

# **Biovolatilization of Arsenic as Arsines from Seawater**

Savage, L., Carey, M., Williams, P. N., & Meharg, A. A. (2018). Biovolatilization of Arsenic as Arsines from Seawater. *Environmental science & technology*, *52*(7), 3968. Advance online publication. https://doi.org/10.1021/acs.est.7b06456

Published in: Environmental science & technology

**Document Version:** Peer reviewed version

**Queen's University Belfast - Research Portal:** Link to publication record in Queen's University Belfast Research Portal

#### Publisher rights

© 2018 American Chemical Society. This work is made available online in accordance with the publisher's policies. Please refer to any applicable terms of use of the publisher.

#### General rights

Copyright for the publications made accessible via the Queen's University Belfast Research Portal is retained by the author(s) and / or other copyright owners and it is a condition of accessing these publications that users recognise and abide by the legal requirements associated with these rights.

#### Take down policy

The Research Portal is Queen's institutional repository that provides access to Queen's research output. Every effort has been made to ensure that content in the Research Portal does not infringe any person's rights, or applicable UK laws. If you discover content in the Research Portal that you believe breaches copyright or violates any law, please contact openaccess@qub.ac.uk.

#### **Open Access**

This research has been made openly available by Queen's academics and its Open Research team. We would love to hear how access to this research benefits you. – Share your feedback with us: http://go.qub.ac.uk/oa-feedback

## 1 Biovolatilization of arsenic as arsines from seawater

2	Laurie Savage <sup>+</sup> , Manus Carey <sup>+,</sup> Paul N. Williams <sup>+</sup> , Andrew A. Meharg <sup>*,+</sup>
4	<sup>+</sup> Institute for Global Food Security, Queen's University Belfast, David Keir Building, Malone
5 6	Road, Belfast, BT9 5BN, Northern Ireland.
7 8	* Corresponding author.
9	
10	
11	
12	
13	
14	

## 16 ABSTRACT

17	Marine sources of arsenic to the atmosphere are normally dismissed as minor. Here we
18	show that arsenic can be biovolatilized from seawater, and that biovolatilzation is based on
19	organic arsenic species present in the seawater. Even though inorganic arsenic is in great
20	excess in seawaters, it is trimethylarsine (TMA) that is the primary biovolatilized product,
21	with dimethylarsine (DMA) also observed if dimethylarsinic acid (DMAA) is spiked into
22	seawaters. With respect to budgets, 0.04% of the total arsenic in the seawater was
23	biovolatilized over a 2-week incubation period. To test the environmental significance of this
24	finding, wet deposition was analyzed for arsenic species at coastal locations, one of which
25	was the origin of the seawater. It was found that the oxidized product of TMA,
26	trimethylarsine oxide (TMAO), and to a less extent DMAA, were widely present. When
27	outputs for arsines (0.9 nmol/m <sup>2</sup> /d) from seawater and inputs from wet deposition (0.3-0.5
28	nmol/m <sup>2</sup> /d) were compared, they were of the same order of magnitude. These findings
29	provide impetus to reexamining the global arsenic cycle as there is now a need to determine
30	the flux of arsines from the ocean to the atmosphere.
31	

