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1 Laboratory and numerical investigation of saline intrusion in fractured coastal aquifers
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7 **Highlights**

- 8 • Saline intrusion in laboratory fractured aquifers was studied for the first time
- 9 • The closer the horizontal fractures were to the sea the larger the saline wedge was
- 10 • Fracture depth determines the extend of saline intrusion in the vertical direction
- 11 • Vertical fractures lead to a significant increase in the width of the mixing zone

12 **Abstract**

13 Laboratory scale experiments and numerical modelling were employed in this study to investigate
14 saltwater intrusion in fractured aquifers. Saline intrusion was initiated in one homogeneous and six
15 fractured experimental aquifers containing individual discontinuities of varying length and
16 orientation. Automated image processing enabled high precision quantification of three intrusion
17 variables, the toe length of the saline wedge, the width of the mixing zone and the aquifer fraction
18 occupied by saltwater. A dual porosity model was successfully utilized to recreate the experimental
19 data and expand the study's findings through rigorous sensitivity analysis. The presence of fractures
20 significantly impacted all three intrusion variables under consideration. The length of intrusion was
21 negatively correlated to the horizontal fracture's distance from the systems' seaward boundary. It was
22 demonstrated that for the same fractured aquifer, the presence of a discontinuity can either limit or
23 augment saline intrusion, depending on the applied hydraulic gradient. For gradients steeper than a
24 critical head difference, at which the toe length was the same for both the fractured and the
25 equivalent homogeneous aquifer, intrusion was suppressed further seaward, while for milder ones it
26 intensified. The distance of horizontal fracture from the aquifer's base determined the extend of
27 intrusion in the vertical direction. In general, the longer the discontinuities were, the more significant
28 their impact on groundwater dynamics. In the case of vertical factures, whenever the saline wedges
29 reached their position, the discontinuities contributed significantly in the widening of the mixing zone,
30 while having limited effect on the other two intrusion characteristics. In aquifers with discontinuities
31 adjacent to their side boundaries, a distinct distribution of saltwater concentration was identified,
32 distinguishing them from the rest of the aquifers.

33 **Keywords:**

34 Saline intrusion, Fractures, Sandbox experiments, SUTRA, Dual-porosity model

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39 **1. Introduction**

40 One fourth of the global population depends on freshwater pumped from fractured aquifers (Ford
41 and Williams, 2007), leading to a significant effort to map them and study their hydrogeological
42 characteristics (Bakalowicz et al., 2007; Custodio, 2009; Chen et al., 2017; Montiel et al., 2018). The
43 expected increase in average water consumption, alongside sea level rise due to global warming, is
44 anticipated to intensify saltwater intrusion (SWI) in coastal aquifers. Since fractured systems are more
45 vulnerable to saltwater intrusion than other aquifer types (Arthur et al., 2007), understanding the
46 physical mechanisms governing this phenomenon is of utmost importance. Saline intrusion has been
47 studied in fractured aquifers on multiple locations worldwide, including Canada (Allen et al., 2003),
48 France (Arfib and Charlier, 2016), Greece (Arfib and Marsily, 2004), Ireland (Perriquet et al., 2014),
49 Italy (Masciopinto, 2006), Korea (Lim et al., 2013), the United Kingdom (MacAllister et al., 2018) and
50 the United States (Xu et al., 2016).

51 Numerical modelling is a long established tool for simulating groundwater flow in real world systems.
52 The numerical approaches employed in the simulation of fractured porous media can be grouped into
53 two categories, implicit and explicit fracture representation models (Sebben et al., 2015; Berre et al.,
54 2018). Implicit fracture representation approaches include single continuum or equivalent porous
55 medium (EPM) models and dual porosity models (DPM). EPM models (Scalnon et al., 2003; Giudici et
56 al., 2012) are single continuum models in which the fractures are represented by adapting the
57 permeability of the porous medium, according to the characteristics of the existing fractures. Dual
58 porosity models (Larsbo and Jarvis, 2005; Fahs et al., 2014) consist of two continua; the matrix, which
59 represents the porous medium and the fracture, which usually has much higher porosity and
60 permeability values. Explicit fracture representation approaches include discrete fracture matrix
61 (DFM) and discrete fracture network (DFN) models. In DFM models the fluid is located in the explicitly
62 represented fractures and the surrounding porous medium alike. DFM models preserve some
63 fractures, while the rest are upscaled and replaced by averaged porous medium quantities. Two
64 dimensional DFM models incorporate fractures as 2-dimensional elements (Jaffré et al. 2005; Angot
65 et al. 2009; Flemisch et al. 2017; Koohbor et al. 2020). DFN models explicitly represent the majority of
66 discontinuities present in a fractured aquifer, and fluid flow occurs mainly within the fracture network.
67 Two-dimensional DFN models represent the geometrical properties of each individual fracture,
68 incorporating them, as 1-dimensional line boundaries between elements (Quinn et al., 2006;
69 Papadopoulou et al., 2008; Hirthe and Graf, 2015; Sebben and Werner, 2016; Ren et al., 2017).
70 Similarly, in 3-d DFN models, fractures are expressed as 2-dimensional planes. The application of each
71 method is associated with specific benefits and limitations (Samardzioska and Popov, 2005; Blessent
72 et al., 2013). EPM approaches are characterized by low levels of computational and model complexity,
73 coming at the cost of oversimplification, that results in inability to represent groundwater flow
74 dynamics in more complex problems. Dual porosity models are generally more successful in simulating
75 the impact of fractures on groundwater flow, to do so though, they introduce different parameters,
76 such as hydraulic conductivity inside the discontinuities and exchange rate coefficients between the
77 matrix and fractures, which are difficult to determine accurately. By contrast, DFM and DFN models
78 are computationally intensive, while the explicit mapping of discontinuities in real-world aquifers is
79 oftentimes problematic.

