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Comparing classical and Bayesian ²¹⁰Pb dating models in human-impacted aquatic environments

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Abstract

Chronologies are an essential tool to place natural archives of environmental changes on a calendar scale. Because of this, studies that compare and assess the accuracy and precision of available dating models are essential. ^{210}Pb is a radioactive isotope which is used to date recent sediments (<150 yr). Here we contrast the chronologies resulting from two different ^{210}Pb dating models: the Constant Flux model (also known as the Constant Rate of Supply model) and the recently developed Bayesian *Plum* model. This comparison was implemented by using four sediment cores from contrasting environmental settings, and showed several benefits of using a Bayesian approach. This allows to infer variables, such as the supported levels of ^{210}Pb , crucial to the chronology and commonly estimated through either samples where an asymptotic behaviour is observed or through ^{226}Ra measurements, which themselves contain some level of uncertainty. Another step of traditional methods is the selection of the equilibrium depth in order to calculate total inventories, which carries ~~out~~ strong consequences for the resulting age depth model. *Plum*, on the other hand, infers the ^{210}Pb flux, which eliminates the need to select an equilibrium depth and allows for dating cores with incomplete inventory. When traditional methods address cores with missing inventory, ^{137}Cs time markers are commonly used to correct the chronology, but *Plum* can use these time markers to improve the chronology and to reduce the uncertainty. These aspects show that the Bayesian approach results in more robust and objective chronologies, and thus in better reconstructions of environmental change over the past centuries.

Keywords: ^{210}Pb sediment dating, Bayesian Statistics, Uncertainty, Marine, Coast, Lake.

1. Introduction

Historical records of global change provide essential information to identify the main stressors affecting ecosystems. Such environmental information can feed models of potential impact scenarios and guide the development of appropriate mitigation strategies. However, because long-term instrumental records are still scarce, natural archives such as sediment cores from aquatic ecosystems that accumulate continuously, have become one of the most relevant resources of retrospective information to establish trends of environmental changes on time-scales from years to millennia or more. This information, which includes rates of change, ages of events and timing differences between records, should be based on reliable chronologies.

^{210}Pb is the most widely used radionuclide to date sediment records within the past 110-150 years, a period that includes the Anthropocene, with huge global changes associated with the need to provide food and energy resources to a rapidly growing human population (Waters et al., 2016). ^{210}Pb ($t_{1/2}=22.23$ years) belongs to the ^{238}U natural decay series, and its abundance in sediments ($^{210}\text{Pb}_{\text{tot}}$) results from the mixture of (a) supported ^{210}Pb ($^{210}\text{Pb}_{\text{sup}}$) produced *in situ* through the decay of ^{226}Ra in the lithosphere, and (b) excess ^{210}Pb ($^{210}\text{Pb}_{\text{ex}}$) produced after ^{222}Rn decay and transferred to sediments through dry and wet deposition, water column scavenging and runoff. Because $^{210}\text{Pb}_{\text{ex}}$ is deposited at the sediment's surface and decays over time, its activity decreases with increasing sediment depth in cores from undisturbed sediment deposits, enabling the establishment of chronologies (Appleby and Oldfield, 1992).

The Constant Flux (CF) model, also known as the Constant Rate of Supply (CRS), (Robbins, 1978; Appleby, 1978) is a robust and widely used ^{210}Pb dating model. It is based on the fundamental equation that establishes that $^{210}\text{Pb}_{\text{ex}}$ activities depend on the $^{210}\text{Pb}_{\text{ex}}$ flux to the sediment surface and the sediment mass loading (Krishnaswamy et al., 1971; Sanchez-Cabeza and Ruiz-Fernández, 2012). The main hypothesis behind the CF model is that the $^{210}\text{Pb}_{\text{ex}}$ flux is constant, and this can be used to infer estimates of temporal variations of sediment accumulation rates (SAR, cm year^{-1}) and mass accumulation rates (MAR, $\text{g cm}^{-2} \text{ year}^{-1}$). Age calculations are based on the total inventory of $^{210}\text{Pb}_{\text{ex}}$ (i.e. the accumulated deposit per unit area), which requires estimates of $^{210}\text{Pb}_{\text{ex}}$ data for each core section until reaching the equilibrium depth (where $^{210}\text{Pb}_{\text{tot}}$ and $^{210}\text{Pb}_{\text{sup}}$ activities become indistinguishable). Therefore, missing $^{210}\text{Pb}_{\text{ex}}$ data should be estimated by interpolation or, in the case of cores not reaching equilibrium, missing inventory extrapolated (Appleby, 2002). These restrictive requirements might cause problems in sub-optimal cores (e.g., those with incomplete $^{210}\text{Pb}_{\text{ex}}$ inventory, poorly defined ^{210}Pb equilibrium depth, or varying $^{210}\text{Pb}_{\text{sup}}$). Therefore, some CF-derived age-models may not provide reliable chronologies, especially given the extensive global land use changes that have taken place over the past century.

Recently, Aquino-López et al. (2018) developed a new model for ^{210}Pb dating called *Plum*, which is also based on the fundamental equation and the CF hypothesis. *Plum* is a Bayesian forward model that simultaneously integrates two different processes: (1) the behaviour of the ^{210}Pb flux and the variation of $^{210}\text{Pb}_{\text{sup}}$ with depth, and (2) an age-depth function (based on Bacon, [*a Bayesian piece-wise linear model constrained by prior information on accumulation rate and variability*](#);

Blaauw and Christen 2011). These processes are parameterized, which allows the use of a Bayesian approach. With the help of millions of Monte Carlo Markov Chain (MCMC) iterations, a posterior sample can be obtained. This posterior sample combines *a priori* information of sediment accumulation rates, $^{210}\text{Pb}_{\text{ex}}$ flux (typically observed in a given sampling site type) and data ($^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra laboratory measurements, and potentially additional information from other radioactive isotopes such as ^{137}Cs and ^{14}C). Within any of the iterations, *Plum* chooses values for both the age-depth function, which assigns ages to any depth, and the ^{210}Pb flux and $^{210}\text{Pb}_{\text{sup}}$. With these parameters and the given ^{210}Pb 's known $t_{1/2}$, values for $^{210}\text{Pb}_{\text{tot}}$ can be calculated and compared to the measurements at any depth. *Plum* only requires experimental data on sediment density, $^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra activities (or at least a good approximation of $^{210}\text{Pb}_{\text{sup}}$) and the outputs include estimates of $^{210}\text{Pb}_{\text{xs}}$ flux, supported ^{210}Pb and sediment ages. Any of these parameters are expressed as posterior distributions and 95% confidence intervals.

In comparison to the traditional implementation of the CF model, which is a reverse and deterministic model (e.g. ages are directly inferred from $^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra laboratory data), *Plum* is a much more flexible forward model that allows coping with non-ideal ^{210}Pb depth profiles (variable compaction poorly-defined equilibrium depth). In addition, it can easily handle gaps in ^{210}Pb data or multiple measurements even from the same depth, and can include other types of dating information (e.g. ^{14}C dates, ^{137}Cs peak and tephtras).

In this work, we present results of a test on the efficacy and reliability of the CF and *Plum* models using data of previously published ^{210}Pb -dated sediment cores from contrasting environments (deep and shallow marine areas, a crater lake, and a

saltmarsh) where dating difficulties were encountered. Our hypothesis was that *Plum* is better than CF in managing the assumptions involved with the ^{210}Pb CF model and in estimating age uncertainties. The age models obtained were contrasted with stratigraphic markers (e.g. ^{137}Cs , Pu isotopes) for validation.

2. Study area

The sediment cores were collected from four contrasting environments in Mexico, in order to reconstruct the temporal trends of a range of global change indicators (Figure 1). Core IXW-500 (Ruiz-Fernández et al., 2019a) was collected from a deep-sea area in the southern Gulf of Mexico (Atlantic Coast) characterized by strong fluvial input, natural oil seeps, intense maritime traffic and offshore oil extraction activities, and was used to assess trace metal contamination.

Core Tehua-II (Ruiz-Fernández et al., 2009b) was sampled from a shallow marine area in the Tehuantepec Gulf (Pacific Coast) characterized by high productivity associated with fluvial inputs and intense upwelling, which sustains relevant commercial fisheries (e.g. tuna, shrimp and shark) and experiences strong shipping traffic owing to oil production and distribution activities. For this core, temporal variations of sediment accumulation rates were determined in order to reconstruct land use changes.

Core AIII (Ruiz-Fernández et al., 2016) was sampled to estimate rates of sea level rise at a saltmarsh in Estero de Urías (Gulf of California, Pacific Coast) an anthropized lagoon that hosts a harbour with shipping activities ranging from cabotage and fishing (mainly tuna and shrimp) to passenger cruises, and which also receives untreated domestic and industrial wastes.

Finally, core SAMO2014-2 (Ruiz-Fernández et al., 2019b) was collected to determine mass accumulation rates in Lake Santa Maria del Oro, a 2-km diameter and 65m deep oligomictic and mesotrophic crater lake (Cardoso-Mohedano et al., 2019). The lake is mainly used for recreational activities, and it is located within the Trans-Mexican Volcanic Belt at 65 km from the Pacific Coast.

3. Methods

The sediment cores were collected with transparent PVC tubes of different internal diameter and assisted by different types of samplers (see Table 1 for sampling details). The cores were extruded and cut into contiguous 1 cm thick sections. Samples were weighed before and after freeze-drying; sediment bulk density (g cm^{-3}) was calculated as the ratio of the total dry mass and the volume (from diameter and thickness) of each core section.

3.1. Laboratory analysis

All sediment cores were analysed in the Sediment Dating Academic Service at UNAM, following the methodology described in previous publications (Ruiz-Fernández et al., 2009, 2016, 2019a, 2019b). To summarise, $^{210}\text{Pb}_{\text{tot}}$ activities were measured using alpha spectrometry (Alpha Ensemble Ortec/Ametek) according to Ruiz-Fernández and Hillaire-Marcel (2009). $^{210}\text{Pb}_{\text{sup}}$ activities were estimated using gamma spectrometry measurements of ^{226}Ra activities (through its daughter radionuclide ^{214}Pb , 352 keV) using a low-background Ortec HPGe well-detector (Ruiz-Fernández et al., 2014). $^{210}\text{Pb}_{\text{xs}}$ activities were determined using the difference between $^{210}\text{Pb}_{\text{tot}}$ and $^{210}\text{Pb}_{\text{sup}}$ activities. To validate the ^{210}Pb -derived chronologies, the activity profiles of ^{137}Cs against depth were determined by gamma-ray spectrometry (662 keV) in all cores, and plutonium isotopes were determined by i)

alpha-particle spectrometry for core SAMO14-2 (Ruiz-Fernández et al., 2019b), ii) low-energy accelerator mass spectrometry for core EUIII (in Centro Nacional de Aceleradores, Spain; Ruiz-Fernández et al., 2016) and iii) mass spectrometry in Spiez Laboratory (Switzerland) for core IXW-500.

3.2. ^{210}Pb -derived chronologies

In the original papers, chronologies, mass and sediment accumulation rates were estimated with the constant flux (CF) model (Appleby and Oldfield, 1978; Robbins, 1978; Sanchez-Cabeza and Ruiz-Fernández, 2012). Dating uncertainties in core Tehua-II were calculated using quadratic propagation uncertainty, whereas in the other cores they were estimated by Monte Carlo simulation with 10^5 simulations (Sanchez-Cabeza et al., 2014).