### 34 • INTRODUCTION

35 The biogeochemical cycle of arsenic, particularly the fluxes through the atmosphere, is poorly described (1, 2). These fluxes are thought to be biovolatilization, salt spray 36 37 entrainment, entrainment of soil dusts, and industrial emissions (1). The uncertainty regarding inputs of arsenic into the atmosphere is due to limited actual field measurements, 38 39 using outdated analytical approaches, with limited arsenic speciation conducted (1). The 40 observation that trimethylarsine oxide (TMAO) constitutes ~20% of arsenic in wet 41 deposition, from weather whose back-trajectories are marine, raised the hypothesis that 42 TMAO in the atmosphere could be of marine origin (3). TMAO is the oxidized form of trimethylarsine (TMA), the main form of arsenic biovolatilized by fungi and bacteria (4-6). 43 44 However, a previous observation that TMAO was bound to marine sampled atmospheric 45 particulates, sampled on a remote Japanese island in the S. China Sea, was attributed to 46 evolution of arsines from (mainland Japanese) soil (7). Trace levels of monomethylarsonic 47 acid (MMAA), dimethylarsinic acid (DMAA) and TMAO have been detected on atmospheric 48 particulates collected on Atlantic and Pacific cruises, at similar concentrations for each 49 ocean, circa. 10 pg/m<sup>3</sup>, 50-500-fold lower than inorganic arsenic (8). There was not a clear 50 pattern in MMAA, DMAA and TMAO concentrations, which contrasts with a later 51 Argentinian terrestrial study that found TMAO at concentrations in particulates at 60 pg/m<sup>3</sup>, 52 with DMAA 5-fold, and MMAA 10-fold, lower (9). Soils have been widely shown to produce 53 arsines, with TMA being the dominant species produced (5,6). This soil source hypothesis 54 for TMAO in marine sampled particulates (7), though, ignored the fact that Japan, and other 55 sampling locations (8,9), are surrounded by, or are beside, oceans, and that the prevailing 56 weather trajectories originate from the oceans.

58	The literature on marine production of arsines is scant, though marine production of arsines
59	is predicted to be important (2). There is only one study on marine algal isolates, showing
60	that they could biovolatilize arsenic, though speciation of these arsines was not ascertained
61	(4). Wet deposition studies that speciate arsenic are limited to one seaboard (3) and one
62	continental (10) setting; with the continental low, and the seaboard one high, relatively, in
63	TMAO. The metabolic precursors of TMA/TMAO have been found widely in seawaters,
64	MMAA and DMAA (11, 12). TMAO can also be the product of arsenobetaine (AsB)
65	degradation (13), a dominant arsenic species in marine biomass. Assessing arsenic
66	speciation in seawater is difficult as the high salt content interferes with chromatography,
67	and the chromatography is often not well resolved with key species co-eluting; such as for
68	anion exchange chromatography where AsB, arsenite and TMAO can co-elute (3). Also,
69	given that there are limited analytical standards in circulation for TMAO, it is not widely
70	reported in the literature (3). In a recent study where multiple marine water certified
71	reference materials (CRMs) were speciated by arsine generation, with those species cryo-
72	trapped and the cryo-trap interfaced with inductively coupled plasma – mass spectrometry
73	(ICP-MS), TMAO was quantified at $\sim$ 0.2 nM concentrations (14). As well as seawater,
74	TMAO/TMA can emanate from soils (15), and can have a volcanic/geothermal origin (16),
75	though the contribution of soils and geothermal sources to the wider environment is
76	relatively low (3, 15,16).
77	

To test the hypothesis of Savage et al. (3) that TMAO in wet deposition is primarily from
marine origin, natural seawaters have to be shown to produce arsines, an experiment that
has not been conducted, to date. In this study, we trap arsines emanating from seawater

- 81 under both un-amended and DMAA amended conditions. The DMAA addition will show if
- 82 arsines production is methylated substrate availability inhibited, and whether any DMAA is
- 83 converted to TMA, to ascertain if DMMA is a precursor to biovolatilization. Furthermore, we
- 84 establish organic arsenic species in wet deposition at coastal locations, rather than the
- 85 inland sites used in previous wet deposition speciation studies (3, 10).
- 86

### 87 • MATERIALS AND METHODS

#### 88 Generation of arsines from seawater

Fresh seawater was obtained from Belfast Lough, Irish Sea, 54.644° N, 5.812° W (Figure 1) 89 90 during September 2017. The seawater was transported directly back to the laboratory (30 91 minutes from sampling site) and immediately transferred into 1 L Quickfit reaction vessels 92 fitted with a Drechsel head with a sintered glass bubbler, with the bubbler positioned below 93 the water line. Connected to the reaction chamber inlets and outlets, were silica gel 94 sampling tubes, dosed with silver nitrate according to Mestrot et al. (15). A personnel gas 95 sampling pump (Pocket Pump, SKC, Dorset, UK) was then joined to the outlet tube, with air drawn into the chamber under negative pressure at a rate of 200 mls/min. Eight such 96 97 microcosms were set up and placed in an Memmert AtmoCONTROL incubator, set at a 98 temperature of 20°C, and lit with LED lights at an intensity of 8,000 lux. After 1-week the 99 inlet and outlet tubes were replaced, and DMAA added to four of the microcosms at a 100 dosing of 1 nM. Chemical and instrument suppliers, and analytical instruments suppliers and 101 conditions, and are given for all analysis in Table S1. This concentration of DMAA (1nM) is 102 near the upper limit of what we recorded in Irish wet deposition as outlined in this study. 103 The microcosms were then incubated for a further week, with traps removed and analyzed.