80 EPM models are prevalent in regional level studies of SWI in fractured coastal aquifers (Nocchi and
81 Salleolini, 2013; Romanazzi et al., 2015; De Filippis et al., 2016; Steiakakis et al., 2016; Zhao et al.,
82 2016). Dual porosity modelling is more commonly utilized to simulate saline intrusion in aquifers with
83 existing conduit networks, this bigger discontinuity size permits a more precise determination of the
84 network's position inside the aquifers. In the majority of these investigations, flow is considered
85 Darcian both in the porous medium and in the fractures (Xu et al., 2018; Feo et al., 2019; Kreyns et al.,
86 2020). Nevertheless, multiple numerical approaches have been proposed to incorporate turbulent

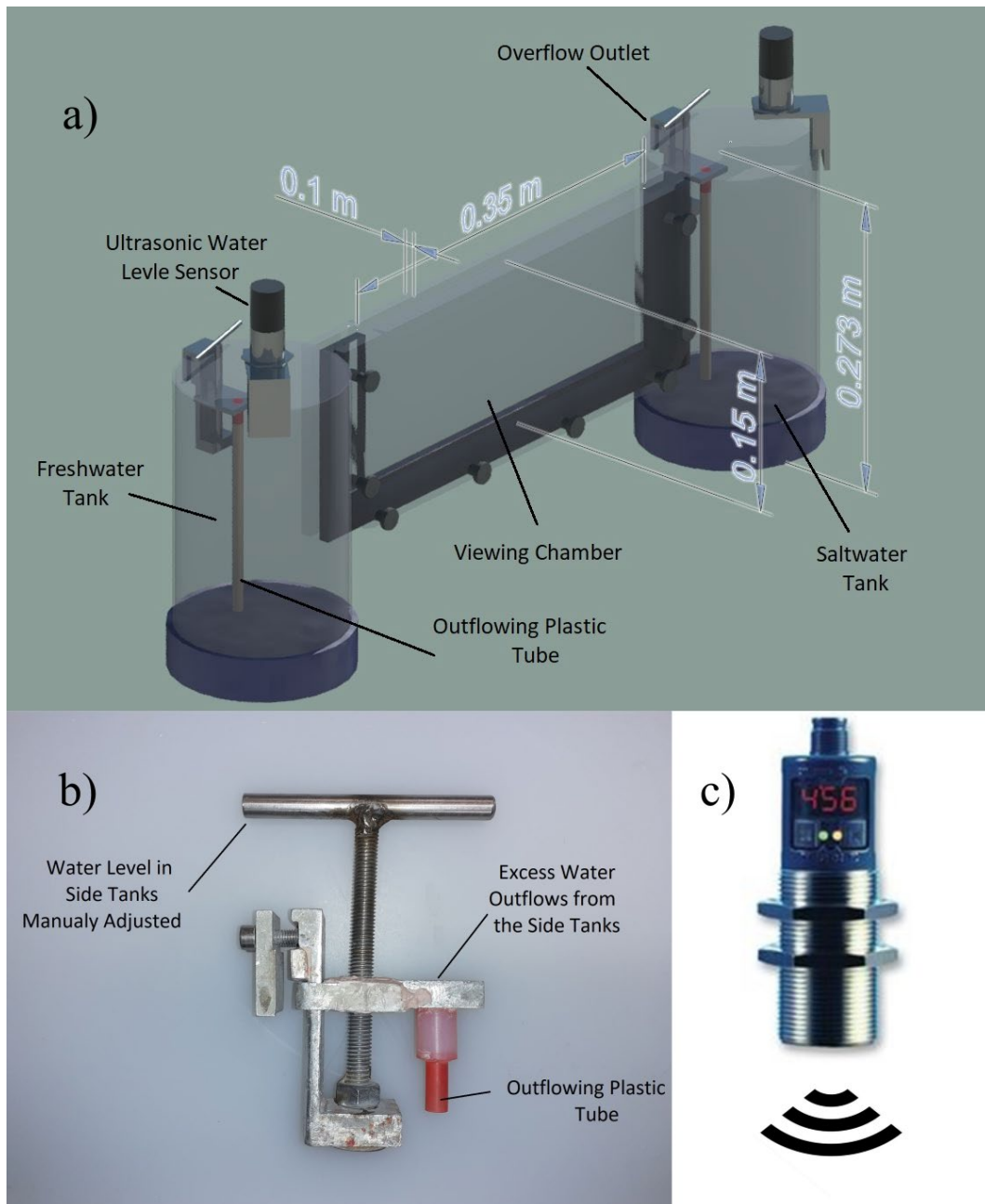
87 flow inside the conduits (Arfib and Marsily, 2004; Xu and Hu, 2017; Xu et al., 2019;). Sebben et al.
88 (2015) and Mozafari et al. (2018) simulated a variation of the well-known Henry problem in 2-
89 dimensional aquifers with well-defined networks of individual fractures using DFN modelling. Koohbor
90 et al. (2019) quantified the impact that the uncertainty associated with the position and density of
91 fractures, in discrete fracture networks, has on the simulated saline intrusion dynamics. Both Dokou
92 and Karantzas (2012) and Khadra and Stuyfzand (2018) created hybrid models combining the EPM and
93 DFN approaches in order to optimize SWI simulation in karst systems.