The *Plum* age-depth models were obtained using the default settings. These settings dictate the prior distributions used for the supported ^{210}Pb , ^{210}Pb flux, and sedimentation rate and its variability (used within each *Bacon* section), as well as the thickness of *Bacon* sections. By default, *Plum* uses 1-cm thick *Bacon* sections (whereas *Bacon*'s default is 5, which is probably too restrictive to model sedimentation close to the surface). The default prior distribution for the supported ^{210}Pb parameters is a gamma distribution with a mean of 15 Bq kg^{-1} and a shape parameter of 2. For core IXW500, *Plum* used the ^{226}Ra measurements to obtain estimates for supported ^{210}Pb at multiple depths, whereas for the other cores supported ^{210}Pb was assumed to be constant throughout each core. For the ^{210}Pb flux, *Plum* uses a gamma distribution with a mean of $50 \text{ Bq m}^{-2} \text{ yr}^{-1}$ and a shape parameter of 2. *Plum* also uses an upper age limit, for which it uses a parameter set to $A_i=0.1$ by default, which is defined as the remaining unmeasured $^{210}\text{Pb}_{\text{ex}}$. This

variable allows for faster convergence of the MCMC, and limits the chronology according to the inferred $^{210}\text{Pb}_{\text{ex}}$ flux. *Plum* also can use information from time-markers such as the human-made peak in ^{137}Cs to refine the age-depth model, but this was ignored as we wanted to observe only the behaviour of the ^{210}Pb data on their own.

4. Results

Figures 2, 3, 4 and 5 show the comparisons between the ^{210}Pb -derived chronologies using the CF and *Plum* dating models.

4.1. TEHUA-II

In the case of TEHUA-II (Figure 2), both models agreed well with each other throughout the chronologies. The CF confidence intervals overlapped *Plum*'s credible intervals, i.e. *Plum* provided narrower intervals. In this case, *Plum* was able to extend the chronology by up to two decades, because fewer measurements were discarded in order to infer the supported ^{210}Pb . Both models agreed quite well with the 1963 time-marker.

4.2. EU-III

Figure 3 shows the comparison between the CF and *Plum* age-models of core EU-III. In this case, the chronologies agreed well with each other, providing similar age estimates and similar credible and confidence intervals. Again, *Plum* was able to extend the chronology for several decades as several more data points are used to infer the chronology: CF inferred a chronology up to a depth of 27.5 cm, whereas *Plum* reached a depth of 39 cm. Regarding the 1963 time marker, *Plum* seems to agree better as the CF model provided younger ages and its uncertainty was insufficient to enclose the time marker.

4.3. SAMO14-2

SAMO14-2's chronology (Figure 4) showed that the age estimates are in good agreement and both models enclose the time marker within its intervals. On the other hand, the model's uncertainty estimates show large differences, with *Plum* providing more conservative uncertainty estimates. Regarding the length of the chronology, *Plum* was again able to extend the chronology by several decades by its use of fewer measurements to infer the supported ^{210}Pb value.

4.4. IXW-500

In this case, the CF- and *Plum*- derived age models resulted in significant differences (Fig. 5) at different core depths. In comparison with CF ages, *Plum* ages were older in the segments surface – 4 cm and 8 – 12 cm, but comparable within the segments 5 – 8 cm and below 12 cm. *Plum* provided a more conservative age estimate, with larger and more credible intervals than the CF model. In this particular case, The *Plum*-derived age model comprised almost twice the CF age period (age estimate of 160 years at 15 cm depth and 308 years at 29 cm depth). This was caused by CF having to discard the 5 bottom measurements, whereas *Plum* was able to extract information from them.

The considerable differences between the models could be caused by the peculiar ^{226}Ra profile of this core (shown as blue dots in the supported ^{210}Pb panel in Figure 4). The largest difference between the age models was observed at mid-depth, where a minimum $^{210}\text{Pb}_{\text{ex}}$ in activity occurs, caused by large ^{226}Ra activity. This also resulted in both models presenting a change in behaviour at this depth, with a larger accumulation rate at depths lower than 9 cm. In this case, refining the chronology using more information (e.g. ^{137}Cs) could be of help, as *Plum* can include

these time markers to the likelihood function in order to obtain an integrated chronology.

It is important to note that both total ^{210}Pb and ^{226}Ra at the five deepest samples were very similar, i.e. equilibrium was reached. In this case, *Plum* uses the information provided by the rest of the core and the prior distribution to provide an informed extrapolation.

5. Discussion

5.1. Tehua-II and EU-III

As all sections of these cores were measured down to equilibrium, as expected both models provided similar results. In the case of TEHUA-II, *Plum* provided narrower credible intervals, which are a result of the use of the *Bacon* age-depth function, instead of the decay equation used by the CF model. On the other hand, the age-models for EU-III agreed well and reconstructed similar uncertainty intervals. *Plum* was able to extend the chronology by a couple of decades in the case of TEHUAII and by more than 100 yr in the case of EU-III. This is a result of *Plum*'s integration of supported ^{210}Pb in the inference process, removing the need to discard the lowermost few measurements.

5.2. SAMO14-2

SAMO14-2 presents an interesting case as relatively few data points were available for the inference. Both models presented overlapping results. In this case, *Plum* provided wider intervals as a result of the low number of samples measured for this core. The CF model provided very precise age estimates, apparently a positive outcome, but considering that less than 25% of the core sections were measured (10 measurements of 1 cm sections were used to infer the chronology to a depth of

41 cm), the more conservative intervals provided by *Plum* are more coherent. This is an important consequence of the interpolations used to estimate the $^{210}\text{Pb}_{\text{ex}}$ inventories, which do not consider the error caused by the absence of measurements. Interestingly, something similar appears to happen in simulated and real-world ^{14}C -dated cores modelled using classical and Bayesian age-models (Blaauw et al. 2018): whereas Bayesian age-models become more precise with increasing dating densities (they ‘learn’), classical age-models such as linear interpolation do not.

5.3. IXW-500

This core presented significant differences between each model. *Plum* inferred lower levels of supported ^{210}Pb at 8 – 10 cm depths in comparison to the measured ^{226}Ra (see Figure 5). It is at these depths where both models show their largest discrepancies. In this case, more measurements around the spike of ^{226}Ra could help *Plum* to model these unusual spikes. Also, *Plum* can use other time markers such as ^{137}Cs to refine the chronology.

5.4. *Plum* extended chronologies

As shown, *Plum* is able to obtain longer chronologies. In some cases, the extension is a result of fewer samples being used exclusively to infer supported ^{210}Pb and/or the use of the final sample with $^{210}\text{Pb}_{\text{ex}}$ not being discarded (which is common practice when using the CF model). These extensions provide an accurate chronology with an uncertainty similar to the rest of the chronology. In other cases, like IXW-500, it is the result of using ^{226}Ra concentrations combined with a sediment with ^{210}Pb in equilibrium. *Plum* provides a much larger extension in these chronologies, but this should be taken with extra caution, since at depths below

where $^{210}\text{Pb}_{\text{ex}}$ is available, age uncertainties become considerably higher and precision decreases. These extensions are a linear simplification, which uses information from both the data and the prior distributions for the accumulation rate and variability. Under simulation, we have observed that these extensions become more inaccurate with depth, but the credible intervals (larger than in the rest of the chronology) are able to capture the true age. These extensions are rather qualitative and should not be used to obtain conclusions on sedimentary processes and records. Because simulations are not within the scope of the present study, these are left for future research.

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5.5. CF rapid accumulation

The CF model was constructed using the decay equation as its basis, which resulted in a logarithmic age-depth function; $t(x) = \lambda^{-1} \log \left(\frac{A(0)}{A(x)} \right)$. This age-depth function works well in the most recent part of the sediment, but as the remaining activity $A(x)$ gets close to 0, the age function rapidly tends to infinity, resulting in an artificial increase of accumulation rates. This increase is not caused by natural sources but it is a result of the mathematical structure of the model and should not be included in interpretations. On the other hand, because *Plum* uses an independent age-depth function, it does not show this artificial increase in accumulation, providing a more realistic age-depth function at deeper parts of the sediment.

5.6. Overall considerations

When CF model users are confronted with a new core, a number of decisions must be made based on the available information. Because of resource constraints, typically not all core sections are measured and interpolation techniques are needed, which may induce an experimental error. By design, *Plum* does not need any kind of interpolation and solely relies on the existing data. While both the CF model and *Plum* should produce better age-models with more data (i.e. more information), *Plum* results are undoubtedly more reliable when some data are missing.

The reliability of measurement can also be sometimes an issue. For example, the CF radiochronologist might decide that a measurement is an outlier, or that concentrations close or below 0 Bq kg⁻¹ are zero. In the case of *Plum*, this is not an issue, as the system will choose, based on Bayesian statistics, which is the best set

of activities and, in some cases, disregard an outlier or ignore low (or negative) concentrations, without human intervention.

One of the critical decisions that must be taken when using the CF model is the selection of the equilibrium depth of the core, i.e., the depth where the $^{210}\text{Pb}_{\text{ex}}$ reaches zero and from where the total $^{210}\text{Pb}_{\text{ex}}$ inventory can be calculated. The equilibrium depth mainly depends on the unambiguous knowledge of the $^{210}\text{Pb}_{\text{sup}}$ value. This is often difficult to ascertain owing to i) large relative uncertainties of $^{210}\text{Pb}_{\text{ex}}$, when calculated by subtraction of ^{210}Pb and ^{226}Ra (or an estimated $^{210}\text{Pb}_{\text{sup}}$), which can be very similar at the core bottom, and ii) changes in sediment composition which can produce variations of $^{210}\text{Pb}_{\text{sup}}$ activities along a core (e.g. Carnero-Bravo et al. 2016), which will affect the total $^{210}\text{Pb}_{\text{xs}}$ inventory and thus the core age, as higher inventories will cause younger ages and vice versa. Therefore, depending on the core's ^{210}Pb activity profile, small variations of specific activities could produce age variations amounting to decades. A relevant feature of *Plum* is that $^{210}\text{Pb}_{\text{sup}}$ is not a a value to be provided by the user ~~user's input~~ but a parameter to be inferred. *Plum* uses data to infer $^{210}\text{Pb}_{\text{sup}}$, which can be ^{226}Ra measurements or ^{210}Pb which reached equilibrium. If only ^{210}Pb is available, *Plum* runs a pre-analysis based on a linear regression, where the segment which provides the largest p-value (and thus lower variability) is suggested to be used exclusively to infer $^{210}\text{Pb}_{\text{sup}}$. Because the p-value is affected by sample size, this pre-analysis should be confirmed by the user, but it has provided reasonable results in several simulations and real cores. This reduces considerably the calculation efforts and decisions of the user, and maximizes the confidence in the $^{210}\text{Pb}_{\text{sup}}$ value.

Sometimes the ^{210}Pb profile is incomplete, i.e. ^{210}Pb does not reach equilibrium. In this case, CF users need to estimate the missing inventory from either a bottom constant accumulation rate, or by using a reference date (e.g. the ^{137}Cs maximum). In the case of *Plum* this is not needed, as even in these cases *Plum* will estimate the total $^{210}\text{Pb}_{\text{ex}}$ flux by Bayesian statistics.

In overall, these and other decisions may cause different scientists, or even the same scientist at different moments, to produce different chronologies with the same data set and using the same model. As *Plum* uses existing information, age models should be identical (as long as the prior parameters are the same) and the user-based bias should largely reduced. In this sense, *Plum* is clearly more objective and we suggest that it should be widely used by ^{210}Pb radiochronologists.