104

Arsenic species trapped on the silver nitrate impregnated silver gel were eluted with ~ 5ml boiling water, volume accurately measured, for subsequent chromatographic quantification of species (15). Arsenic species in the seawater were determined in triplicate following 10fold dilution and oxidation of any arsenite to arsenate with 0.3% final concentration of H<sub>2</sub>O<sub>2</sub>, to move arsenite away from the co-eluting species, AsB and MMAA. For speciation, a

110 Thermo Scientific IC5000 Ion Chromatography system interfaced with a Thermo ICAP-Q ICP-111 MS was used. Chromatography utilized a 2x250mm Thermo Scientific AS7 anion-exchange 112 column with a 2x50mm AG7 guard column and ammonium carbonate over a linear gradient 113 from 20nM to 200nM was the mobile phase. Purchased standards, suppliers detailed in 114 Table S1, of arsenobetaine, inorganic arsenic, MMAA, DMAA, and TMAO were run to 115 ascertain retention times, and for quantification of those species. Total arsenic was 116 determined in triplicate seawater samples following Millipore membrane (0.45 micron) 117 filtering on 10-fold diluted samples.

118

For total arsenic determination, the Thermo Scientific ICAP-Q ICP-MS was used in helium collision cell mode to quantify arsenic against authentic samples with rhodium used as an external standard, spiked to give a 10  $\mu$ g/l concentration. For all analytical methods, blanks were prepared from >18.2 MΩ.cm (at 25°C) water in the same manner as the samples and included with each instrument batch. Limit of Detection (LoD) was calculated as average concentration of the blank + 3 times standard deviation of a species mix dilution series.

125

126 Arsenic speciation in wet deposition

Wet deposition was collected during July and August 2015 at two sampling locations, the
Belfast location used for seawater collection, and an Atlantic coastal location, Caherdaniel,
Co. Kerry, Republic of Ireland, 51.769° N, 10.099° W. Bulk deposition was collected in 56 mm
mouth diameter, pre-weighed, polypropylene 30 ml sampling tubes, suspended 1.5 m off
the ground to avoid splash back, for 24-h and immediately frozen on collection. Volume of
wet deposition collected was measured gravimetrically, and then converted to a deposition
volume on a surface area basis. Defrosted samples were 0.45 micron filtered, H<sub>2</sub>O<sub>2</sub> added

- to 1%, and arsenic species quantified by IC-ICP-MS, as detailed in Savage et al. (3). The 72-h
- 135 air trajectories for each location at each date was modelled by Hysplit (17) as per Savage et
- 136 al. (3).
- 137
- 138

### 139 • **RESULTS**

#### 140 **Biovolatilization of arsenic from seawater**

141 The total arsenic present in the fresh seawater was 26.0 ± 0.6 nM. Speciation was quantified 142 using anion-exchange in a 1/10 seawater dilution. The sum of species was 18.5 nM, with 7.5 143 nM difference in total arsenic being due to arsenic species retained on column or peak 144 broadened below LoD due to salinity. Speciated compounds were inorganic arsenic at 16.5 ± 145 0.7 nM, and DMAA at 1.4  $\pm$  0.7 nM. A peak was observed where TMAO/AsB eluted at 0.59  $\pm$ 146 1.1 nM, but was not apparent on the cation exchange chromatographic runs. TMAO has 147 severe peak broadening on the cationic column in the presence of seawater and is not quantifiable. AsB is not peak broadened on the cationic column, suggesting the unidentified 148 149 peak is TMAO. The AsB/TMAO peak merged into the DMAA peak at lower sea water 150 dilutions (X1 and X2) and was below LoD (0.08 nM) at higher dilutions (X100). No MMAA 151 above LoD was present in any of the dilutions.