94 Experimental sandbox setups have been used successfully over the years to investigate saline
95 intrusion in homogeneous (Robinson et al., 2016, Li et al., 2018; Takahasi et al., 2018; Armanuos et
96 al., 2019) and heterogeneous aquifers (Konz et al., 2008; Vithanage et al., 2012; Dose et al, 2014; Liu
97 et al., 2013; Mehdizadeh et al., 2014;). Multiple studies have indicated the significant impact of
98 heterogeneity on SWI dynamics (Abdulhalik and Ahmed, 2017a, b; Houben et al., 2017). Nevertheless,
99 saline intrusion in fractured aquifers has never been studied on a sandbox setup. Laboratory
100 investigations of solute transport in fractured systems have been conducted by Li (2004) and Faulkner
101 et al. (2009). In both studies, glass beads were utilized to recreate the porous medium. In the first
102 investigation, the aquifer had a central vertical conduit running from top to bottom, while in the
103 second study a single horizontal channel was present at the aquifer's lower boundary.

104 The current study employed sandbox experiments coupled with numerical modelling to study saline
105 intrusion in fractured coastal systems. To the best of the authors' knowledge, this is the first time that
106 SWI in fractured aquifers has been methodically investigated on a laboratory scale. Head induced
107 saline intrusion was initiated across six fractured aquifers, with the experimental data obtained being
108 benchmarked against measurements done on a homogeneous system. The study quantified the
109 impact that the position, size and orientation of individual fractures have on the standard saltwater
110 intrusion characteristics of toe length (TL) of the intruding wedge and the width of the mixing zone
111 (WMZ). Unlike in previous investigations, these intrusion variables were measured with high accuracy,
112 utilizing advanced image analysis techniques (Robison et al., 2015; Etsias et al., 2020 a). The laboratory
113 data were successfully recreated using numerical simulations so that the sandbox measurements were
114 supplemented by an in-depth sensitivity analysis, further expanding the conclusions derived by the
115 experimental observations. This study constitutes a contribution towards outlining the basic
116 mechanisms of SWI in fractured porous media, while its results could assist in the successful
117 management and protection of real-world coastal fractured aquifers.

118 **2. Experimental setup and its correlation with real world fractured aquifers**

119 The sandbox apparatus depicted in Figure 1a was employed in deriving the experimental data of this
120 study. It comprised two cylindrical tanks and a thin central viewing chamber of dimensions 0.38 m ×
121 0.15 m × 0.01 m. Red food colouring (E129 Allura Red AC Granular) was mixed with saltwater at a
122 concentration of 0.15 g/L. The density of the dyed saltwater was equal to 1025 kg/m³. Transparent
123 general-purpose glass beads with a diameter of 1090 μm, supplied by Whitehouse Scientific®, were
124 siphoned into the viewing chamber to recreate the porous medium of the aquifer. Freshwater was
125 constantly introduced at the bottom of the left chamber, while the right cylinder was filled with dyed
126 saltwater. The water level at each side chamber was regulated using two adjustable height outflow
127 outlets (Figure 1b) placed at the top of the side tanks, while it was constantly monitored by two
128 ultrasonic sensors (Figure 1c) with an accuracy of 0.2 mm. Two laser-cut mesh screens were utilized
129 to secure the glass beads inside the central chamber. Experimental investigations were performed in
130 a dark room; the necessary illumination was provided by two Camtree® 600 LED panels. A Nikon D850
131 Digital SLR Camera recorded the experimental images at five-minute intervals. The laboratory
132 apparatus was presented in detail by Robinson et al. (2015).

