown in Fig. 7, the 291 temporal variations of P and  $H_s$  were basically coherent, at both hourly, daily 292 and monthly time scales. This can be explained from the relation of P and 293  $H_s$ , since P is proportional to the square of  $H_s$ . For the period 2014-2016, 294 P was mostly confined to the range of  $10^2$  to  $10^3$  W m<sup>-1</sup> at B1, B4 and B5. 295 while at B2, B3 and B6 wave power was considerably less, with values mostly 296 less than  $100 \mathrm{W m}^{-1}$ . 297

Statistics of annual wave climate at the six buoys were calculated and 298 are summarized in Table 1, in terms of annual mean significant wave height 299 and its standard deviation  $((H_s)_{mean} \pm \text{std.dev.})$ , maximum significant wave 300 height  $((H_s)_{max})$ , annual mean wave period  $((T_e)_{mean})$ , and annual mean 301 and maximum wave power (i.e.,  $P_{mean}$  and  $P_{max}$ ). It was not surprising that 302  $(H_s)_{mean}$  averaged over the period 2014–2016 was very small, with values less 303 than 0.5 m for the buoys except B4, and  $(H_s)_{mean}$  was only about 0.1 m at B3 304 and B6. B4 displays a relatively  $(H_s)_{mean}$ , but was still very small (0.6 m). 305 The largest  $(H_s)_{max}$  was found at B4, with value of over 4.0 m, occurring on 306 September 16, 2014 when Typhoon Kalmaegi (No. 1415) passed through. 307 For Inner Dapeng Bay, represented by B1 and B2,  $(H_s)_{max}$  was, respectively, 308 2.5 m and 2.4 m, and about 1.7 m at B5. Smallest  $(H_s)_{max}$  was still found at 309 B3 and B6 (less than 1 m).  $(T_e)_{mean}$  was relatively larger at B1, B2, B4 and 310

B5 (about 4.5 s), and smaller at B3 and B6 (3.5 s). The spatial distribution of wave power was similar to that of  $H_s$ : annual mean and maximum wave power,  $P_{mean}$  and  $P_{max}$ , were also largest at location B4 (1.25 kW m<sup>-1</sup> and 88.1 kW m<sup>-1</sup>, respectively), followed by B5, B1, and B2 (respectively, 0.46, 0.39, 0.26 kW m<sup>-1</sup> for  $P_{mean}$ ), while they were still smallest at B3 and B6 (0.03 kW m<sup>-1</sup> for  $P_{mean}$ , and 1.2 and 1.6 kW m<sup>-1</sup> for  $P_{max}$ , respectively).

Monthly and seasonal wave climate variability are of importance for wave energy resource assessment. Monthly mean and seasonal characteristics of  $H_s$  and  $T_e$  as well as P were thus investigated for the study region, based on the buoy observations over the period 2014-2016.

Fig. 9 shows the variability in monthly values of  $H_s$ ,  $T_e$  and P observed at 321 the 6 buoy locations, while also displaying spatial difference in these variables 322 between buoys. The monthly mean  $H_s$  for 2014–2016 can be briefly divided 323 into three groups: smallest monthly mean  $H_s$  was found at B6 and B3, with 324 values mostly less than 0.1 m; monthly mean values of  $H_s$  at B1, B2 and B5 325 were mostly in the range of 0.2 m - 0.5 m; largest monthly values  $H_s$  occurred 326 at B5, which was between 0.5 m - 0.8 m. Monthly mean values of  $T_e$  were also 327 smallest at B3 and B6 (mostly around  $3.5 \,\mathrm{s}$ ), followed by  $T_e$  at B5 (between 328 4s-5s), while monthly  $T_e$  at B4, B2 and B1 were relatively large with values 329 ranging between  $4.5 \,\mathrm{s} - 6.0 \,\mathrm{s}$ . Correspondingly, monthly mean wave power, 330 P, for the period 2014-2016 at the 6 locations can also be also divided into 331 three groups by magnitude: largest at B4, with values ranging from about 332  $1 \,\mathrm{kW \, m^{-1}}$  to  $3.5 \,\mathrm{kW \, m^{-1}}$ ; monthly P were all less than  $1 \,\mathrm{kW \, m^{-1}}$  at B1, B2 333 and B5; while very small values were recorded at B3 and B6, with magnitudes 334 of all less than  $0.1 \,\mathrm{kW}\,\mathrm{m}^{-1}$ . 335