152

153 Only 3 arsenic species, TMAO, DMAA and inorganic arsenic were present on seawater 154 microcosm outlet chemo-traps; while only two were detectable on inlets, DMAA and 155 inorganic arsenic (Figure 2). The ½ Limit of Detection (LoD) for TMAO was calculated at 64 156 pM in solution extract; the trap was diluted to ~ 5ml. The inlet concentration was subtracted 157 from the outlet concentration and the data expressed on the quantity of arsines 158 biovolatilized from 1 L of seawater over one week (Figure 3). During the first week of 159 sampling, AsH<sub>3</sub> was produced/trapped at a rate of 2.9 pmol/L/wk, which was significant 160 based on one-way anova analysis at P= 0.05, but fluxes were negligible for week 2, 161 regardless of DMAA treatment. DMA biovolatilization was negligible for weeks 1 and 2 162 without DMAA addition, but production was 15.8 pmol/L/wk for the DMAA treatment,

significant at P<0.0001. TMA was produced at a rate of 3.0 pmol/L/wk for the first week 163 164 incubation, and increased to 8.3 pmol/L/wk in the second week without DMAA addition, 165 and 14.7 pmol/L/wk with DMAA addition, P=0.013 for the one-way anova for the 3-way 166 comparisons, with the only pairwise comparison that was significant being week 1 versus 167 week 2 with the DMA amendment (P=0.0124). The sum of ( $\Sigma$ ) arsenic species increased from 5.1 to 9.0 pmol/L/wk in the absence of additional DMAA in week 2, and with DMAA 168 169 this figure was 32.2 pmol/L/wk, with one-way-anova being highly significant (P=0.0008). The 170  $\Sigma$  arsenic species biovolatilized was 0.01% of total seawater arsenic for week one and 0.03% 171 for week two without DMAA. The additional  $\Sigma$  arsenic species biovolatilized in the week 2 172 DMAA treated minus the non-DMAA treated microcosms, expressed as a percentage of the 173 DMAA added, was 23.5%.

174

## 175 Arsenic speciation in wet deposition

176 Two sampling sites were used to study arsenic inputs in wet deposition (Figure 1). The Kerry site, in the S.W. of Ireland on the Atlantic coast, is one of the most remote and unpolluted 177 178 regions of Europe, dominated with Atlantic originating weather trajectories delivering wet 179 deposition (Figure 1). Belfast is situated on the Irish Sea, and while still remote in a 180 European context, has a history of industrialization and is closer to Europe, but still receives 181 predominantly wet deposition from Atlantic originating trajectories (Figure 1). The Kerry 182 location had higher concentrations of TMAO and DMAA throughout, and at lower 183 depositional volumes, also higher inorganic arsenic concentrations (Figure 4 and Table 1). 184 TMAO and inorganic arsenic were routinely above the LoD (LoDs given in Table 1) at deposition volumes below 5 L/m2/d, while DMAA had only one sample above LoD for 185 186 Belfast, but was above LoD for half the Kerry samples. DMAA was too low to observe any

187 clear washout, i.e. the relationship between arsenic species concentration and depositional 188 volume, but for TMAO at both sites, and also for inorganic arsenic at the Kerry site, this was 189 clear and steep, fitting exponential decay functions. Inorganic arsenic was more erratic, and 190 could be high at higher depositional volumes, for Belfast, with no clear exponential decay 191 relationship for washout.

192

193 Considering depositional inputs of the 3-arsenic species (Table 1). Inorganic arsenic inputs 194 were 2.6-fold higher for Belfast, though not significantly so (P>0.05). DMAA was similar for 195 both sites, but this is an estimate given that the Belfast sites were mainly below LoD, and 1/2 196 LoD was used in this calculation. TMAO was significantly (P<0.0001) higher, 1.6-fold, in Kerry 197 compared to Belfast. Comparing these wet depositional inputs with seawater 198 biovolatilization for TMA (Figure 3), assuming that biovolatilization is occurring to a depth of 199 1 M, and averaging TMA production for week 1 and 2, then 1 m<sup>2</sup> of seawater can produce 200  $0.86 \text{ nmol/m}^2/d$ . Wet depositional inputs of TMAO was at a rate of 0.31 and 0.50 201 nmol/m<sup>2</sup>/d for Belfast and Kerry respectively, i.e. just lower than potential seawater TMAO 202 production.