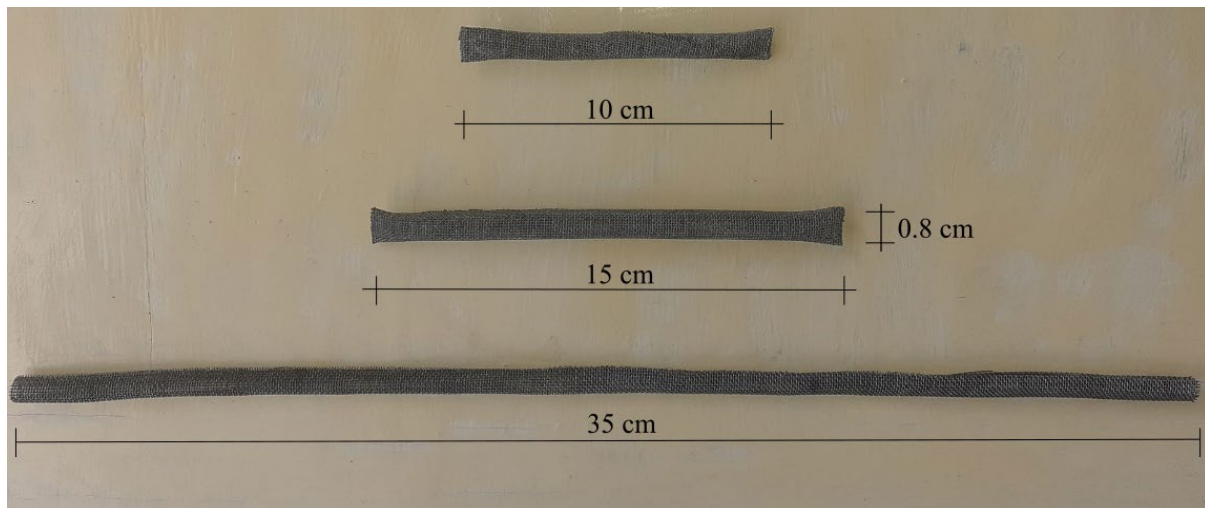


133

134 Figure 1: a) 3-dimensional representation of the utilized experimental setup alongside photos of the
 135 b) adjustable overflow outlets and c) ultrasonic sensors

136 Stainless steel mesh screens with 1 mm openings were used to recreate the aquifer fractures. The
 137 mesh was manually moulded around a steel rod with an 8 mm diameter. Three hollow mesh cylinders
 138 with lengths of 35 cm, 15 cm and 10 cm were created (Figure 2). These structures were placed at pre-
 139 determined positions inside the sandbox's viewing chamber, recreating six unique laboratory
 140 fractured aquifers (Figure 3). In four cases, the fractures were horizontally oriented, while in the
 141 remaining two they were vertical in orientation. In the horizontally fractured aquifers, the cylindrical
 142 mesh was placed at the middle of the viewing chamber's height. The first aquifer included the longer
 143 fracture (35 cm), crossing most of the porous medium, while in the remaining three test configurations
 144 the 15 cm long cylinder was located centrally on the test area mid-point and at the left and right
 145 aquifer boundaries respectively. In the vertically fractured test cases, the 10 cm long fracture was
 146 placed at the test mid-point and at a distance of 10 cm from the right boundary of the aquifer. In the

147 interests of clarity, the laboratory aquifers will henceforth be referred to according to their
148 corresponding fracture position: horizontal-long (Figure 3b), horizontal-middle (Figure 3c), horizontal-
149 left (Figure 3d), horizontal-right (Figure 3e), vertical-middle (Figure 3f) and vertical-right (Figure 3g).
150 Testing was also conducted on a homogeneous aquifer, without any fractures, that served as a
151 benchmark case for this investigation (Figure 3a).



152
153 Figure 2: The three hollow mesh cylinders employed to recreate the aquifer fractures during
154 experimental measurements

155 This is the first ever laboratory study of saltwater intrusion in fractured aquifers. In the absence of any
156 similar investigations, sandbox studies of SWI, alongside fieldwork investigations of saline intrusion in
157 real world fractured hydrological systems, were employed to validate the suitability of the
158 experimental setup in the approximation of fractured aquifers. Glass beads of comparable size have
159 been utilized in multiple laboratory studies of saline intrusion. Zhang et al. (2001 and 2002) used glass
160 beads with a diameter of 725 μm , Goswami and Clement (2007) and Chang and Clement (2013)
161 employed 1.1 mm wide glass spheres, while Konz et al. (2009) studied SWI in heterogeneous aquifers
162 using glass beads of sizes varying between 0.6 mm and 2.2 mm. An intrinsic flow test on the
163 experimental domain allowed calculation of the permeability of the porous media using Darcy's law.
164 Permeability and porosity were equal to $1.83 \times 10^{-9} \text{ m}^2$ and 0.385, respectively. These results agreed
165 with the values reported by Robinson et al. (2015, 2016) for the same laboratory apparatus. Glass
166 beads of similar size have been utilised by Abdelgawad et al. (2017) and Abdoulhalik and Ahmed
167 (2017a, b) in experimental investigations of saltwater upconing.

168 The permeability measured in the laboratory porous medium was compared with values of
169 permeability in real world fractured aquifers. De Fillipis et al. (2016) presented a study of saline
170 intrusion in a karstic coastal aquifer in the Taranto area of Northern Italy. The lowest hydraulic
171 conductivity values (of the order of 10^{-4} m/sec) were calculated inland and along the extremely
172 western coast, while the central and eastern coastline part of the aquifer had a hydraulic conductivity
173 of 0.1 m/sec, or approximately $9.07 \times 10^{-9} \text{ m}^2$. In an investigation of SWI in a fractured aquifer in Crete,
174 Greece (Steiakakis et al., 2016) a permeability of $1.35 \times 10^{-10} \text{ m}^2$ was reported for the limestone portion
175 of the hydrological system. A comparable permeability value was identified for the fracture-karst
176 aquifer in Zhoushuizi district of Dalian City in northern China (Zhao et al., 2016). Xu and Hu (2017)
177 introduced a numerical model for simulating seawater intrusion to a coastal karst aquifer with a
178 conduit system. A porous medium permeability equal to $2.4 \times 10^{-9} \text{ m}^2$ was derived from previous field
179 scale studies (Loper et al., 2005; Kincaid and Werner, 2008; Xu et al., 2016) of the Floridan aquifer in

180 the Woodville Karst Plain, in Florida (USA). Finally, a study of the coastal carbonate aquifer in western
181 Cuba (Hernandez-Diaz et al., 2019) reported an aquifer permeability equal to $1.1 \times 10^{-9} \text{ m}^2$. The real-
182 world fractured aquifer permeabilities values presented in this paragraph, have a difference of less
183 than an order of magnitude from the permeability of the utilised glass bead medium. This validates
184 the current laboratory setup as an acceptable approximation of saltwater dynamics in fractured
185 hydrological systems.