~~*Plum* produces ^{210}Pb chronologies up the maximum core length (Fig. 5) based on the assumption that mass accumulation rates ($\text{g cm}^{-2} \text{yr}^{-1}$) are constant beyond the core section from where $^{210}\text{Pb}_{\text{ex}}$ is available and using both data from the most recent part of the sediment and the prior distribution. This is a useful *Plum* feature allowing practical applications (e.g. selection of core sections from pre-industrial times, a rough estimation of the maximum core age, correlation among core sections). However, caution is advised when using these extrapolated ages beyond where $^{210}\text{Pb}_{\text{ex}}$ is available, since age uncertainties become considerably higher and precision decreases. The behaviour of these extensions can only be analysed with help of simulations. Because simulations are not within the scope of this study, these are left for a future study.~~

In this comparison study, only ^{210}Pb measurements were used to produce chronologies. However, for core IXW-500 additional nuclides were measured,

including ^{137}Cs , which has known historical peaks caused by radioactive releases into the atmosphere. Such peaks are regularly used to validate or adjust ^{210}Pb -based chronologies. *Plum* can integrate such information, as well as ^{14}C and other dates, into the ^{210}Pb chronology.

6. Conclusions

The application of the CF ^{210}Pb sediment dating model requires from the radiochronologist several inputs, such as the equilibrium depth and data interpolation when not all sections are measured, a rather common case. However, with the Bayesian approach used by *Plum* this is no longer needed, thus making the dating process less subjective and more reproducible.

The comparison of both approaches show that *Plum* provides similar results to the classical CF model, with the added benefits of providing more precise uncertainty estimates when sampling density is high (like in the case of TEHUA-II and EU-III), and providing more realistic uncertainty estimates where few samples are available (SAMO14-2). Also, *Plum* is not affected by artificial increases of sedimentation in the deepest part of the sediment, unlike the CF model where the logarithmic age-function tends to infinity as background is reached. This allows the user to interpret dates deeper in the core without being affected by artificial increases in the sedimentation.

Plum' older ages, obtained by the usability of more data, appears to be in disagreement with previous chronology limits (Appleby, 1998). This point was previously addressed by Aquino-López et al. (2018), as by using simulations it was shown that these older ages remained accurate when the core is fully sampled. These older ages are also accompanied by a more realistic uncertainty (much larger

credible intervals), and should be handled with caution, even more when the core was is not fully sampled. Users and/or studies which required less uncertainty should consider discarding these older ages or providing the model with other time markers, which will improve the levels of uncertainty and help the model provide a more precise chronology.

The ability of adding extra information such as ^{137}Cs , radiocarbon dates and/or tephra dates can be used for both validating the chronology or improving it. Because *Plum* uses raw ^{210}Pb to create the chronology, multiple isotope chronologies are possible without having to pre-model the ^{210}Pb dates. This quality should not be underestimated, as the effects of the bias in the deepest part of the sediment presented by the CF model have not yet been studied. Because of all these considerations, we believe that *Plum* has the potential of providing more robust and objective ^{210}Pb chronologies.

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Data Availability

416 Sediment data are available as supplementary material.

417 **Supplementary material**

418 - Data used for ^{210}Pb dating with *Plum*.

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492

493 **Figure captions**

494 **Figure 1.** Geographical locations where sediment cores were collected.

495 **Figure 2.** *Plum* and CF age-depth models for TEHUA-II. *Plum* 95% credible intervals
496 and mean are shown as red dashed lines, and dark shadow show some iterative
497 chronologies obtained by the MCMC. The CF age-depth model is shown in blue,
498 using linear interpolation between neighbouring data points used to construct the
499 model. Bottom boxes show depths where samples were analyzed. The orange box
500 shows the ^{137}Cs 1963 time marker.

501 **Figure 3.** *Plum* and CF age-depth models for core EU-III. For explanation of symbols
502 and colours, see Fig. 2.

503 **Figure 4.** *Plum* and CF age-depth models for SAMO14-2. For explanation of
504 symbols and colours, see Fig. 2.

505 **Figure 5.** *Plum* and CF age-depth models for core IXW-500. For explanation of
506 symbols and colours, see Fig. 2. In the “Supported ^{210}Pb ” upper panel, black dots
507 represent *Plum* simulations, blue dots are the mean of the simulated values and red
508 dots are the ^{226}Ra measurements

509

Table 1. Sampling information for sediment cores from Mexican marine and lacustrine areas.

Core ID	Coordinates	Sampling date and depth (m)	Length (cm)	Sediment characteristics	Corer type and internal diameter (cm)
IXW-500 ^a	19°26.649' N	06/08/2015	25	Marine muddy	Multicorer
	93°53.323' W	1010		>94% silt + clay	9.5
TEHUA-II ^b	15° 59.987' N	15/10/2004	18	Marine sandy	Reineck
	94° 48.469' W	67		>75% sand	7
EU-III ^c	23°09.322'N	05/05/2012	50	Coastal hypersaline	Push core
	106°19.672'W	0		>60% silt	10
SAMO14-2 ^d	21°22.187' N	28/04/2019	78	Lacustrine	Uwitec™
	104°34.335'W	48		>80% silt	8.6

^aRuiz-Fernández et al., 2019a; ^bRuiz-Fernández et al., 2009b, ^cRuiz-Fernandez et al., 2016; ^dRuiz-Fernandez et al., 2019b.

Comparing classical and Bayesian ^{210}Pb dating models in human-impacted aquatic environments

*Marco A. Aquino-López, Ana Carolina Ruiz-Fernández, Maarten Blaauw, Joan-Albert Sanchez-Cabeza**

Answers to reviewers

Dear Editor,

We thank you for your support to revise our manuscript, and thank the reviewers for their efforts in reviewing this manuscript.

We have addressed each issue mentioned by the reviewers, provide an answer in this document, and the manuscript changes are shown in track mode. We have also used the opportunity to make some minor editorial changes.

We sincerely hope that the manuscript is now acceptable for publication.

Best regards,

Prof. J.A. Sanchez-Cabeza
UNAM

Reviewer #1

As the time interval covered by the Pb-210 lasts only 100 years, any extrapolation beyond this range is quite speculative. Despite the warnings in the present manuscript, it is quite dangerous to propose, as valid ages, extrapolations down to 300-400 years. Even applying C-14 ages together, it will be difficult to validate this extrapolation, because of the marine reservoir effect. Therefore, I would like to propose to restrict ages not older than 150 years.

This point was addressed in the original paper (Aquino-López et al. 2018). In the section named Chronology Limit, the use of a dynamic age limit was discussed, which takes into consideration not only a single value like the one presented by Appleby (1998), or alternative values such as 100 or 150 yr. Aquino-López et al. (2018) presented a series of simulated data where we know the true age values, and this allows us to observe the accuracy of the model. These simulations show that the model does not lose accuracy even in older ages, even though error estimates become larger. Users which are uncomfortable with large uncertainties can decide to discard portions of the age-depth model, but we feel this should be left to user discretion and to the levels of uncertainty relevant to each study.

In fact, the authors of this manuscript are in debate about the validity of extensions of several decades up to centuries. Further research based on simulations is planned to properly test this important issue. However, this is outside of the scope of the current manuscript, and we have clarified this in the revised version.

In order to clarify the issue and focus the reader interest, we have summarized sub-section 5.4 *Plum extended chronologies* and the second last paragraph of the discussion section as:

5.4. Plum extended chronologies

As shown, Plum is able to obtain longer chronologies. In some cases, the extension is a result of fewer samples being used exclusively to infer supported ^{210}Pb and/or the use of the final sample with $^{210}\text{Pb}_{\text{ex}}$ not being discarded (which is common practice when using the CF model). These extensions provide an accurate chronology with an uncertainty similar to the rest of the chronology. In other cases, like IXW-500, it is the result of using ^{226}Ra concentrations combined with a sediment with ^{210}Pb in equilibrium. Plum provides a much larger extension in these chronologies, but this should be taken with extra caution, since at depths below where $^{210}\text{Pb}_{\text{ex}}$ is available, age uncertainties become considerably higher and precision decreases. These extensions are a linear simplification, which uses information from both the data and the prior distributions for the accumulation rate and variability. Under simulation, we have observed that these extensions become more inaccurate with depth, but the credible intervals (larger than in the rest of the chronology) are able to capture the true age. These extensions are rather qualitative and should not be used to obtain conclusions on sedimentary processes and records. Because simulations are not within the scope of the present study, these are left for future research.

Also, we have included a paragraph in the Conclusions section:

Plum' older ages, obtained by the usability of more data, appears to be in disagreement with previous chronology limits (Appleby, 1998). This point was previously addressed by Aquino-López et al. (2018), as by using simulations it was shown that these older ages remained accurate when the core is fully sampled. These older ages are also accompanied by a more realistic uncertainty (much larger credible intervals), and should be handled with caution, even more when the core was not fully sampled. Users and/or studies which require less uncertainty should consider discarding these older ages or providing the model with other time markers, which will improve the levels of uncertainty and help the model provide a more precise chronology.

Note that not all C-14 dates in the case studies were from marine sites.

Reviewer #2

This publication is an example of the use of the new Bayesian PLUM model to date sedimentary boreholes using lead-210 measurements. It perfectly illustrates the interest of this model compared to more classical models (here the CRS model).

I would have liked an example of the Bayesian approach also using artificial radionuclides to see the complete interest of the model as well as to understand the integration of additional proxy in PLUM. In the end, is the approach completely independent of the model user or does the user tune the model in the same way as Bacon. In this case, the most important thing is that the users of the lead-210 data clearly explain their assumptions, basic data etc. (Mustaphi et al. 2019). This work deserves prompt publication.

Many thanks to the reviewer for the positive comments. We decided to not incorporate data points of artificial radionuclides into the Plum age-models, so as to provide fair comparisons with the CRS model which uses such time markers as validation only. The manuscript states: "Plum also can use information from time-markers such as the human-made peak in ^{137}Cs to refine the age-depth model, but this was ignored as we wanted to observe only the behaviour of the ^{210}Pb data on their own." The incorporation of other dating techniques is left to future studies where we will discuss the effect and benefits this additional information has on the resulting chronology.

As for the comment on 'tuning': Plum, like Bacon, is a Bayesian approach that uses prior distributions to reconstruct the age-depth relationship. Although the default settings have been set to allow for a wide variety of sites to be age-modelled, for some cores these values need to be modified. We believe that no approach can produce entirely automated age-models without any user input. If so desired however, the approach could be modified by replacing the *Bacon* component by, say, *BChron* which has fewer user-provided settings.