### 203 DISCUSSION

204 The results presented here show that arsines, particularly, TMA, can be biovolatilized from 205 seawaters at rates commensurate with inputs from wet deposition, suggesting very close 206 linking between the two, and explaining observations that TMAO deposition occurs when 207 weather trajectories are of marine origin (3). It also calls into question that TMAO observed 208 in marine particulates are of terrestrial origin, as has previously been hypothesized (7). The 209 fact that higher concentrations in wet deposition, and higher depositional rates of TMAO 210 were observed in a remote S.W. Irish location, as compared to an industrialized N.E. Irish 211 location, points to a natural origin of this TMA/TMAO for both sites, given that Belfast is also on an Irish Sea coastal location. That S.W. Irish rainwater concentrations of TMAO were 212 213 higher could be due to both Kerry sitting on a larger water body, the Atlantic, and due to the 214 fact that Belfast weather trajectories, as for Kerry, are predominantly westerly, but that 215 they also have to travel over the Irish landmass, for most trajectory pathways. The concentrations of  $\Sigma$  arsenic species recorded here relate well to previous 20<sup>th</sup> century 216 217 observations for total arsenic for Bermuda, Eastern Atlantic (18) and Western Pacific 218 locations (19), and to our own Indian Ocean measurements (3). These illustrate the ubiquity 219 of arsenic in maritime derived wet deposition.

220

TMAO is rapidly washed out of the atmosphere, and thus longer weather paths over land would decrease atmospheric burdens of TMAO when precipitation is occurring. This steep washout curve, and the ubiquity of TMAO in wet deposition samples, had been shown in our previous Indian Ocean investigations (3). For the Indian Ocean (3), TMAO inputs were lower in winter as compared to summer, even though weather trajectories come from the same locations. As organic arsenic species in the water column are seasonal in the Indian

Ocean, with DMA being higher during summer months (13), production and availability of
organo-arsenic species in the water column may be key to understanding seawater
biovolatilization of arsenic. Atlantic algal productivity has also been linked to total arsenic
deposition at a west Britain location, when air trajectories were so aligned (20).

231

232 Previous studies on arsenic speciation in Atlantic waters find that DMAA concentrations of 233 0.1-0.3 nM are typical, rapidly decreasing with depth (>50 m) to below LoD. (12,21,22). 234 Higher DMAA concentrations, similar to ours at ~1 nM, were reported from the continental 235 shelf off the American east coast (12) and multiple European west coasts continental shelf locations (23). This suggests that methylation of arsenic is more active in a shelf setting. In 236 237 deeper ocean surface-waters, the lower concentration of DMAA does not appear to be due 238 to limitation in total or inorganic arsenic, a metabolic precursor (4), as inorganic arsenic is 239 typically 10-20 nM for both open ocean and coastal shelf waters (12, 21-23). MMAA is less 240 commonly reported. It is mainly below LoD for a French coast study (23), as for our Irish Sea 241 location here. It is below LoD in other Atlantic studies (21,22), and averaging 0.13 nM in 242 another (12). It should be noted that only direct analysis of arsenic speciation was made 243 using the French study (23), while the other Atlantic studies used indirect assessment of 244 arsenic species (photo-ionization detection) (12,21,22). Using a highly sensitive and accurate 245 speciation approach, on seawater certified reference materials, Matousek et al. (14) 246 measured TMAO in coastal and off-shore waters at concentrations ranging from 0.16-0.3 247 nM, higher than MMAA (0.03-0.12nM), whereas DMAA ranged from 0.53-1.64 nM, and 248 inorganic arsenic was ~14 nM. Thus, when specifically looked for, i.e. using appropriate 249 analytical methodologies, TMAO appears to be present at lower concentrations than DMAA, 250 but higher than MMAA (14). The fact that TMAO is always higher than DMAA in wet

251 deposition, with DMAA often below LoD, yet DMAA concentrations are higher in seawaters, 252 argues for TMAO coming from a biovolatilization route rather than salt spray, as has been 253 previously postulated to be the source of arsenic in wet deposition (1). Also, MMAA and 254 TMAO appear to have similar concentrations in seawaters (14), yet MMAA is never recorded 255 in wet deposition, probably as it is below LoD, not just absent. MMAA <LoD, again, argues 256 for TMAO in wet deposition being from a biovolatilization, as opposed to a salt spray, 257 source. Low TMAO concentrations in the water column may be expected if TMAO is rapidly 258 converted to TMA and lost from the water column.