186 The sandbox fracture permeability was determined via sensitivity analysis, conducted using the
187 numerical model presented in the next session. The fractures (steel mesh tube) were approximately
188 100 times more permeable than the surrounding porous medium (glass beads). This permeability ratio
189 is in agreement with fieldwork investigations of fractured aquifers. McAllister et al. (2018) in their
190 study of the Seaford and Lewes Nodular Chalk formations in the UK, reported intrinsic fracture
191 permeabilities 100 to 150 times larger than the rest of the aquifer. Similarly, Xu and Hu (2017)
192 identified a conduit network that was 250 times more permeable than the porous medium of the
193 Floridian karst aquifer.

194 In the initial stage of each experiment, freshwater occupied the whole aquifer. Saltwater was
195 introduced to the system by applying a hydraulic head difference $dH = 6 \text{ mm}$ between the two side
196 chambers of the laboratory apparatus. After the stabilization of the saline wedge, a new phase of
197 saline intrusion was initiated by modifying the hydraulic head difference from 6mm to 4mm. In the
198 final part of the experiments saltwater retreat was generated by applying a steeper hydraulic gradient
199 ($dH = 4 \text{ mm} - 5 \text{ mm}$) in the aquifer. The resulting hydraulic gradients were similar to those documented
200 for various real world aquifers (Attanayake and Sholley, 2007; Ferguson and Gleesson, 2012). The total
201 measurement time varied between 160 and 180 minutes in the seven investigated cases, depending
202 on the time needed for the saline wedge to stabilize after changing the head difference. During the
203 experiment, water in the right cylinder was monitored using a YSI Professional Plus Instrument (Pro
204 Plus) water quality meter. The salinity values for all six experiments, obtained at five-minute intervals,
205 demonstrated minimum variation (Figure S1 of the supplementary material) indicating that water
206 salinity in the right chamber did not reduce due to the outflow of freshwater through the porous
207 medium.

208 The automated image analysis algorithms introduced by Robinson et al. (2015) and Etsias et al. (2020
209 a and b) were utilized to successfully post-process the acquired experimental images. An Artificial
210 Neural Network (ANN) with a single hidden layer, consisting of 10 neurons, was employed to recreate
211 saltwater concentration fields from the corresponding Light Intensity (LI) values of the laboratory test
212 images (Figure 3). Two variables were calculated utilizing image processing to quantify the effect of
213 fractures on aquifer saline intrusion: TL and WMZ. TL was deemed equivalent to the horizontal
214 distance between the saltwater boundary of the laboratory aquifer and the point where the 50 %
215 saltwater concentration isoline intersected with the bottom of the aquifer, while WMZ equated to the
216 average vertical distance between the 25 % and 75 % saltwater concentration isolines. The automated
217 calculation of these variables was described in detail by Robinson et al. 2015.

218 **3. Numerical modelling**

219 The experimental saltwater flow fields were simulated utilizing a 2-dimensional, saturated-
220 unsaturated, variable-density groundwater flow model in SUTRA (Voss and Provost, 2010), where the
221 fractures were modelled discretely as 2-dimensional elements (DFM 2D/2D). Permeability and
222 porosity were equal to $1.83 \times 10^{-9} \text{ m}^2$ and 0.385, respectively. Fracture permeability was determined,
223 via sensitivity analysis, being 100 times larger than that of the matrix, while the porosity of fractures
224 was set equal to 1. Since the size of the mesh tubes was approximately 0.8 cm, and turbulent flow is

225 not relevant inside fractures with a diameter of less than 1 cm (White et al., 2019), Darcian flow was
 226 considered in both media (matrix and fracture). The highest velocity inside the fracture (0.0337 m/sec)
 227 was recorded in the horizontal-long case for an applied hydraulic gradient of $DH = 6\text{mm}$. This results
 228 in a low Reynolds number, equal to 256.3, clearly indicating laminar flow inside the mesh tube. The
 229 x-z two-dimensional flow model, with dimensions of $0.38\text{ m} \times 0.134\text{ m}$, was discretized with a finite
 230 element mesh with quadrangular elements of a size of $1.22 \times 10^{-3}\text{ m}$. The element size complied with
 231 the Peclet number criterion (Voss and Souza 1987), while the applied dispersivity values were within
 232 the range introduced by Abarca and Clement (2009). Unsaturated flow was simulated utilizing the van
 233 Genuchten equation (van Genuchten, 1980). Van Genuchten parameters were determined according
 234 to values measured in laboratory testing for glass beads of comparable size (Benson et al. 2014,
 235 Sweijan et al. 2017). A hydrostatic freshwater ($C = 0\%$) boundary condition was applied on the left
 236 boundary while a hydrostatic saltwater ($C = 100\%$) boundary condition was implemented on the right
 237 boundary of the aquifer. A necessary step towards the successful numerical recreation of the
 238 laboratory data was the precise identification of the position of the fracture in each test case. This was
 239 achieved through image analysis of the original experimental figures and the subsequent distribution
 240 of all the corresponding elements of the numerical model in either the matrix or the fracture medium.
 241 The simulation time for each test case was similar to the equivalent experimental duration, while the
 242 time-step was equal to 1 sec. Model parameters are listed in Table 1. A comparison between
 243 experimental and numerical results is presented in Figures 3 and 4.