Seasonal values of  $H_s$ ,  $T_e$  and P also displayed spatial variability be-336 tween the six buoy locations (see Fig. 9). At locations B3 and B6, seasonal 337 variations in  $H_s$ ,  $T_e$  and P were very small and limited in 0.1 m, 4 s and 338  $0.1 \,\mathrm{kW \, m^{-1}}$ , respectively. Seasonal differences in  $H_s$ ,  $T_e$  and P were evident 339 at locations B1 and B5: seasonal values at B1 and B5 were relatively large 340 in autumn and winter, and smaller in spring and summer months. Relative 341 to B1 and B5, reversed seasonality was observed at B2, i.e., large values oc-342 curred in summer months with smaller values observed in winter. Largest 343 seasonal values of  $H_s$  (around 0.6 m) were observed at B4, with relatively 344 small variability. However, larger values of  $T_e$  and P were recorded in sum-345 mer and autumn months, with smaller values in winter and spring. From the 346 comparison between seasonal and monthly winds (Fig. 5) and waves (Fig. 9), 347 it is not surprising that the variability in the wind climate were not coherent 348 with those of the wave climate in the study region, since, in general, waves 349 in the coastal area might be not generated by local winds. 350

In addition to the numerical values of significant wave height and energy 351 period, their frequency of occurrence is also important for assessment of 352 wave energy resources. The combined scatter and energy diagrams, in terms 353 of  $H_s$ ,  $T_e$ , and P, can provide convenient and comprehensive information 354 for conveying the characteristics of wave energy resources. Fig. 10 shows the 355 diagrams for the 6 sites, averaged at the same observational time (hourly) and 356 all based on the data 2014-2016. In the figure, the significant wave height 357 has been divided into intervals of one third of a meter in the range of 0-3 m, 358 and the energy period has been divided into intervals of 1s ranging from 1s 359 to 10 s. The colors on the diagrams show the proportion of incident energy 360

expected in one year, with numerical values given by the colour bar. The black curves, n the values of  $1 \text{ kW m}^{-1}$ ,  $3 \text{ kW m}^{-1}$ , and  $5 \text{ kW m}^{-1}$ , represent isolines of wave power calculated from Eq. 2. The numerical values on the diagrams in the figure represent the occurrence of a combination of  $H_s$  and  $T_e$  within the corresponding range, in number of hours per year.

In Inner Dapeng bay, represented by buoys B1 and B2, Fig. 10 shows 366 that, for buoy B1, sea states in the range of 0.3-0.6 m for  $H_s$  and 5-6 s 367 for  $T_e$  occurred most frequently, providing the largest contribution to the 368 total annual wave energy (more than 30%), and the second largest contri-369 bution (about 18%) was from sea states with  $H_s$  between 0.3–0.6 m and  $T_e$ 370 between 4-5 s, which also displayed high frequency of occurrence; at B1, the 371 frequency of occurrence of sea states with  $H_s$  between 0 m and 0.3 m and 372  $T_e$  between 4s and 6s was also very high, while the contribution to the to-373 tal wave energy was less than 20%, since their values are relatively small; 374 moreover, sea states of  $H_s$  between 0.3 m and 0.6 m and  $T_e$  between 6 s and 375 7s at B1 provided a relatively high contribution of about 15% to the wave 376 energy, due to the relatively larger values in terms of  $H_s$  and  $T_e$ , even though 377 their frequency of occurrence was relatively low; for buoy B2, sea states with 378 highest frequency of occurrence were, respectively, in the range  $0.0-0.3 \,\mathrm{m}$ 379 for  $H_s$  and 4-6 s for  $T_e$ , and together they contributed more than 30% of 380 total wave energy resources, while sea states of  $H_s$  between 0.3-0.6 m and 381  $T_e$  between 5–6s provided a significant contribution (more than 20%) to the 382 wave energy resource, followed by the contribution from sea states of  $H_s$  and 383  $T_e$ , respectively, in the range of 0.3-0.6 m and 6-7 s. 384