Minor comments:

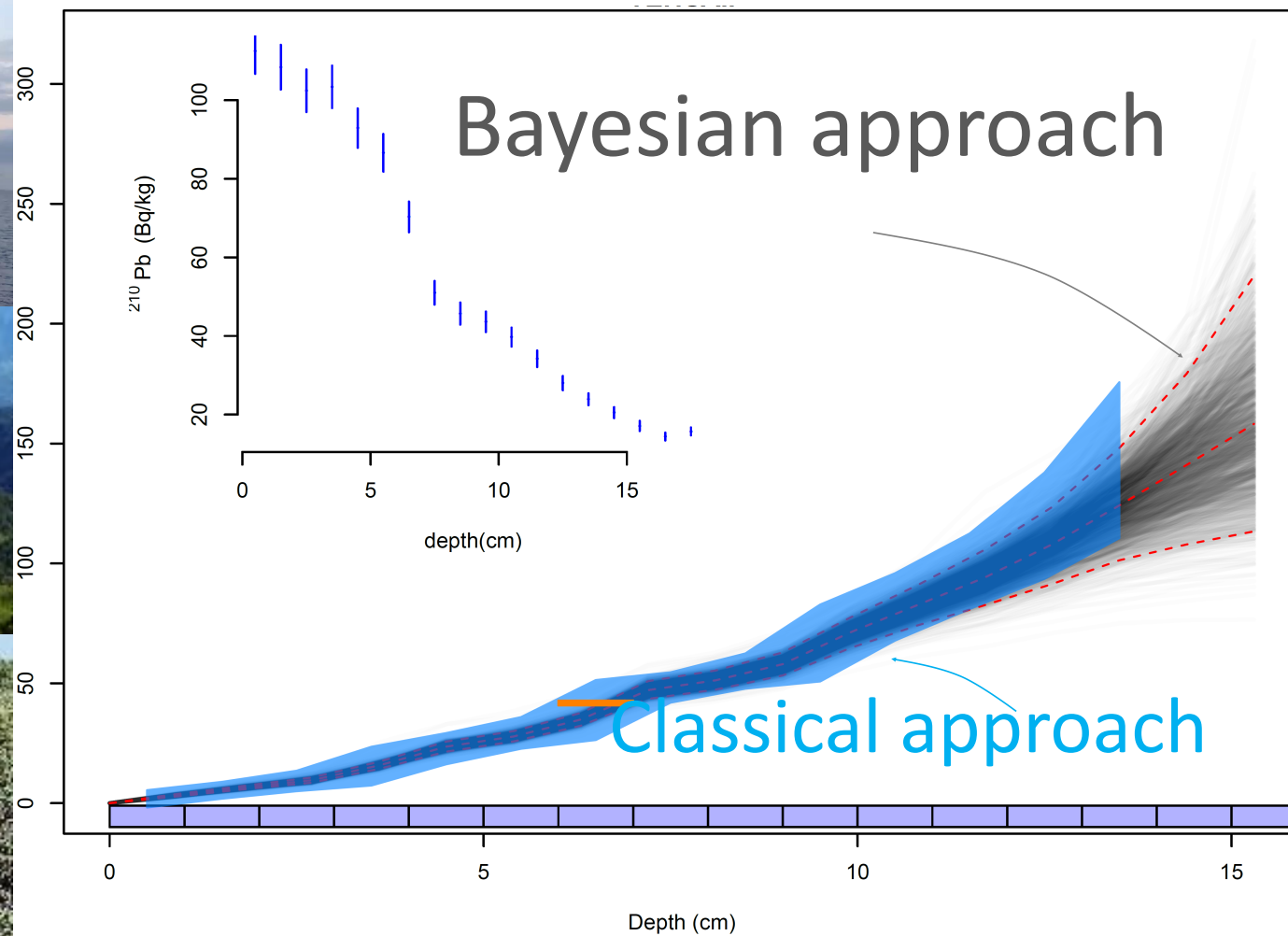
I. 83 explain BACON acronym

Bacon is not an acronym, but we have now clarified the model:

... (based on Bacon, a Bayesian piece-wise linear model constrained by prior information on accumulation rate and variability; Blaauw and Christen 2011).

I.316: user's input: please rephrase.

Rephrased to "... is not a value to be provided by the user but..."



Comparing classical and Bayesian ²¹⁰Pb dating models in human-impacted aquatic environments

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Abstract

Chronologies are an essential tool to place natural archives of environmental changes on a calendar scale. Because of this, studies that compare and assess the accuracy and precision of available dating models are essential. ^{210}Pb is a radioactive isotope which is used to date recent sediments (<150 yr). Here we contrast the chronologies resulting from two different ^{210}Pb dating models: the Constant Flux model (also known as the Constant Rate of Supply model) and the recently developed Bayesian *Plum* model. This comparison was implemented by using four sediment cores from contrasting environmental settings, and showed several benefits of using a Bayesian approach. This allows to infer variables, such as the supported levels of ^{210}Pb , crucial to the chronology and commonly estimated through either samples where an asymptotic behaviour is observed or through ^{226}Ra measurements, which themselves contain some level of uncertainty. Another step of traditional methods is the selection of the equilibrium depth in order to calculate total inventories, which carries strong consequences for the resulting age depth model. *Plum*, on the other hand, infers the ^{210}Pb flux, which eliminates the need to select an equilibrium depth and allows for dating cores with incomplete inventory. When traditional methods address cores with missing inventory, ^{137}Cs time markers are commonly used to correct the chronology, but *Plum* can use these time markers to improve the chronology and to reduce the uncertainty. These aspects show that the Bayesian approach results in more robust and objective chronologies, and thus in better reconstructions of environmental change over the past centuries.

Keywords: ^{210}Pb sediment dating, Bayesian Statistics, Uncertainty, Marine, Coast, Lake.

1. Introduction

Historical records of global change provide essential information to identify the main stressors affecting ecosystems. Such environmental information can feed models of potential impact scenarios and guide the development of appropriate mitigation strategies. However, because long-term instrumental records are still scarce, natural archives such as sediment cores from aquatic ecosystems that accumulate continuously, have become one of the most relevant resources of retrospective information to establish trends of environmental changes on time-scales from years to millennia or more. This information, which includes rates of change, ages of events and timing differences between records, should be based on reliable chronologies.

^{210}Pb is the most widely used radionuclide to date sediment records within the past 110-150 years, a period that includes the Anthropocene, with huge global changes associated with the need to provide food and energy resources to a rapidly growing human population (Waters et al., 2016). ^{210}Pb ($t_{1/2}=22.23$ years) belongs to the ^{238}U natural decay series, and its abundance in sediments ($^{210}\text{Pb}_{\text{tot}}$) results from the mixture of (a) supported ^{210}Pb ($^{210}\text{Pb}_{\text{sup}}$) produced *in situ* through the decay of ^{226}Ra in the lithosphere, and (b) excess ^{210}Pb ($^{210}\text{Pb}_{\text{ex}}$) produced after ^{222}Rn decay and transferred to sediments through dry and wet deposition, water column scavenging and runoff. Because $^{210}\text{Pb}_{\text{ex}}$ is deposited at the sediment's surface and decays over time, its activity decreases with increasing sediment depth in cores from undisturbed sediment deposits, enabling the establishment of chronologies (Appleby and Oldfield, 1992).

The Constant Flux (CF) model, also known as the Constant Rate of Supply (CRS), (Robbins, 1978; Appleby, 1978) is a robust and widely used ^{210}Pb dating model. It is based on the fundamental equation that establishes that $^{210}\text{Pb}_{\text{ex}}$ activities depend on the $^{210}\text{Pb}_{\text{ex}}$ flux to the sediment surface and the sediment mass loading (Krishnaswamy et al., 1971; Sanchez-Cabeza and Ruiz-Fernández, 2012). The main hypothesis behind the CF model is that the $^{210}\text{Pb}_{\text{ex}}$ flux is constant, and this can be used to infer estimates of temporal variations of sediment accumulation rates (SAR, cm year^{-1}) and mass accumulation rates (MAR, $\text{g cm}^{-2} \text{ year}^{-1}$). Age calculations are based on the total inventory of $^{210}\text{Pb}_{\text{ex}}$ (i.e. the accumulated deposit per unit area), which requires estimates of $^{210}\text{Pb}_{\text{ex}}$ data for each core section until reaching the equilibrium depth (where $^{210}\text{Pb}_{\text{tot}}$ and $^{210}\text{Pb}_{\text{sup}}$ activities become indistinguishable). Therefore, missing $^{210}\text{Pb}_{\text{ex}}$ data should be estimated by interpolation or, in the case of cores not reaching equilibrium, missing inventory extrapolated (Appleby, 2002). These restrictive requirements might cause problems in sub-optimal cores (e.g., those with incomplete $^{210}\text{Pb}_{\text{ex}}$ inventory, poorly defined ^{210}Pb equilibrium depth, or varying $^{210}\text{Pb}_{\text{sup}}$). Therefore, some CF-derived age-models may not provide reliable chronologies, especially given the extensive global land use changes that have taken place over the past century.

Recently, Aquino-López et al. (2018) developed a new model for ^{210}Pb dating called *Plum*, which is also based on the fundamental equation and the CF hypothesis. *Plum* is a Bayesian forward model that simultaneously integrates two different processes: (1) the behaviour of the ^{210}Pb flux and the variation of $^{210}\text{Pb}_{\text{sup}}$ with depth, and (2) an age-depth function (based on *Bacon*, a Bayesian piece-wise linear model constrained by prior information on accumulation rate and variability;

Blaauw and Christen 2011). These processes are parameterized, which allows the use of a Bayesian approach. With the help of millions of Monte Carlo Markov Chain (MCMC) iterations, a posterior sample can be obtained. This posterior sample combines *a priori* information of sediment accumulation rates, $^{210}\text{Pb}_{\text{ex}}$ flux (typically observed in a given sampling site type) and data ($^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra laboratory measurements, and potentially additional information from other radioactive isotopes such as ^{137}Cs and ^{14}C). Within any of the iterations, *Plum* chooses values for both the age-depth function, which assigns ages to any depth, and the ^{210}Pb flux and $^{210}\text{Pb}_{\text{sup}}$. With these parameters and the given ^{210}Pb 's known $t_{1/2}$, values for $^{210}\text{Pb}_{\text{tot}}$ can be calculated and compared to the measurements at any depth. *Plum* only requires experimental data on sediment density, $^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra activities (or at least a good approximation of $^{210}\text{Pb}_{\text{sup}}$) and the outputs include estimates of $^{210}\text{Pb}_{\text{xs}}$ flux, supported ^{210}Pb and sediment ages. Any of these parameters are expressed as posterior distributions and 95% confidence intervals.

In comparison to the traditional implementation of the CF model, which is a reverse and deterministic model (e.g. ages are directly inferred from $^{210}\text{Pb}_{\text{tot}}$ and ^{226}Ra laboratory data), *Plum* is a much more flexible forward model that allows coping with non-ideal ^{210}Pb depth profiles (variable compaction poorly-defined equilibrium depth). In addition, it can easily handle gaps in ^{210}Pb data or multiple measurements even from the same depth, and can include other types of dating information (e.g. ^{14}C dates, ^{137}Cs peak and tephtras).

In this work, we present results of a test on the efficacy and reliability of the CF and *Plum* models using data of previously published ^{210}Pb -dated sediment cores from contrasting environments (deep and shallow marine areas, a crater lake, and a

saltmarsh) where dating difficulties were encountered. Our hypothesis was that *Plum* is better than CF in managing the assumptions involved with the ^{210}Pb CF model and in estimating age uncertainties. The age models obtained were contrasted with stratigraphic markers (e.g. ^{137}Cs , Pu isotopes) for validation.

2. Study area

The sediment cores were collected from four contrasting environments in Mexico, in order to reconstruct the temporal trends of a range of global change indicators (Figure 1). Core IXW-500 (Ruiz-Fernández et al., 2019a) was collected from a deep-sea area in the southern Gulf of Mexico (Atlantic Coast) characterized by strong fluvial input, natural oil seeps, intense maritime traffic and offshore oil extraction activities, and was used to assess trace metal contamination.

Core Tehua-II (Ruiz-Fernández et al., 2009b) was sampled from a shallow marine area in the Tehuantepec Gulf (Pacific Coast) characterized by high productivity associated with fluvial inputs and intense upwelling, which sustains relevant commercial fisheries (e.g. tuna, shrimp and shark) and experiences strong shipping traffic owing to oil production and distribution activities. For this core, temporal variations of sediment accumulation rates were determined in order to reconstruct land use changes.