244 Table 1: Summary of the discrete fracture matrix model parameters

Input parameters	Values
Domain length (m)	0.38
Domain height (m)	0.134
Element size (m)	1.22×10^{-3}
Freshwater density (kg/m^3)	1000
Saltwater density (kg/m^3)	1025
Freshwater head (m)	0.134
Saltwater head (m)	0.128 -0.13
Porous medium (matrix)	
Permeability (m^2)	1.83×10^{-9}
Porosity	0.385
Fracture	
Permeability (m^2)	1.83×10^{-7}
Porosity	1
Longitudinal dispersivity (m)	10^{-3}
Transverse dispersivity (m)	3.75×10^{-5}
Van Genuchten parameters	
α (1/Pa)	8.45×10^{-4}
n	4.5
Time step (sec)	1
Simulation time (min)	
Simulating exp. aquifers	50-70
Sensitivity Analysis	80

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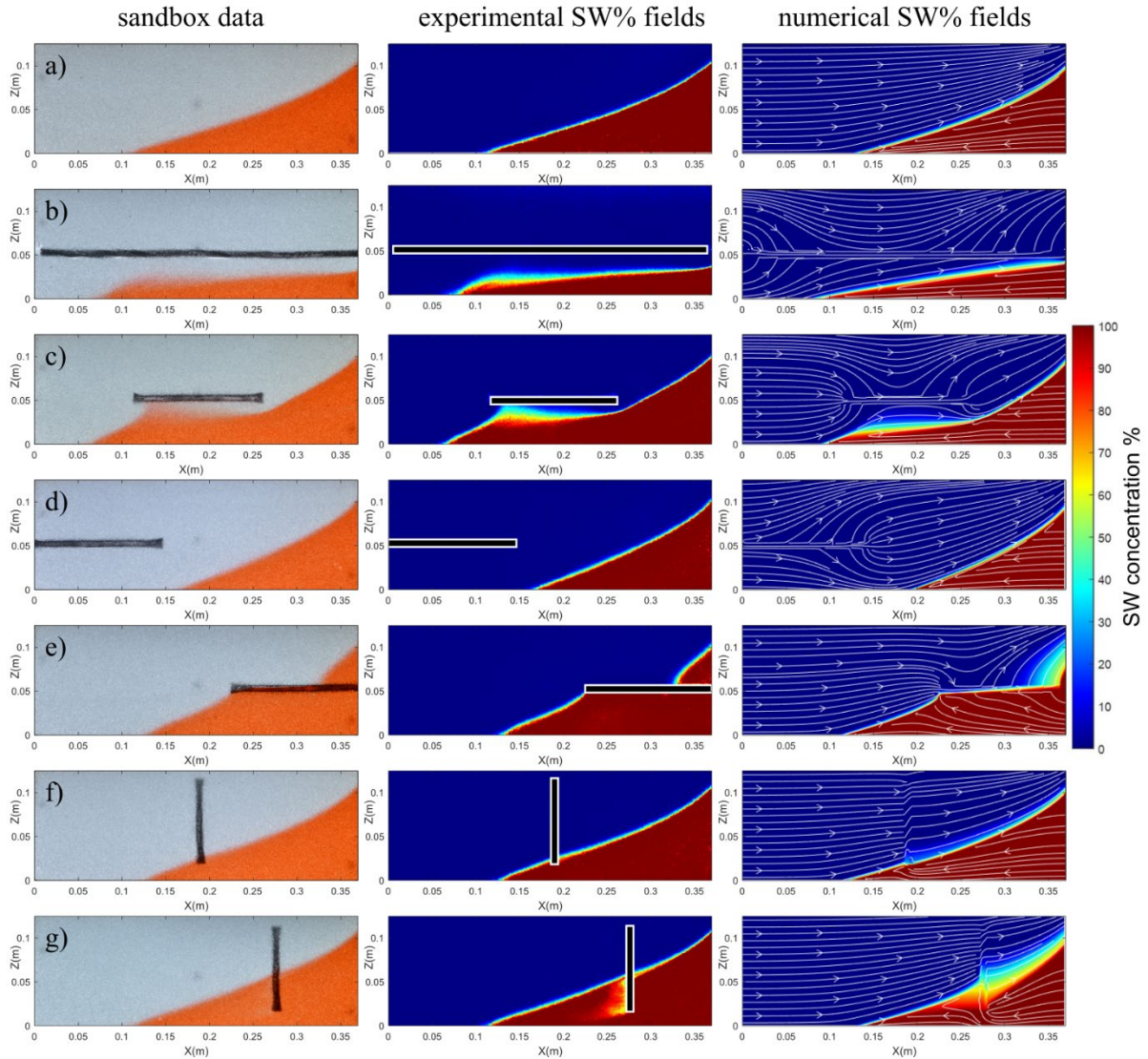
246 The conclusions about the impact of fractures on SWI, derived from experimental observations, were
 247 investigated further through a rigorous sensitivity analysis using the aforementioned numerical
 248 model. In total four different model setups were created, in three of them the fractures' orientation

249 was horizontal while in the fourth one it was vertical. This sensitivity analysis expanded the study's
250 findings on the effect of the fractures' position, orientation and length on aquifer saltwater dynamics.