Concerning sea states, Fig. 10 clearly shows that for B3, located in Aoza-

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ixia Bay, within Daya bay, almost all sea states were below 0.3 m and in 386 the range 6-7 s, in terms of  $H_s$  and  $T_e$ , respectively, and contributing more 387 than 60% of total wave energy resources there, with the second most signifi-388 cant contribution coming from sea states of  $H_s$  less than 0.3 m and  $T_e$  4 s to 389 5s. For the other buoy station in Daya Bay, B5, sea states with the high-390 est frequency of occurrence were in the range 0.0-0.3 m and 4-5 s for  $H_s$ 391 and  $T_e$ , respectively, providing the largest contribution to the total annual 392 wave energy (more than 43%); with the second largest contribution (about 393 30%) coming from sea states with  $H_s$  and  $T_e$  between 0–0.3 m and 5–6 s, 394 respectively; the frequency of occurrence of sea states with  $H_s$  below 0.3 m 395 were also significantly high, but their contribution to the total annual energy 396 resources was relatively small. 397

For buoy B4, located off the coast of Dapeng peninsula and in the deepest 398 location of the 6 buoys, sea states in the range of 0.3-0.9 m for  $H_s$  and 4-6 s 399 for  $T_e$  provided the largest contribution to the total annual wave energy (all 400 together more than 70%), while sea states with larger  $H_s$ , between 0.9-1.2 m, 401 and  $T_e$  between 5–6s contributed about 8% of total wave energy resources. 402 For B6, located in the narrow Shenzhen Bay, sea states in terms of  $H_s$  and 403  $T_e$ , and their contributions to total wave energy, were similar to those at 404 location B3, where water depth is also very shallow as well. 405

Overall, for all buoys, most sea states in terms of  $H_s$  were below 0.6 m,. Concerning  $T_e$ , most sea state values were between 4–6 s for B1–B2, and B4–B5, while values of between 3–4 s were found for B3 and B6. A basic knowledge of significant wave height values informs that values below 0.6 m in the ocean are quite small. It is not surprising that  $H_s$  is small in the study region, since water depths are correspondingly low and the region is
relatively closed off from open seas.

Wave direction plays an important role in wave energy assessment. Fig. 11 413 shows seasonal distributions of incoming wave direction for the 6 buoys in 414 the study region. It is apparent from this figure that there was no obvious 415 seasonal change in wave direction for the 6 buoys. Wave directions at these 416 buoys locations were not in accordance with the corresponding wind direc-417 tions (see Fig. 6), but they mainly reflected the wave propagating directions 418 of propagating away from generation sources in the open region towards the 419 coast, so that none at all was from the IV quadrant due to the coastline 420 orientations at the buoys. Moreover, in all directions wave power was mostly 421 less than  $1 \,\mathrm{kW}\,\mathrm{m}^{-1}$ . 422

In Dapeng bay, most waves came from the II quadrant: for B1, the pre-423 vailing wave direction was from the SE, with a very minor contribution from 424 the SSE; for B2, southerly waves prevail in all seasons except in Winter, fol-425 lowed by SSW, whereas SSW waves a little more occurred in winter, with a 426 little lower occurrence (about 31%) from the south. For Dava bay, at B3, 427 most of the wave energy was provided by waves from the I quadrant, and 428 the prevailing wave direction was NE, followed by ENE and NNE, which did 429 not match the prevailing wind direction (see Fig. 6). This indicates that 430 the waves at B3 were mostly not generated by local winds, but mostly came 431 from the open region where the waves propagated to the coast. As for the 432 other buoy in Daya bay, B5, most of the wave energy was contributed by 433 easterly waves: ESE waves prevailed in winter and autumn, followed by N, 434 whereas northerly waves occurred more in Spring and summer, followed by 435

ESE: a minor contribution (less than 5%) was due to waves from the ENE. 436 At B4, Fig. 6 reveals that most waves came from the II quadrant, which 437 did not match the prevailing wind directions, indicating that the waves were 438 not generated by local winds, but from waves propagating away from gen-439 eration sources in the open region. For B6 located in Shenzhen Bay, most 440 of the waves (more than 50% on average for the all seasons) came from III, 441 as expected, with southwesterly and WSW directions prevailing. However, 442 it is noted that occurrence from other directions was rare, especially from 443 the opposite direction, i.e., NE and ENE, which might be due to the fact 444 that this is a relatively narrow area, and where the influence of reflected and 445 refracted waves might be of more significance. 446

Previous studies have shown that waves transport energy supplied to them over vast distances, and dissipative effects may play only a smaller role in deep water, as opposed to the surf zone; nearshore waves are nonlinearly related to the strength, fetch, and duration of the wind [62]. Therefore, the local wave climate can be frequently affected by strong incident waves or wind fetch both inside and outside the study region.