Core AIII (Ruiz-Fernández et al., 2016) was sampled to estimate rates of sea level rise at a saltmarsh in Estero de Urías (Gulf of California, Pacific Coast) an anthropized lagoon that hosts a harbour with shipping activities ranging from cabotage and fishing (mainly tuna and shrimp) to passenger cruises, and which also receives untreated domestic and industrial wastes.

Finally, core SAMO2014-2 (Ruiz-Fernández et al., 2019b) was collected to determine mass accumulation rates in Lake Santa Maria del Oro, a 2-km diameter and 65m deep oligomictic and mesotrophic crater lake (Cardoso-Mohedano et al., 2019). The lake is mainly used for recreational activities, and it is located within the Trans-Mexican Volcanic Belt at 65 km from the Pacific Coast.

3. Methods

The sediment cores were collected with transparent PVC tubes of different internal diameter and assisted by different types of samplers (see Table 1 for sampling details). The cores were extruded and cut into contiguous 1 cm thick sections. Samples were weighed before and after freeze-drying; sediment bulk density (g cm^{-3}) was calculated as the ratio of the total dry mass and the volume (from diameter and thickness) of each core section.

3.1. Laboratory analysis

All sediment cores were analysed in the Sediment Dating Academic Service at UNAM, following the methodology described in previous publications (Ruiz-Fernández et al., 2009, 2016, 2019a, 2019b). To summarise, $^{210}\text{Pb}_{\text{tot}}$ activities were measured using alpha spectrometry (Alpha Ensemble Ortec/Ametek) according to Ruiz-Fernández and Hillaire-Marcel (2009). $^{210}\text{Pb}_{\text{sup}}$ activities were estimated using gamma spectrometry measurements of ^{226}Ra activities (through its daughter radionuclide ^{214}Pb , 352 keV) using a low-background Ortec HPGe well-detector (Ruiz-Fernández et al., 2014). $^{210}\text{Pb}_{\text{xs}}$ activities were determined using the difference between $^{210}\text{Pb}_{\text{tot}}$ and $^{210}\text{Pb}_{\text{sup}}$ activities. To validate the ^{210}Pb -derived chronologies, the activity profiles of ^{137}Cs against depth were determined by gamma-ray spectrometry (662 keV) in all cores, and plutonium isotopes were determined by i)

alpha-particle spectrometry for core SAMO14-2 (Ruiz-Fernández et al., 2019b), ii) low-energy accelerator mass spectrometry for core EUIII (in Centro Nacional de Aceleradores, Spain; Ruiz-Fernández et al., 2016) and iii) mass spectrometry in Spiez Laboratory (Switzerland) for core IXW-500.

3.2. ^{210}Pb -derived chronologies

In the original papers, chronologies, mass and sediment accumulation rates were estimated with the constant flux (CF) model (Appleby and Oldfield, 1978; Robbins, 1978; Sanchez-Cabeza and Ruiz-Fernández, 2012). Dating uncertainties in core Tehua-II were calculated using quadratic propagation uncertainty, whereas in the other cores they were estimated by Monte Carlo simulation with 10^5 simulations (Sanchez-Cabeza et al., 2014).

The *Plum* age-depth models were obtained using the default settings. These settings dictate the prior distributions used for the supported ^{210}Pb , ^{210}Pb flux, and sedimentation rate and its variability (used within each *Bacon* section), as well as the thickness of *Bacon* sections. By default, *Plum* uses 1-cm thick *Bacon* sections (whereas *Bacon*'s default is 5, which is probably too restrictive to model sedimentation close to the surface). The default prior distribution for the supported ^{210}Pb parameters is a gamma distribution with a mean of 15 Bq kg^{-1} and a shape parameter of 2. For core IXW500, *Plum* used the ^{226}Ra measurements to obtain estimates for supported ^{210}Pb at multiple depths, whereas for the other cores supported ^{210}Pb was assumed to be constant throughout each core. For the ^{210}Pb flux, *Plum* uses a gamma distribution with a mean of $50 \text{ Bq m}^{-2} \text{ yr}^{-1}$ and a shape parameter of 2. *Plum* also uses an upper age limit, for which it uses a parameter set to $A_i=0.1$ by default, which is defined as the remaining unmeasured $^{210}\text{Pb}_{\text{ex}}$. This

variable allows for faster convergence of the MCMC, and limits the chronology according to the inferred $^{210}\text{Pb}_{\text{ex}}$ flux. *Plum* also can use information from time-markers such as the human-made peak in ^{137}Cs to refine the age-depth model, but this was ignored as we wanted to observe only the behaviour of the ^{210}Pb data on their own.

4. Results

Figures 2, 3, 4 and 5 show the comparisons between the ^{210}Pb -derived chronologies using the CF and *Plum* dating models.

4.1. TEHUA-II

In the case of TEHUA-II (Figure 2), both models agreed well with each other throughout the chronologies. The CF confidence intervals overlapped *Plum*'s credible intervals, i.e. *Plum* provided narrower intervals. In this case, *Plum* was able to extend the chronology by up to two decades, because fewer measurements were discarded in order to infer the supported ^{210}Pb . Both models agreed quite well with the 1963 time-marker.

4.2. EU-III

Figure 3 shows the comparison between the CF and *Plum* age-models of core EU-III. In this case, the chronologies agreed well with each other, providing similar age estimates and similar credible and confidence intervals. Again, *Plum* was able to extend the chronology for several decades as several more data points are used to infer the chronology: CF inferred a chronology up to a depth of 27.5 cm, whereas *Plum* reached a depth of 39 cm. Regarding the 1963 time marker, *Plum* seems to agree better as the CF model provided younger ages and its uncertainty was insufficient to enclose the time marker.

4.3. SAMO14-2

SAMO14-2's chronology (Figure 4) showed that the age estimates are in good agreement and both models enclose the time marker within its intervals. On the other hand, the model's uncertainty estimates show large differences, with *Plum* providing more conservative uncertainty estimates. Regarding the length of the chronology, *Plum* was again able to extend the chronology by several decades by its use of fewer measurements to infer the supported ^{210}Pb value.

4.4. IXW-500

In this case, the CF- and *Plum*- derived age models resulted in significant differences (Fig. 5) at different core depths. In comparison with CF ages, *Plum* ages were older in the segments surface – 4 cm and 8 – 12 cm, but comparable within the segments 5 – 8 cm and below 12 cm. *Plum* provided a more conservative age estimate, with larger and more credible intervals than the CF model. In this particular case, The *Plum*-derived age model comprised almost twice the CF age period (age estimate of 160 years at 15 cm depth and 308 years at 29 cm depth). This was caused by CF having to discard the 5 bottom measurements, whereas *Plum* was able to extract information from them.

The considerable differences between the models could be caused by the peculiar ^{226}Ra profile of this core (shown as blue dots in the supported ^{210}Pb panel in Figure 4). The largest difference between the age models was observed at mid-depth, where a minimum $^{210}\text{Pb}_{\text{ex}}$ in activity occurs, caused by large ^{226}Ra activity. This also resulted in both models presenting a change in behaviour at this depth, with a larger accumulation rate at depths lower than 9 cm. In this case, refining the chronology using more information (e.g. ^{137}Cs) could be of help, as *Plum* can include

these time markers to the likelihood function in order to obtain an integrated chronology.

It is important to note that both total ^{210}Pb and ^{226}Ra at the five deepest samples were very similar, i.e. equilibrium was reached. In this case, *Plum* uses the information provided by the rest of the core and the prior distribution to provide an informed extrapolation.

5. Discussion

5.1. Tehua-II and EU-III

As all sections of these cores were measured down to equilibrium, as expected both models provided similar results. In the case of TEHUA-II, *Plum* provided narrower credible intervals, which are a result of the use of the *Bacon* age-depth function, instead of the decay equation used by the CF model. On the other hand, the age-models for EU-III agreed well and reconstructed similar uncertainty intervals. *Plum* was able to extend the chronology by a couple of decades in the case of TEHUAII and by more than 100 yr in the case of EU-III. This is a result of *Plum*'s integration of supported ^{210}Pb in the inference process, removing the need to discard the lowermost few measurements.

5.2. SAMO14-2

SAMO14-2 presents an interesting case as relatively few data points were available for the inference. Both models presented overlapping results. In this case, *Plum* provided wider intervals as a result of the low number of samples measured for this core. The CF model provided very precise age estimates, apparently a positive outcome, but considering that less than 25% of the core sections were measured (10 measurements of 1 cm sections were used to infer the chronology to a depth of

41 cm), the more conservative intervals provided by *Plum* are more coherent. This is an important consequence of the interpolations used to estimate the $^{210}\text{Pb}_{\text{ex}}$ inventories, which do not consider the error caused by the absence of measurements. Interestingly, something similar appears to happen in simulated and real-world ^{14}C -dated cores modelled using classical and Bayesian age-models (Blaauw et al. 2018): whereas Bayesian age-models become more precise with increasing dating densities (they ‘learn’), classical age-models such as linear interpolation do not.

5.3. IXW-500

This core presented significant differences between each model. *Plum* inferred lower levels of supported ^{210}Pb at 8 – 10 cm depths in comparison to the measured ^{226}Ra (see Figure 5). It is at these depths where both models show their largest discrepancies. In this case, more measurements around the spike of ^{226}Ra could help *Plum* to model these unusual spikes. Also, *Plum* can use other time markers such as ^{137}Cs to refine the chronology.

5.4. *Plum* extended chronologies

As shown, *Plum* is able to obtain longer chronologies. In some cases, the extension is a result of fewer samples being used exclusively to infer supported ^{210}Pb and/or the use of the final sample with $^{210}\text{Pb}_{\text{ex}}$ not being discarded (which is common practice when using the CF model). These extensions provide an accurate chronology with an uncertainty similar to the rest of the chronology. In other cases, like IXW-500, it is the result of using ^{226}Ra concentrations combined with a sediment with ^{210}Pb in equilibrium. *Plum* provides a much larger extension in these chronologies, but this should be taken with extra caution, since at depths below

where $^{210}\text{Pb}_{\text{ex}}$ is available, age uncertainties become considerably higher and precision decreases. These extensions are a linear simplification, which uses information from both the data and the prior distributions for the accumulation rate and variability. Under simulation, we have observed that these extensions become more inaccurate with depth, but the credible intervals (larger than in the rest of the chronology) are able to capture the true age. These extensions are rather qualitative and should not be used to obtain conclusions on sedimentary processes and records. Because simulations are not within the scope of the present study, these are left for future research.

5.5. *CF rapid accumulation*

The CF model was constructed using the decay equation as its basis, which resulted in a logarithmic age-depth function; $t(x) = \lambda^{-1} \log \left(\frac{A(0)}{A(x)} \right)$. This age-depth function works well in the most recent part of the sediment, but as the remaining activity $A(x)$ gets close to 0, the age function rapidly tends to infinity, resulting in an artificial increase of accumulation rates. This increase is not caused by natural sources but it is a result of the mathematical structure of the model and should not be included in interpretations. On the other hand, because *Plum* uses an independent age-depth function, it does not show this artificial increase in accumulation, providing a more realistic age-depth function at deeper parts of the sediment.

5.6. *Overall considerations*

When CF model users are confronted with a new core, a number of decisions must be made based on the available information. Because of resource constraints, typically not all core sections are measured and interpolation techniques are needed,

which may induce an experimental error. By design, *Plum* does not need any kind of interpolation and solely relies on the existing data. While both the CF model and *Plum* should produce better age-models with more data (i.e. more information), *Plum* results are undoubtedly more reliable when some data are missing.