251 **4. Results and discussion**

252 **4.1 Experimental results**

253 The experimental saltwater flow fields that were generated inside the sandbox setup, by a hydraulic
254 head difference between the left and right aquifer boundaries equal to $dH = 4\text{mm}$, are presented in
255 Figure 3. The saline concentration fields derived by experimental image analysis and numerical
256 modelling (SUTRA) are displayed alongside them. Alongside saltwater concentration, SUTRA calculates
257 the flow velocity vectors at each individual element. The flow streamlines plotted on top of the
258 numerical saltwater concentration fields in the third column of Figure 3 were generated by feeding
259 the numerical flow velocities into the streamslice.m built-in MATLAB equation. This is a standard
260 function developed by Mathworks and included in the MATLAB R2018a release. The results
261 demonstrated that the position, size and orientation of discontinuities have a significant impact on
262 the extent of saltwater intrusion, as well as the shape of saline wedges in fractured systems. The toe
263 length has been a long established metric of the extend of saline intrusion. Nevertheless, in cases like
264 the horizontal-long aquifer (Figure 3b), even though the observed TL value was relatively big, the
265 actual volume of the aquifer occupied by saltwater was disproportionately small. On that account, the
266 uniquely shaped freshwater - saltwater interfaces, generated by the presence of high permeability
267 fractures, required the use of a supplementary variable to help quantify and assess the successful
268 study of SWI. This novel variable, now termed as the saline volume fraction, corresponded to the
269 fraction of the porous medium in which saltwater concentration was higher than 90%. It was
270 calculated with high accuracy by the aforementioned image analysis algorithms and has been utilized
271 for the first time in this study.