259

The only previous study on biovolatilization from a marine media was a micro-algal pure 260 261 culture study (4). Arsenate was added at 20,000 nM concentration to incubating 262 Ostreococcus tauri under optimal culture conditions, and 0.01% of this arsenic was 263 converted to arsines. Even though our total arsenic concentrations in seawater were a 1000 264 times lower, and that any biovolatilizing microbial community was present at natural 265 densities, we found similar conversion of dissolved seawater arsenic to arsines, 0.04% over 266 2-week incubation. While Zhang et al. (4) did not speciate trapped arsines, they did speciate 267 the incubating medium where they found 2.4% of added inorganic arsenic converted to 268 DMAA, but no MMAA was detected. TMAO and AsB were not looked for, or, at least, too 269 low to report. Given that TMAO and AsB can co-elute with other species (3), and seawater 270 interferes with TMAO detection, TMAO (if present) may not have been observed under the 271 chromatographic conditions used. In a study on freshwater photosynthetic cyanobacteria 272 where the arsenic methylating protein ArsM was isolated, it was shown to convert inorganic 273 arsenic to arsines in vitro, with 2% of arsenic in the incubating medium converted to TMA, 274 alone (24). TMAO, DMAA and MMMA were all detected in culture solution after addition of

275 inorganic arsenic, in both purified ArsM studies and where the ArsM gene was cloned into 276 *Escherichia coli*. When DMAA was added directly to our microcosms we saw a 23% 277 conversion to DMA. DMAA addition enhanced the production of TMA by 77%, showing 278 methylation and reduction of DMAA to biovolatilized TMA. As DMAA was detected in 279 seawaters at circa. the same levels as the amendment, it is likely that the exogenously 280 added DMAA was rapidly metabolized, and the natural DMAA in the seawater was within 281 cells and lysed on preparation for analysis. This is consistent with findings on algae where 282 the bulk of the arsenic in microcosms is within the cells (4,24). Another observation of our 283 study was that TMAO production from both non-amended seawater increased during the 284 second week, suggesting enhanced arsenic species turnover during week 2. The pathways 285 and regulation of arsenic methylation in natural seawaters needs further investigation. It is 286 becoming apparent that seawater contribution to the global arsenic cycle has been 287 underestimated, and marine biovolatilization as a source ignored by all but a handful of 288 studies (3,4). Matschullat (1) estimated that 27 t of arsenic in the global atmosphere was 289 derived from salt-spray. Although our measurements are limited in geography and time, 290 they are the only measurements of TMA output and TMAO input to be conducted to date. 291 There is now a need to estimate the contribution of marine produced arsines to arsenics 292 global arsenic cycle.

293

Wet deposition, reported here, and in Savage et al. (3), and measured seawater TMAO concentrations (14), are in the same range sub-nM. It also has to be considered that wet deposition direct to seawaters, or runoff from terrestrial wet deposition, will also be a source of TMAO to seawaters, and there might be tight recycling of TMAO depositional inputs to TMA outputs, which are then oxidized in the atmosphere to TMAO, before re-