Last but not least, we provide a brief discussion concerning the relation-453 ship between the wind energy and the wave climate for the study area. Fig. 12 454 displays a preliminary correlation between the wind speed/energy and the 455 wave climate based on the local buoy measurements. Results from all buoys 456 other than B5 (not recorded) are displayed. The upper panels of Fig. 12 show 457 the relationship between anomalies of monthly averaged data (monthly val-458 ues minus monthly climatology) of wind speed and significant wave height for 459 the period 2014–2016. For this long-timescale (low frequency) comparison, 460

the threshold correlation coefficient at the 95% and 99% significance levels 461 is about 0.35 and 0.45, respectively (the sample of the time series of the 462 monthly anomalies is about 30). It can be seen from this figure that, except 463 at B3 buoy location, all other locations display a relatively small correla-464 tion between wind strength and wave conditions. Moreover, we attempted 465 to calculate the delay/forward correlations, but they are still insignificant 466 (not shown). The bottom panel in Fig. 12 describes correlations between 467 daily anomalies (daily values minus monthly climatology) of wind speed and 468  $H_s$  for the period 2014–2016. The coefficients were all statistically signifi-469 cant (sample numbers for the time series are all greater than 550, and the 470 threshold correlation coefficient at the 95% and 99% significance levels is less 471 than 0.11). At Dapeng Bay, represented by buoys B1 and B2, correlations 472 were over 0.3 for the period 2014-2016. At B3, located in Daya Bay, the 473 correlation was very high, 0.77. A relatively high correlation is also found at 474 station B4 (0.60). For Shenzhen Bay (B6), the correlation coefficient is about 475 0.39, similar to values for Dapeng Bay. Therefore, at this daily timescale, the 476 local wind energy significantly influences the wave climate in the study area. 477 Thus, care must be taken before we can conclude whether the wave climate in 478 the study region might be more affected by the local wind at daily timescale 479 rather than for long-time statistics when non-local wave signals from outside 480 the study region get more involved. 481

#### 482 4. Summary and Conclusions

In this study, wind and wave climates for the Shenzhen coastal region are evaluated by means of buoy observational data. Buoys were first placed in the region by the city local government in 2014 to observe and monitor the atmospheric and hydrodynamic climate of the region. Six buoys are located in different areas of the study region, including Dapeng Bay, Daya Bay, Shenzhen Bay, and the area off the tip of the Dapeng peninsula. The waters in these areas are very shallow, ranging in depth from about 3 m-22 m.

In terms of wind speed and direction at the buoys (2.5 m above the sea)490 surface), wind climate and potential wind energy resources were assessed in 491 detail for the period 2014-2016. It was found that the annual mean wind 492 speed at the buoy locations for the period 2014-2016 varied from about 493  $3.1 \,\mathrm{m \, s^{-1}}$  to  $4.1 \,\mathrm{m \, s^{-1}}$ , with maximum wind speeds of more than  $17 \,\mathrm{m \, s^{-1}}$  oc-494 curring as a result of tropical cyclones. These winds resulted in annual mean 495 wind powers of about  $37-94 \mathrm{W m^{-2}}$ . Among the buoys, largest averaged 496 wind speed and power were found at B4, located in the relatively open area 497 off the southern coast of Dapeng peninsula. On average, more than 80% of 498 wind speeds were less than  $5 \,\mathrm{m \, s^{-1}}$  in the study region. However, the wind 499 speed was relatively large at location B4, where about 45% of wind speeds 500 were over  $5 \text{ m s}^{-1}$  in winter and autumn, and 20% in summer and spring. 501

Seasonal variability in wind speed and power fluctuated at the different buoy locations over the 2014–2016 period. At B1 and B4, seasonal variability was relatively large in autumn and winter, and smaller in spring and summer months. However, reversed seasonality occurred at location B6, where wind and wind power were relatively large in spring and summer, and were smaller in autumn and winter. Seasonal variations were realtively small at B2 and B3 locations.