The reliability of measurement can also be sometimes an issue. For example, the CF radiochronologist might decide that a measurement is an outlier, or that concentrations close or below 0 Bq kg⁻¹ are zero. In the case of *Plum*, this is not an issue, as the system will choose, based on Bayesian statistics, which is the best set of activities and, in some cases, disregard an outlier or ignore low (or negative) concentrations, without human intervention.

One of the critical decisions that must be taken when using the CF model is the selection of the equilibrium depth of the core, i.e., the depth where the ²¹⁰Pb_{ex} reaches zero and from where the total ²¹⁰Pb_{ex} inventory can be calculated. The equilibrium depth mainly depends on the unambiguous knowledge of the ²¹⁰Pb_{sup} value. This is often difficult to ascertain owing to i) large relative uncertainties of ²¹⁰Pb_{ex}, when calculated by subtraction of ²¹⁰Pb and ²²⁶Ra (or an estimated ²¹⁰Pb_{sup}), which can be very similar at the core bottom, and ii) changes in sediment composition which can produce variations of ²¹⁰Pb_{sup} activities along a core (e.g. Carnero-Bravo et al. 2016), which will affect the total ²¹⁰Pb_{xs} inventory and thus the core age, as higher inventories will cause younger ages and vice versa. Therefore, depending on the core's ²¹⁰Pb activity profile, small variations of specific activities could produce age variations amounting to decades. A relevant feature of *Plum* is that ²¹⁰Pb_{sup} is not a value to be provided by the user but a parameter to be inferred. *Plum* uses data to infer ²¹⁰Pb_{sup}, which can be ²²⁶Ra measurements or ²¹⁰Pb which

reached equilibrium. If only ^{210}Pb is available, *Plum* runs a pre-analysis based on a linear regression, where the segment which provides the largest p-value (and thus lower variability) is suggested to be used exclusively to infer $^{210}\text{Pb}_{\text{sup}}$. Because the p-value is affected by sample size, this pre-analysis should be confirmed by the user, but it has provided reasonable results in several simulations and real cores. This reduces considerably the calculation efforts and decisions of the user, and maximizes the confidence in the $^{210}\text{Pb}_{\text{sup}}$ value.

Sometimes the ^{210}Pb profile is incomplete, i.e. ^{210}Pb does not reach equilibrium. In this case, CF users need to estimate the missing inventory from either a bottom constant accumulation rate, or by using a reference date (e.g. the ^{137}Cs maximum). In the case of *Plum* this is not needed, as even in these cases *Plum* will estimate the total $^{210}\text{Pb}_{\text{ex}}$ flux by Bayesian statistics.

In overall, these and other decisions may cause different scientists, or even the same scientist at different moments, to produce different chronologies with the same data set and using the same model. As *Plum* uses existing information, age models should be identical (as long as the prior parameters are the same) and the user-based bias should largely reduced. In this sense, *Plum* is clearly more objective and we suggest that it should be widely used by ^{210}Pb radiochronologists.

In this comparison study, only ^{210}Pb measurements were used to produce chronologies. However, for core IXW-500 additional nuclides were measured, including ^{137}Cs , which has known historical peaks caused by radioactive releases into the atmosphere. Such peaks are regularly used to validate or adjust ^{210}Pb -based chronologies. *Plum* can integrate such information, as well as ^{14}C and other dates, into the ^{210}Pb chronology.

6. Conclusions

The application of the CF ^{210}Pb sediment dating model requires from the radiochronologist several inputs, such as the equilibrium depth and data interpolation when not all sections are measured, a rather common case. However, with the Bayesian approach used by *Plum* this is no longer needed, thus making the dating process less subjective and more reproducible.

The comparison of both approaches show that *Plum* provides similar results to the classical CF model, with the added benefits of providing more precise uncertainty estimates when sampling density is high (like in the case of TEHUA-II and EU-III), and providing more realistic uncertainty estimates where few samples are available (SAMO14-2). Also, *Plum* is not affected by artificial increases of sedimentation in the deepest part of the sediment, unlike the CF model where the logarithmic age-function tends to infinity as background is reached. This allows the user to interpret dates deeper in the core without being affected by artificial increases in the sedimentation.

Plum' older ages, obtained by the usability of more data, appears to be in disagreement with previous chronology limits (Appleby, 1998). This point was previously addressed by Aquino-López et al. (2018), as by using simulations it was shown that these older ages remained accurate when the core is fully sampled. These older ages are also accompanied by a more realistic uncertainty (much larger credible intervals), and should be handled with caution, even more when the core is not fully sampled. Users and/or studies which require less uncertainty should consider discarding these older ages or providing the model with other time markers,

which will improve the levels of uncertainty and help the model provide a more precise chronology.

The ability of adding extra information such as ^{137}Cs , radiocarbon dates and/or tephra dates can be used for both validating the chronology or improving it. Because *Plum* uses raw ^{210}Pb to create the chronology, multiple isotope chronologies are possible without having to pre-model the ^{210}Pb dates. This quality should not be underestimated, as the effects of the bias in the deepest part of the sediment presented by the CF model have not yet been studied. Because of all these considerations, we believe that *Plum* has the potential of providing more robust and objective ^{210}Pb chronologies.

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Data Availability

Sediment data are available as supplementary material.

Supplementary material

- Data used for ^{210}Pb dating with *Plum*.

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467

468 **Figure captions**

469 **Figure 1.** Geographical locations where sediment cores were collected.

470 **Figure 2.** *Plum* and CF age-depth models for TEHUA-II. *Plum* 95% credible intervals
471 and mean are shown as red dashed lines, and dark shadow show some iterative
472 chronologies obtained by the MCMC. The CF age-depth model is shown in blue,
473 using linear interpolation between neighbouring data points used to construct the
474 model. Bottom boxes show depths where samples were analyzed. The orange box
475 shows the ^{137}Cs 1963 time marker.

476 **Figure 3.** *Plum* and CF age-depth models for core EU-III. For explanation of symbols
477 and colours, see Fig. 2.

478 **Figure 4.** *Plum* and CF age-depth models for SAMO14-2. For explanation of
479 symbols and colours, see Fig. 2.

480 **Figure 5.** *Plum* and CF age-depth models for core IXW-500. For explanation of
481 symbols and colours, see Fig. 2. In the “Supported ^{210}Pb ” upper panel, black dots
482 represent *Plum* simulations, blue dots are the mean of the simulated values and red
483 dots are the ^{226}Ra measurements

484

Table 1. Sampling information for sediment cores from Mexican marine and lacustrine areas.

Core ID	Coordinates	Sampling date and depth (m)	Length (cm)	Sediment characteristics	Corer type and internal diameter (cm)
IXW-500 ^a	19°26.649' N	06/08/2015	25	Marine muddy	Multicorer
	93°53.323' W	1010		>94% silt + clay	9.5
TEHUA-II ^b	15° 59.987' N	15/10/2004	18	Marine sandy	Reineck
	94° 48.469' W	67		>75% sand	7
EU-III ^c	23°09.322'N	05/05/2012	50	Coastal hypersaline	Push core
	106°19.672'W	0		>60% silt	10
SAMO14-2 ^d	21°22.187' N	28/04/2019	78	Lacustrine	Uwitec™
	104°34.335'W	48		>80% silt	8.6

^aRuiz-Fernández et al., 2019a; ^bRuiz-Fernández et al., 2009b, ^cRuiz-Fernandez et al., 2016; ^dRuiz-Fernandez et al., 2019b.

Figure 1

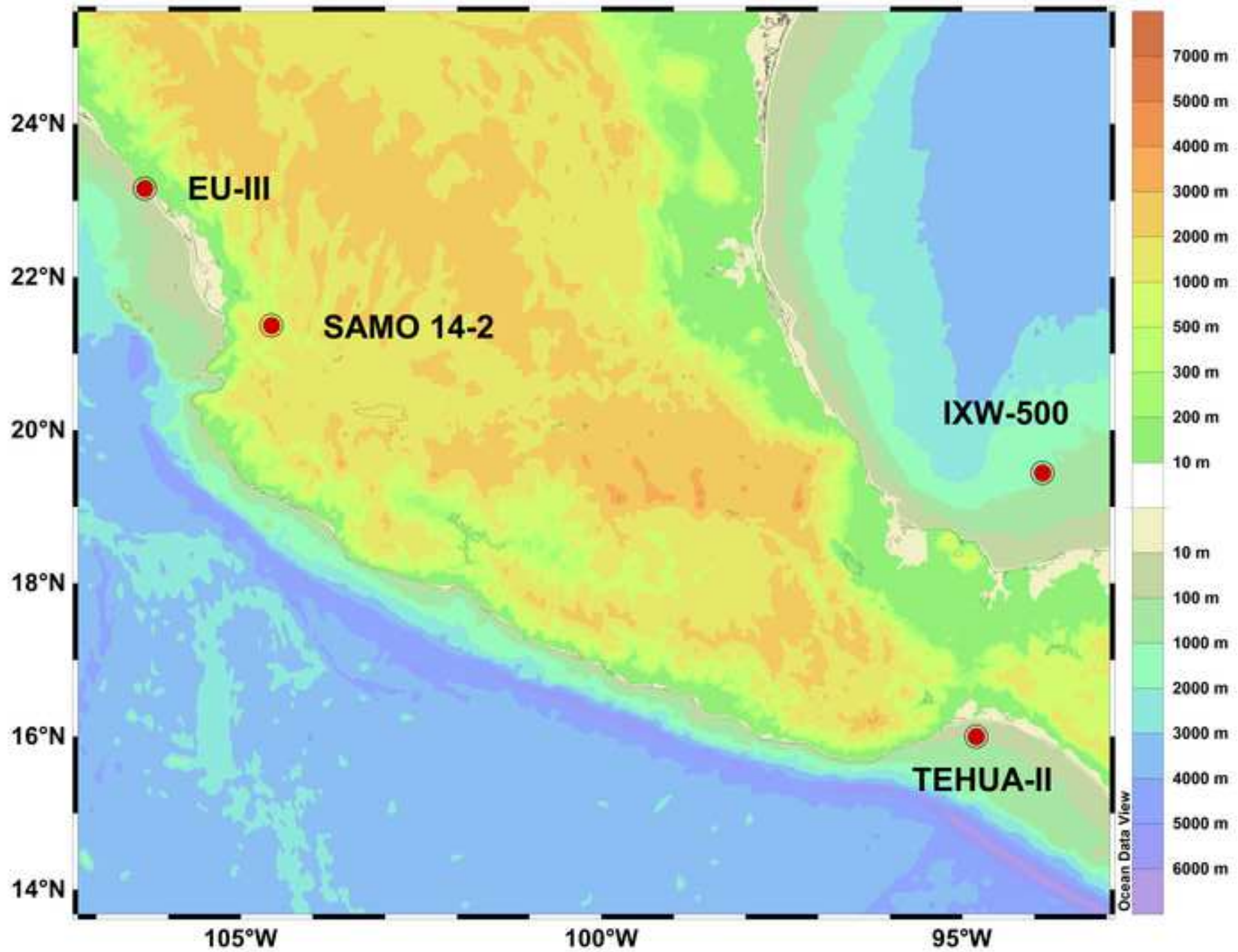
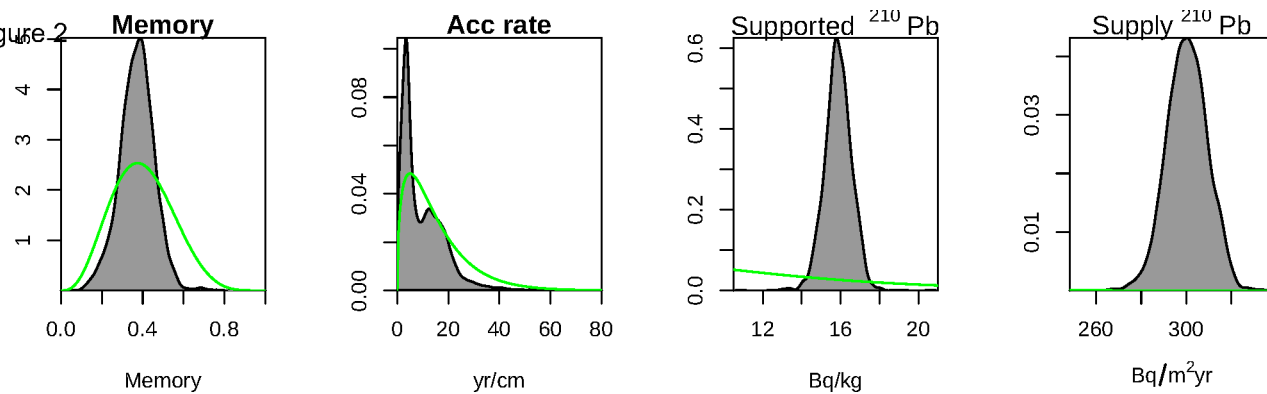