299	deposition. However, demethylation of TMAO also has to be considered as a potential fate				
300	of TMAO in the water column as well (13). TMAO can also be derived from AsB degradation				
301	(13), and the role of this pathway also remains to be investigated in the context of				
302	biomethylation.				
303					
304	ASSOCIATED CONTENT				
305	The Supporting Information contains 1 table pertaining to analytical reagents and analytical				
306	instru	ment parameters.			
307					
308	REFERENCES				
309					
310	1.	Matschullat, J. Arsenic in the geosphere. Sci. Total Environ. 2000, 249, 297-312.			
311	2.	Mestrot, A.; Planer-Friedrich, B; Feldmann, J. Biovolatilisation: a poorly studied			
312		pathway of the arsenic biogeochemical cycle. Environ. Sci. Processes Impacts 2012,			
313		15, 1639.			
314	3.	Savage, L.; Carey, M.; Hossain, M.; Islam, M.R.; de Silva, P.M.C.S.; Williams, P.N.;			
315		Meharg, A.A. Elevated trimethylarsine oxide and inorganic arsenic in northern			
316		hemisphere summer monsoonal wet deposition. Environ. Sci. Technol. 2017, in press			
317		DOI: 10.1021/acs.est.7b04356			
318	4.	Zhang, SY.; Sun, G. –X.; Yin, X. –X.; Rensing, C.; Zhu, Y. –G. Biomethylation and			
319		volatilization of arsenic by the marine microalgae Ostreococcus tauri. Chemosph.			
320		<b>2013</b> , <i>93</i> , 47-53.			

321	5.	Zhu, Y.G.; Xue, X.M; Kappler, A.; Rosen, B.; Meharg, A.A. Linking genes to microbial
322		biogeochemical cycling: lessons from arsenic. Environ. Sci. Technol. 2017, 51, 7326-
323		7339.
324	6.	Zhao, F.J.; Zhu, Y.G.; Meharg, A.A. Methylated arsenic species in rice: geographical
325		variation, origin and uptake mechanisms. Environ. Sci. Technol. 2013, 47, 3957-3966.
326	7.	Mukai, H.; Ambe, Y.; Muku, T., Takeshita, K. Seasonal variation of methylarsenic
327		compounds in airborne particulate matter. Nature 1986, 324, 239-241.
328	8.	Nakamura, M.; Matsuzono, Y.; Tanaka, S.; Hashimito, Y. Chemical form of arsenic
329		compounds and distribution of their concentrations in the atmosphere, Appl.
330		Organomet. Chem., 4, 1990, 223-230.
331	9.	Jakob, R.; Roth, A.; Haas, K.; Krupp, E. M.; Raab, A.; Smichowski, P.; Gomez, D.;
332		Feldmann, J. Atmospheric stability of arsines and the determination of their
333		oxidative products in atmospheric aerosols ( $PM_{10}$ ): evidence of the widespread
334		phenomena of biovolatilization of arsenic. J. Environ. Monit. 2010, 12, 409-416.
335	10.	Huang, JH.; Matzner, E. Biogeochemistry of organic and inorganic arsenic species in
336		a forested catchment in Germany. Environ. Sci. Technol. 2007, 41, 1564-1569.
337	11.	Tanaka, S.; Santosa, S. J.; The concentration distribution and chemical form of
338		arsenic compounds in sea water. In: Biogeochical Processes and Ocean Flux in the
339		Western Pacific. Eds Sakai, H.; Nozaki, Y. pp. 159-170, Terra Scientific Pub. Comp.,
340		Tokyo, <b>1995</b> .
341	12.	Wurl, O.; Shelley, R. U.; Landing, W. M.; Cutter, G. A. Biogeochemistry of dissolved
342		arsenic in the temperate to tropical North Atlantic Ocean. Deep-Sea Res. Il 2015,
343		116, 240-250.