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Seasonal and spatial wind direction variability differed between buoys

B1–B4 (no wind direction data at B5 and B6). At B1, located in the north-510 west end of Dapeng Bay, the prevailing wind direction in winter, spring, sum-511 mer and autumn was, respectively, easterly, easterly to southerly, southerly, 512 and southerly. The prevailing wind direction was from the east in all seasons 513 at location B2. At B3, located in Dava Bay, the largest contribution to wind 514 energy resources was provided by the southerly winds, in all seasons, with 515 SSE and SSW winds also providing considerable contributions. The prevail-516 ing wind direction at B4 varied seasonally, with occurrence of less than 15%517 from any particular direction: in winter, most winds were from the the I 518 quadrant, followed by the II quadrant; westerly winds prevailed in spring, 519 with southerly winds prevailing in autumn. 520

Wave climate and potential wave energy resources in terms of wave height, 521 wave energy period, wave direction, and wave power were also evaluated for 522 the period 2014–2016. The data showed that at the 6 buoy locations,  $H_s$  was 523 mostly less than 1 m, and even less than 0.5 m at B2, B3, and B6 locations. 524 As a result, wave power, P, was mostly limited to the range  $10^2$  to  $10^3$  W m<sup>-1</sup> 525 at B1, B4 and B5, and mostly less than  $100 \,\mathrm{W m^{-1}}$  at the other locations. 526 This may be due to the facts that, first, water depth is very shallow, and, 527 secondly, the buoys are located in bays where the sea is somewhat semi-528 enclosed, all of which are not favorable for the development of wind waves. 529  $T_e$  was mostly in the range of  $3\!-\!7\,\mathrm{s}$  in the study region, with values of more 530 than 10s occasionally occurring at all 6 buoys, mostly during periods of 531 tropical cyclones. 532

It was not surprising that annual mean significant wave height,  $(H_s)_{mean}$ , for the period 2014-2016 was relatively small, with values of less than 1.0 m

for the study region, and largest values of  $(H_s)$  found at B4 of over 4.0 m. 535 In addition, most sea states in terms of  $H_s$  were less than 0.6 m, with  $T_e$ 536 between 4-6s for B1-B2, and B4-B5, and 3-4s for B3 and B6. The 537 annual mean wave energy period,  $(T_e)_{mean}$ , was relatively large at B1, B2, 538 B4 and B5 (about  $4.5 \,\mathrm{s}$ ), and smaller at B3 and B6 ( $3.5 \,\mathrm{s}$ ). Correspondingly, 539 the annual mean wave power was largest at B4  $(1.25 \,\mathrm{kW \, m^{-1}})$ , and between 540 0.26-0.46 kW m<sup>-1</sup> at B1, B2, and B5, and smallest at B3 and B6, with values 541 of only  $0.03 \,\mathrm{kW}\,\mathrm{m}^{-1}$ . Therefore, we can conclude that the potential of the 542 wave energy resource at the buoy locations are very small. 543

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- <sup>712</sup> [62] Zhang Y, Kennedy AB, Tomiczek T, Donahue AS, Westerink JJ. Val<sup>713</sup> idation of Boussinesq-Green-Naghdi Modelling for Surf Zone Hydrody<sup>714</sup> namics. Ocean Engineering 2016;111:299–309.

Table 1: List of measurement buoys in the Shenzhen coastal region with water depth and some fundamental mean wind and wave characteristics. The collected buoy data covered from April 2014 to the end of 2016, with data gap that can be seen in Fig. 2. Main wind and wave characteristics were statistically calculated, including wind speed,  $V \text{ (m s}^{-1})$ ; wind power,  $W \text{ (W m}^{-2})$ ; significant wave height,  $H_s$  (m); wave energy period,  $T_e$  (s); wave power,  $P \text{ (kW m}^{-1})$ .