Figure 2



TEHUAIL

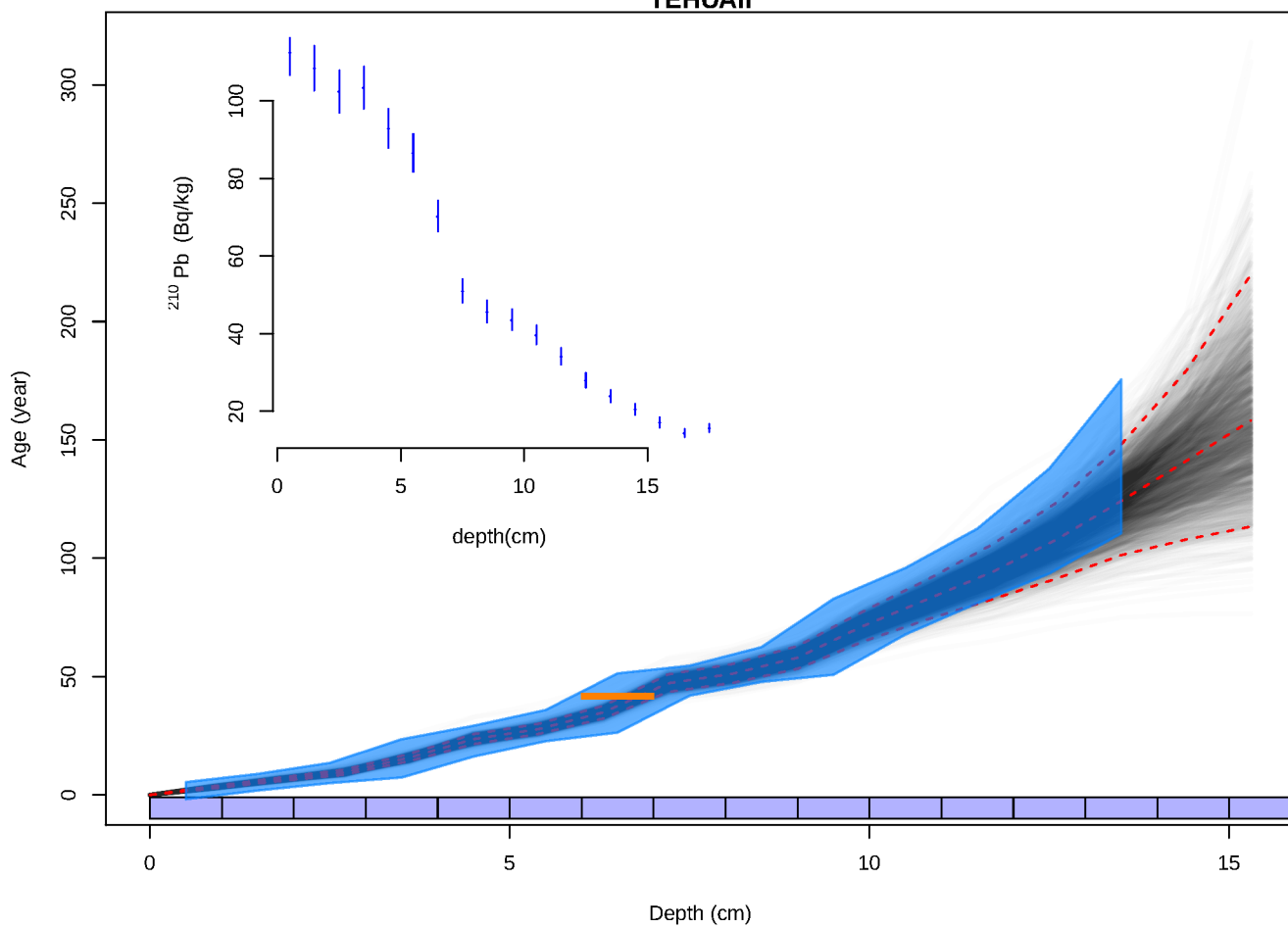
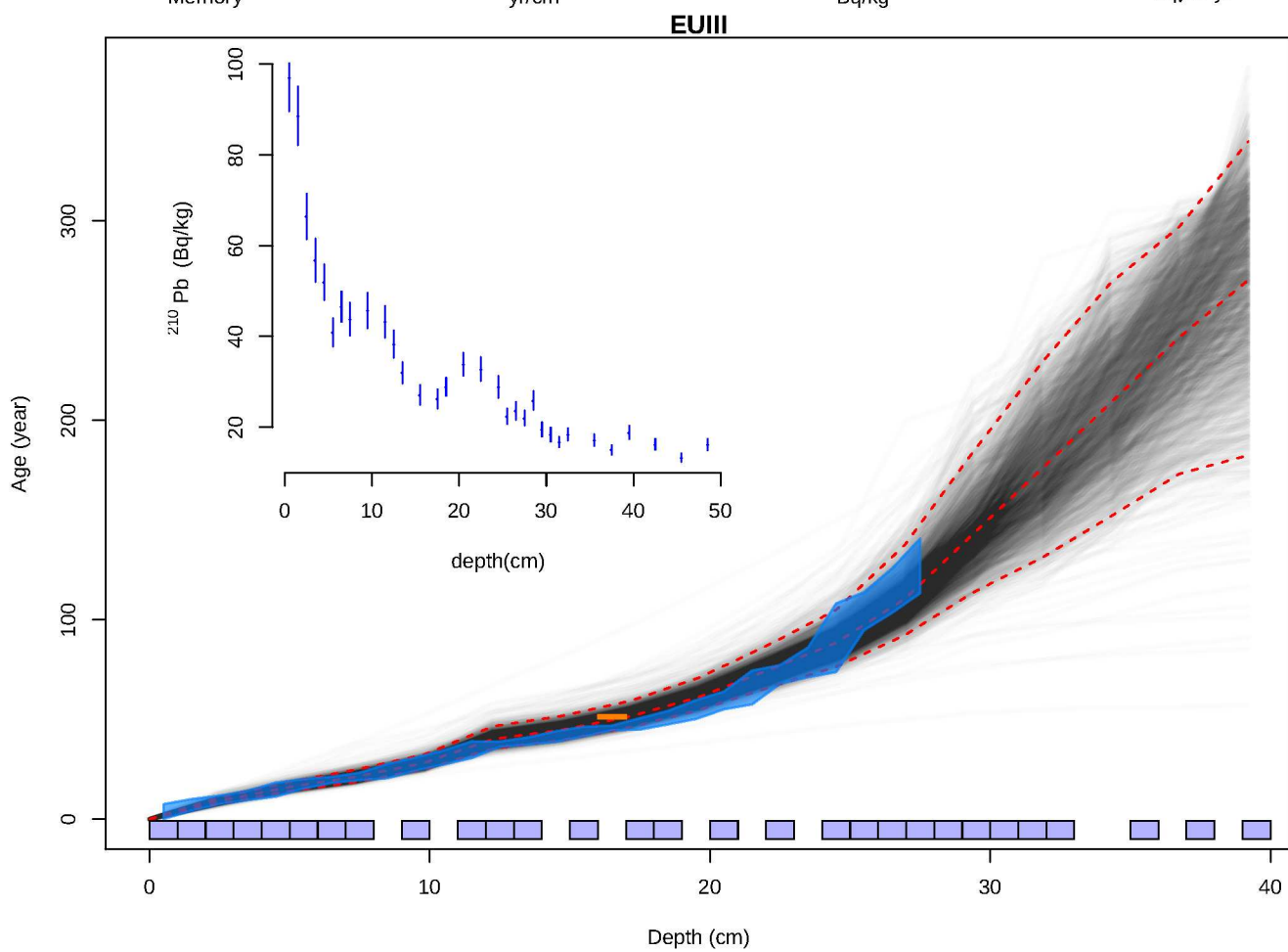
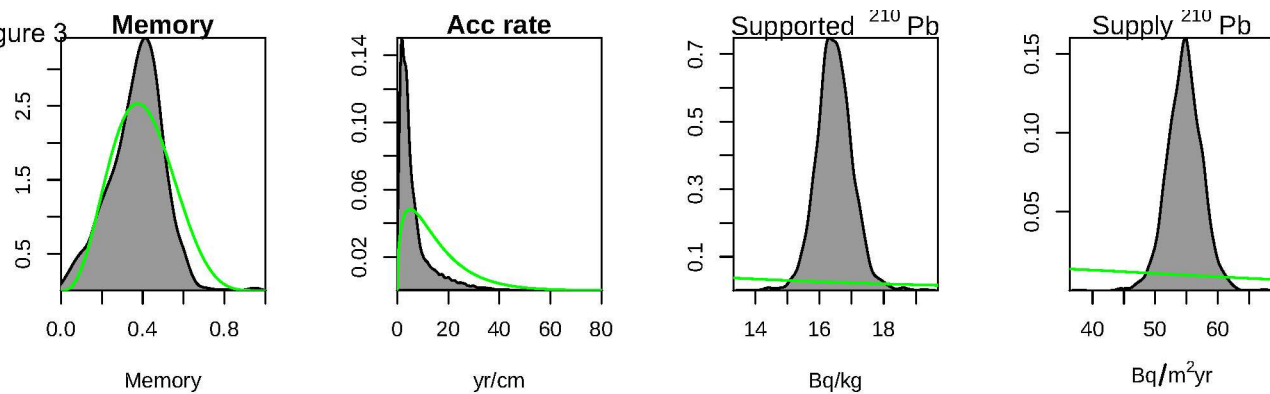
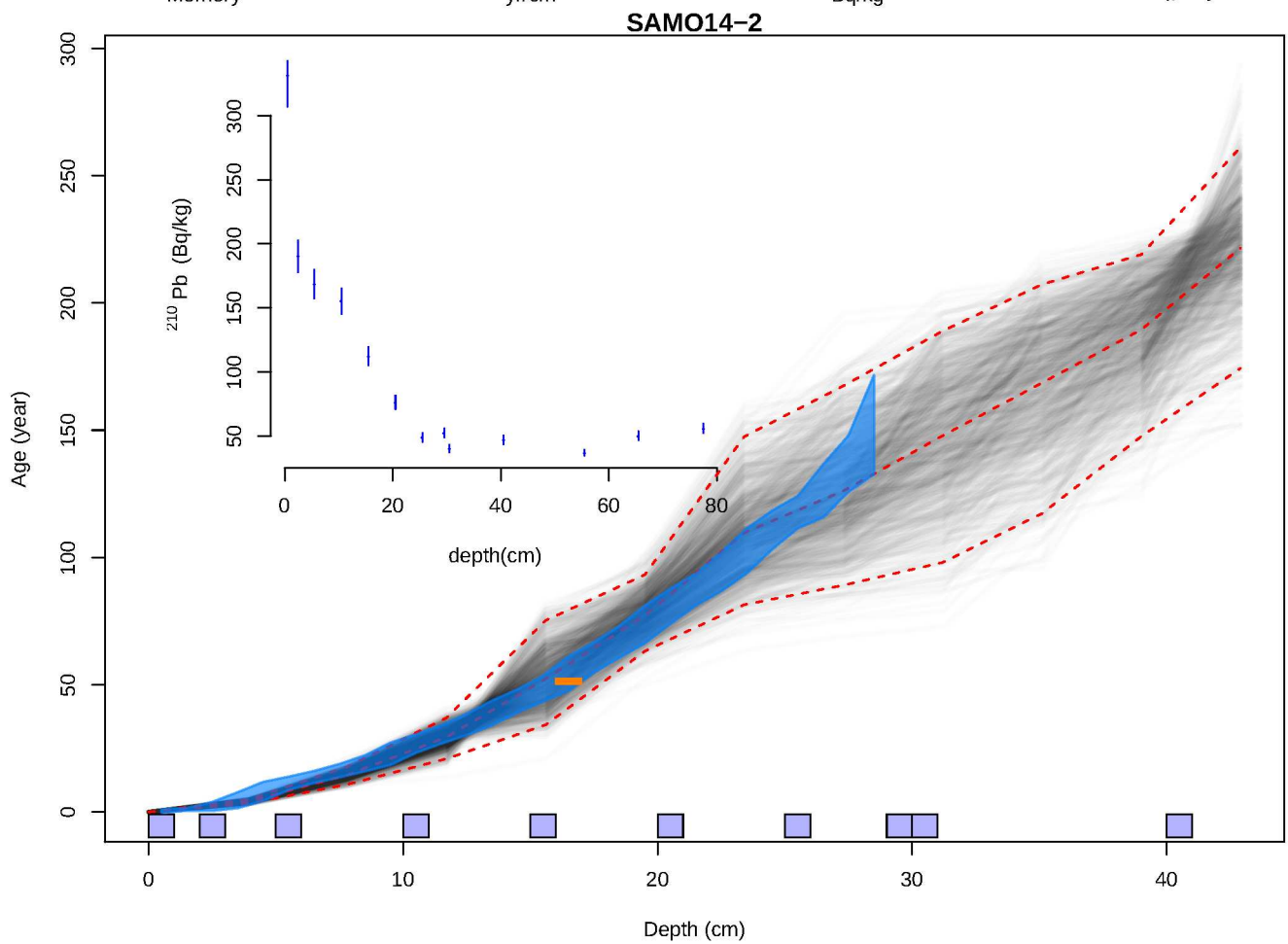
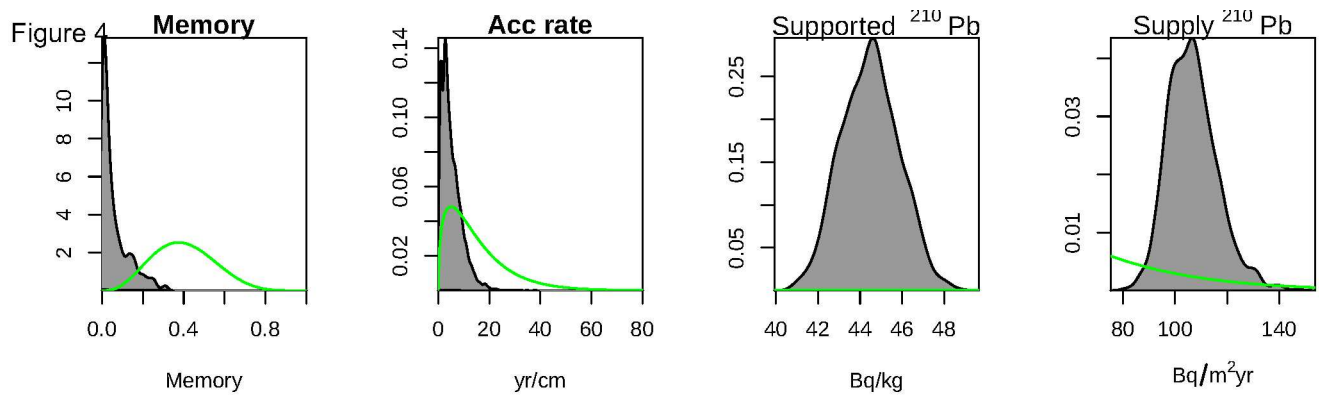
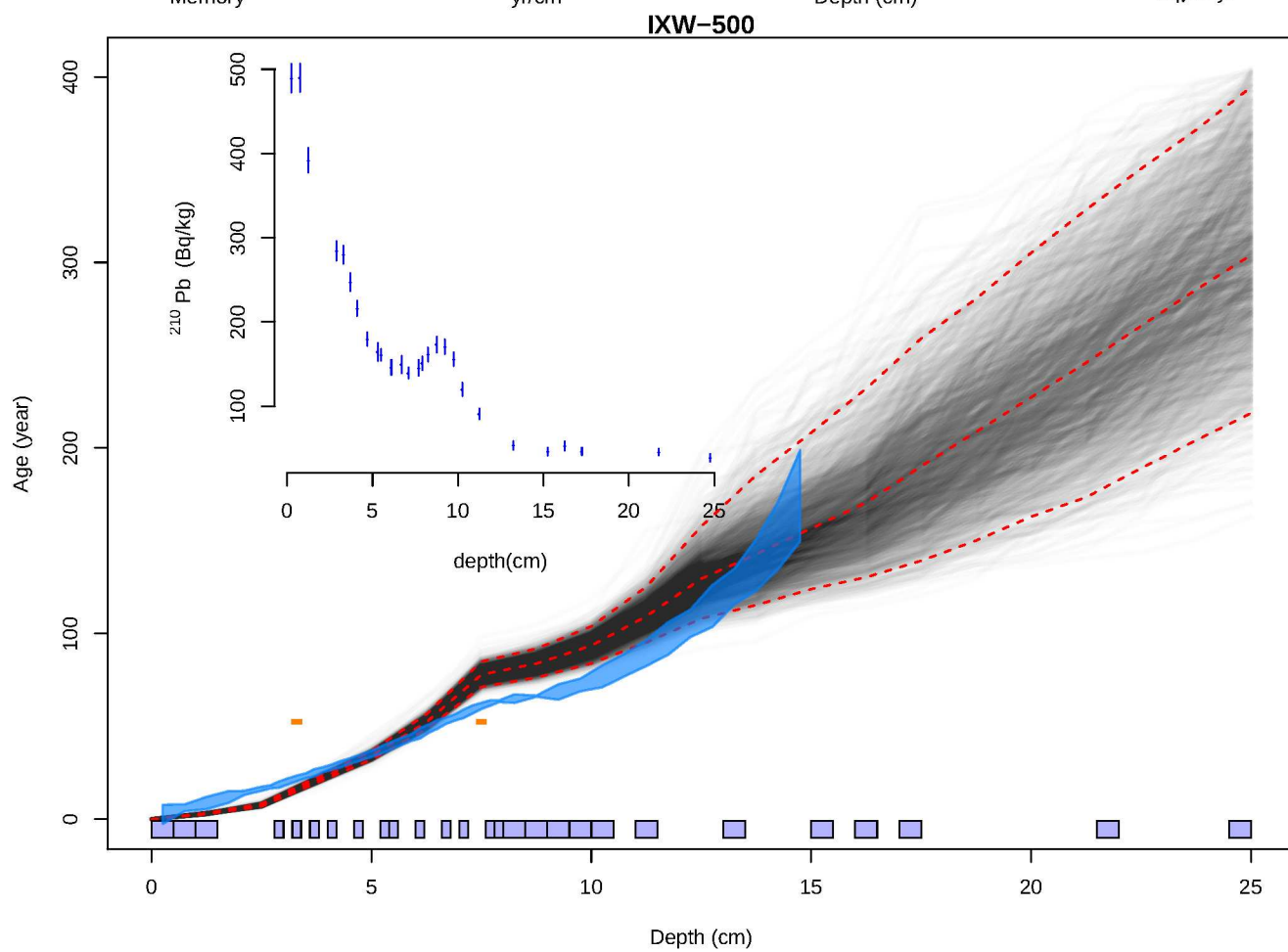
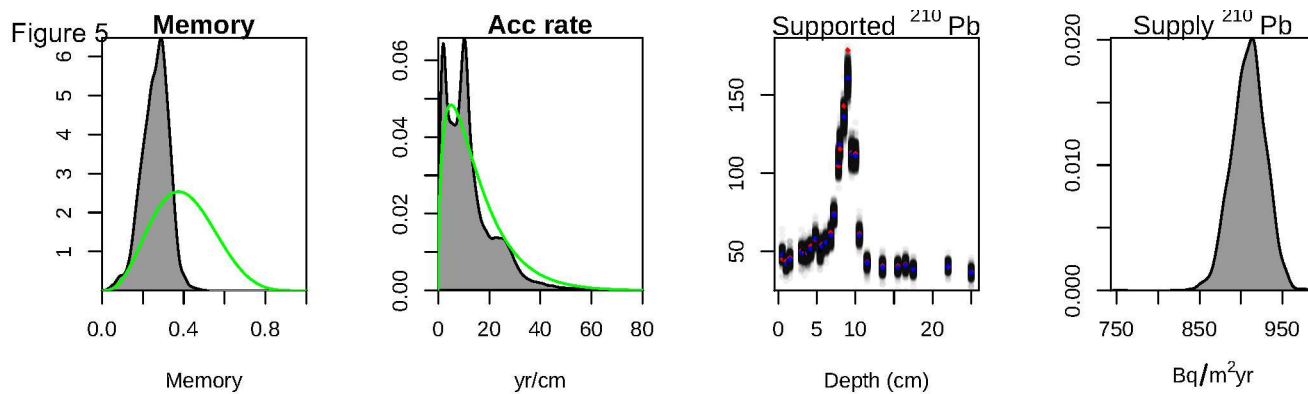
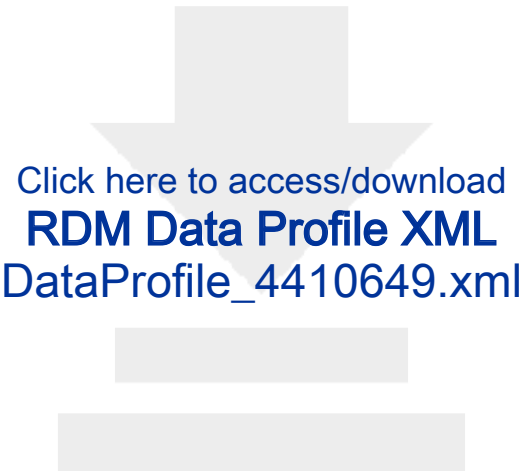


Figure 3









Declaration of interests

- ☒ The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.
- ☐ The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Declarations of interest: none.



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