344	13. Hanaoka, K.; Nakamura, O.; Ohno, H.; Tagawa, S.; Kaise, T. Degradation of
345	arsenobetaine to inorganic arsenic by bacteria in seawater. Hydrobio. 1995, 316, 75-
346	80.
347	14. Matoušek, T.; Currier, J. M.; Trojánková, N.; Saunders, R. J.; Ishida, M. C.; González-
348	Horta, C.; Musil, S.; Mester, Z.; Stýblo, M.; Dědina, J. Selective hydride generation-
349	cryotrapping-ICP-MS for arsenic speciation analysis at pictogram levels: analysis of
350	river and sea water reference materials and human bladder epithelial cells. J. Anal.
351	At. Spectrom. <b>2013</b> , 28, 1456-1465.
352	15. Mestrot, A.; Feldmann, J.; Krupp, E. M; Hossain, M. S.; Roman-Ross, G.; Meharg, A. A.
353	Field fluxes and speciation of arsines emanating from soils. Environ. Sci. Technol.
354	<b>2011</b> , <i>45</i> , 1798-1804.
355	16. Arndt, J.; Ilge, G.; Planer-Friedrich, B. Evaluation of techniques for sampling volatile
356	arsenic on volcanoes. J. Volcan. Geotherm. Res. 2017, 331, 16-25.
357	17. Draxler, R. R.; Hess, G. D. An overview of the HYSPLIT_4 modeling system of
358	trajectories, dispersion, and deposition. Aust. Meteor. Mag. 1998, 47, 295-308.
359	18. Cutter, G. A. Metalloids in wet deposition on Bermuda: concentrations, sources, and
360	fluxes. J. Geophys. Res. <b>1993</b> , 98, 16777-16786.
361	19. Andreae, M. O. Arsenic in rain and the atmospheric mass balance of arsenic. J.
362	Geophys. Res. <b>1980,</b> 85, 4512-4518.
363	20. Blazina, T.; Läderach, A.; Jones, G. D.; Sodemann, H.; Wernli, H.; Kirchner, J. W.;
364	Winkel, H. E. Marine primary productivity as a potential indirect source of selenium
365	and other trace elements in atmospheric deposition. Environ. Sci. Technol. 2016, 51,
366	108-118.

367	21. Cutter, G.A.; Cutter, L.S.; Featherstone, A.M; Lohrenz, S.E. Antimony and arsenic
368	biogeochemistry in the western Atlantic Ocean. Deep-Sea Res. II. 2001, 48, 2895-
369	2915.
370	22. Cutter, G.A.; Cutter, L.S. Metalloids in the high latitude North Atlanic Ocean: internal
371	cycling. <i>Mar. Chem.</i> <b>1998</b> , 61, 25-36.
372	23. Michel, P.; Boutier, B.; Herbland, A.; Averty, B.; Artigas, L.F.; Auger, D. Chartier, E.
373	Behaviour of arsenic on the continental shelf of the Girode estuary: role of
374	phytoplankton in vertical fluxes during spring blood conditions. Oceanolog. Acta
375	<b>1998</b> , 21, 325-333.
376	24. Yin, X.X.; Chen, J.; Qin, J.; Sun, G.X.; Rosen, B.P.; Zhu, Y.G. Biotransformation and
377	volatilization of arsenic by three photosynthetic cyanobacteria. Plant Physiol. 2011,
378	156, 1631-1638.
379 380 381 382	

Table 1. Concentration and rates of deposition in wet precipitation for the Belfast and Kerry
 sampling sites. The LoDs for arsenic species were 0.12, 0.13 and 0.288 nM for DMAA, TMAO
 and inorganic arsenic, respectively.

		concentration (nM)			deposition (nmol/m2/d)		
	<u>n</u>	inorg. As	DMAA	TMAO	inorg. As	DMAA	TMAO
Belfast							
mean	32	0.64	0.05	0.15	2.32	0.19	0.31
S.E.		0.20	0.01	0.03	0.78	0.03	0.04
Kerry							
mean	8	1.14	0.26	0.58	0.89	0.20	0.50
S.E.		0.34	0.11	0.15	0.11	0.04	0.11

- Figure 1. Geographic location of the two sampling sites, Belfast and Kerry, used in this study, along with the weather 72-h back trajectory origins modelled by HySplit. For the back trajectories, black is for Belfast, red is for Kerry. The map image was created in the Mac application MagicMaps.



Figure 2. IC-ICP-MS chromatograms for arsenic species chemo-trapped on silver nitrate from
incubated seawater. Chromatograms are shown in week 1 (upper) and week (2) pairs for
each microcosm. All 8 microcosms were untreated for week 1, and for week 2 half were left
untreated (left-hand column) and half treated with a final concentration of 1 nM DMA. The
red shaded chromatographic trace is for the inlet trap; the black line trace is for the outlet
trap.



Figure 3. Production rate of arsines from seawater, calculated from subtracting the inlet
silver nitrate traps from the outlet trap quantified chromatograms presented in Figure 2.
Graphs show each individual microcosm by week and treatment. The horizontal red line is
the mean, the red vertical is the standard error of the mean.