Buoy	Depth (m)	$V_{mean}$	$V_{max}$	$W_{mean}$	$W_{max}$	$(H_s)_{mean}$	$(H_s)_{max}$	$(T_e)_{mean}$	$P_{mean}$	$P_{max}$
		$\pm$ std.dev.				$\pm$ std.dev.				
B1	11	$3.1\pm2.3$	17.5	58	3282	$0.34 \pm 0.18$	2.5	4.7	0.39	22.5
B2	11	$3.1 \pm 1.7$	17.6	37	3339	$0.25 \pm 0.17$	2.4	4.7	0.26	28.0
В3	5	$3.4 \pm 2.0$	15.6	52	2325	$0.11 \pm 0.05$	0.8	3.5	0.03	1.2
B4	22	$4.1\pm2.5$	17.1	94	3063	$0.62\pm0.30$	4.3	4.5	1.25	88.1
B5	12	/	/	/	/	$0.38 \pm 0.23$	1.7	4.3	0.46	8.7
B6	3	$3.6 \pm 1.8$	15.7	51	2347	$0.13\pm0.05$	0.9	3.5	0.03	1.6



Figure 1: (C): Bathymetry contours (m) for the study area together with locations of the six wind and wave measurement buoys (B1-B6) in the coastal Shenzhen region (bot-tom/main panel). The study area is situated in the northern South China Sea (top panels). The broken line represents the boundary of Hong Kong waters.



Figure 2: Time series of wind speed  $(m s^{-1})$  at the buoy locations (B1-B4 and B6) shown in Fig. 1 for the period 2014–2016. Dashed lines show the time periods of the tropical cyclones (see Fig. 4) passing through the study region. Wind data were missed at B5 buoy.



Figure 3: Time series of wind power  $(W m^{-2})$  at locations B1–B4 and B6 shown in Fig. 1 for the period 2014–2016.



Figure 4: Left: Paths of tropical cyclones (TCs) passing through the northern South China Sea for the period 2014-2016. Right: The TC No. (Name) and landed time. Data source: http://tcdata.typhoon.gov.cn/zjljsjj\_sm.html [61]. The classifications of the TCs in China are as follows: Tropical Depression (TD;  $10.8-17.1 \text{ m s}^{-1}$ ); Tropical Storm (TS;  $17.2-24.4 \text{ m s}^{-1}$ ); Severe Tropical Storm (STS;  $24.5-32.6 \text{ m s}^{-1}$ ); TYphoon (TY;  $32.7-41.4 \text{ m s}^{-1}$ ); Severe TYphoon (STY;  $41.5-50.9 \text{ m s}^{-1}$ ); and Super TYphoon (SuperTY;  $\geq 51.0 \text{ m s}^{-1}$ )



Figure 5: Seasonal (left) and monthly (right) mean wind speed  $(m s^{-1})$  and wind power  $(W m^{-2})$  for the buoys B1-B4 and B6 shown in Fig. 1.



Figure 6: Seasonal wind speed and direction roses based on the measurement from buoys B1-B4 for the period 2014-2016. Wind direction data were missed at B5 and B6 locations.



Figure 7: Time series of significant wave heights  $(H_s \text{ in } m)$  and wave energy periods  $(T_e \text{ in } s)$  at the six buoy locations (B1-B6) for the period 2014-2016.



Figure 8: Time series of wave power  $(Wm^{-1})$ , P, at buoy locations of B1-B6 shown in Fig. 1 for the period 2014-2016.



Figure 9: Seasonal (left) and monthly (right) mean significant wave height  $(H_s)$ , energy period  $(T_e)$ , and wave power (P), for the six buoys (B1-B6) shown in Fig. 1 for the period 2014-2016.



Figure 10: Bivariate distributions of occurrence and energy in terms of significant wave height  $(H_s)$ , and energy period  $(T_e)$  averaged for the period 2014-2016 for at the six stations displayed in Fig. 1. The color scale, as a percentage, represents the contribution of the sea state to the total energy, while the black numbers indicates the occurrence of sea states in number of hours in one year.



Figure 11: Seasonal wave power and direction roses based on the measurement from the buoys B1-B6 for the period 2014-2016.



Figure 12: Top panel: the relationship between anomalies of monthly averaged data (monthly values minus monthly climatology) of wind speed and significant wave height,  $H_s$ , at buoys B1-B4 and B6 for the period 2014-2016. Bottom panel: the relationship between anomalies of daily averaged (daily values minus monthly climatology) of wind speed and Hs at the buoys B1-B4 and B6 (as above). R represents correlation coefficient.