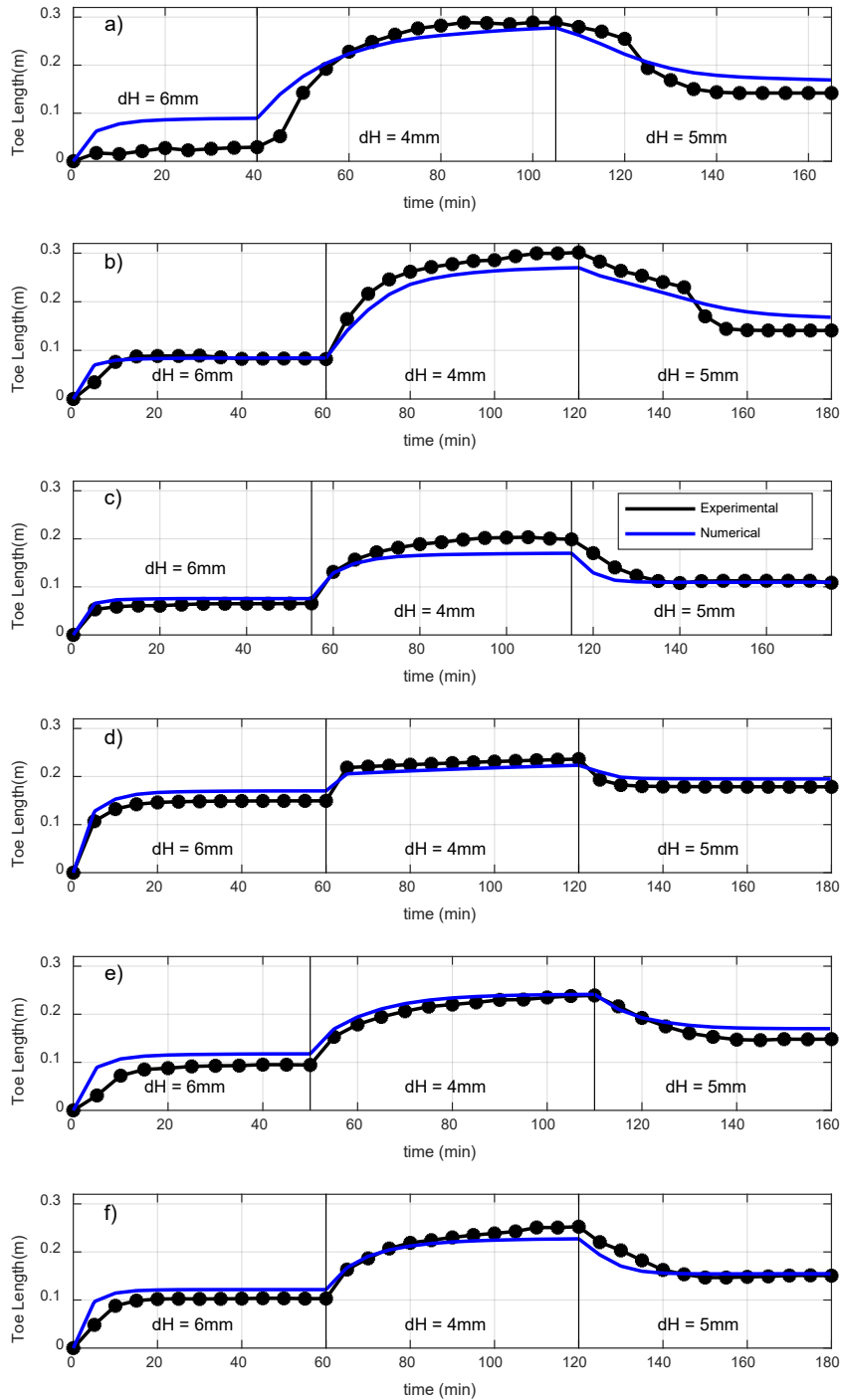


272

273 Figure 3: Photos of laboratory induced saline intrusion alongside experimental and numerical
 274 saltwater concentration fields in the seven investigated aquifers. From top to bottom the investigated
 275 cases are a) homogeneous aquifer, b) horizontal-long, c) horizontal-middle, d) horizontal-left, e)
 276 horizontal-right, f) vertical-middle and g) vertical-right fractured aquifers.

277 The numerical model successfully recreated the laboratory results for all three distinct intrusion and
 278 retreat ($dH = 6 \text{ mm} - 4 \text{ mm} - 5 \text{ mm}$) experimental phases, for all the tested cases (Figure 4). In all
 279 cases, there is excellent temporal agreement between the measured values of TL and those obtained
 280 from the numerical simulations. Furthermore, it proved that using a dual-porosity model to simulate
 281 the sandbox data was a valid choice, and that this model can be safely employed to both interpret the
 282 physical mechanisms affecting SWI in the specific laboratory aquifers, as well as to further expand any
 283 conclusions derived from their study. As expected, the experimental TL was negatively correlated with
 284 the applied hydraulic gradient. The difference between the three steady state TL values (Table 2) was
 285 more significant in the horizontal-long aquifer, where $\Delta TL_{6-4\text{mm}} = +25.95 \text{ cm}$ and $\Delta TL_{4-5\text{mm}} = -11.83 \text{ cm}$,
 286 while it was minimum for the horizontal-right case, $\Delta TL_{6-4\text{mm}} = +8.7 \text{ cm}$ and $\Delta TL_{4-5\text{mm}} = -5.78 \text{ cm}$. This
 287 demonstrated that, the size and position of the discontinuity can significantly affect the impact of
 288 hydraulic head difference on the extend of saline intrusion. Every change in the hydraulic gradient was
 289 followed by an initial, rapid adjustment of the toe length, which subsequently stabilized. The time to

290 reach steady state in each intrusion and retreat phase did not vary much between the six investigated
291 fractured cases. This is in agreement with the laboratory findings of Robinson et al. (2016) indicating
292 that this time depends solely on the permeability of the porous medium, which was the same in all
293 laboratory setups.



295 Figure 4: Experimental and numerical transient toe length values of the a) horizontal-long, b)
 296 horizontal-middle, c) horizontal-left, d) horizontal-right, e) vertical-middle and f) vertical-right
 297 fractured aquifers

298 Table 2: Steady-state toe length values of the seven investigated laboratory aquifers, and the
 299 corresponding relative change in TL generated during the experimental intrusion (dH = 6 – 4 mm) and
 300 retreat (dH = 4 – 5 mm) phases

Aquifer TL (cm)	dH = 6 mm	dH = 4 mm	dH = 5 mm	$\Delta TL\%_{6-4mm}$	$\Delta TL\%_{4-5mm}$
homogeneous	12.10	25.38	14.29	+109.75	-43.70
horizontal-long	2.94	28.89	17.06	+882.65	-40.95
horizontal-middle	8.22	30.14	14.82	+266.67	-50.83
horizontal-left	6.57	20.34	10.87	+209.59	-46.56
horizontal-right	14.94	23.64	17.86	+58.23	-24.45
vertical-middle	9.44	23.91	14.82	+153.28	-38.02
vertical-right	10.30	25.25	15.06	+145.15	-40.36

301

302 The three variables outlining the effect of fractures on saline intrusion: TL, WMZ and the percentage
 303 of aquifer saline volume, were quantified. Their values were benchmarked against the SWI
 304 characteristics in the homogeneous laboratory aquifer with the same permeability and porosity values
 305 (Table 3). It was established that, in comparison to the homogeneous case, TL was significantly longer
 306 for the horizontal-long and horizontal-middle aquifers, by 13.8 % and 18.7 % respectively, while it was
 307 about 19.9 % and 6.9 % shorter for the horizontal-left and horizontal-right cases. No quantifiable
 308 deviation in the TL values was observed for the systems with vertical fracture orientation. The mixing
 309 zone in all the fractured cases was wider than that in the homogeneous aquifer. In particular, for the
 310 horizontal-long, horizontal-middle and vertical-right systems, WMZ was more than double the width
 311 of the benchmark case (220 % ~ 285 % of the homogeneous WMZ), while for the remaining three
 312 fractured aquifers it was between 21 % and 52 % wider. The total volume fraction of the aquifer
 313 occupied by the saline wedge was significantly smaller for the horizontal-long and horizontal-left
 314 cases, being equal to just 49.1 % and 78.2 % of the equivalent saline volume fraction in the
 315 homogeneous case. In the remaining fractured aquifers saline volume fraction did not deviate
 316 significantly, ranging between 93.5 % and 105.9 % of the benchmark case.

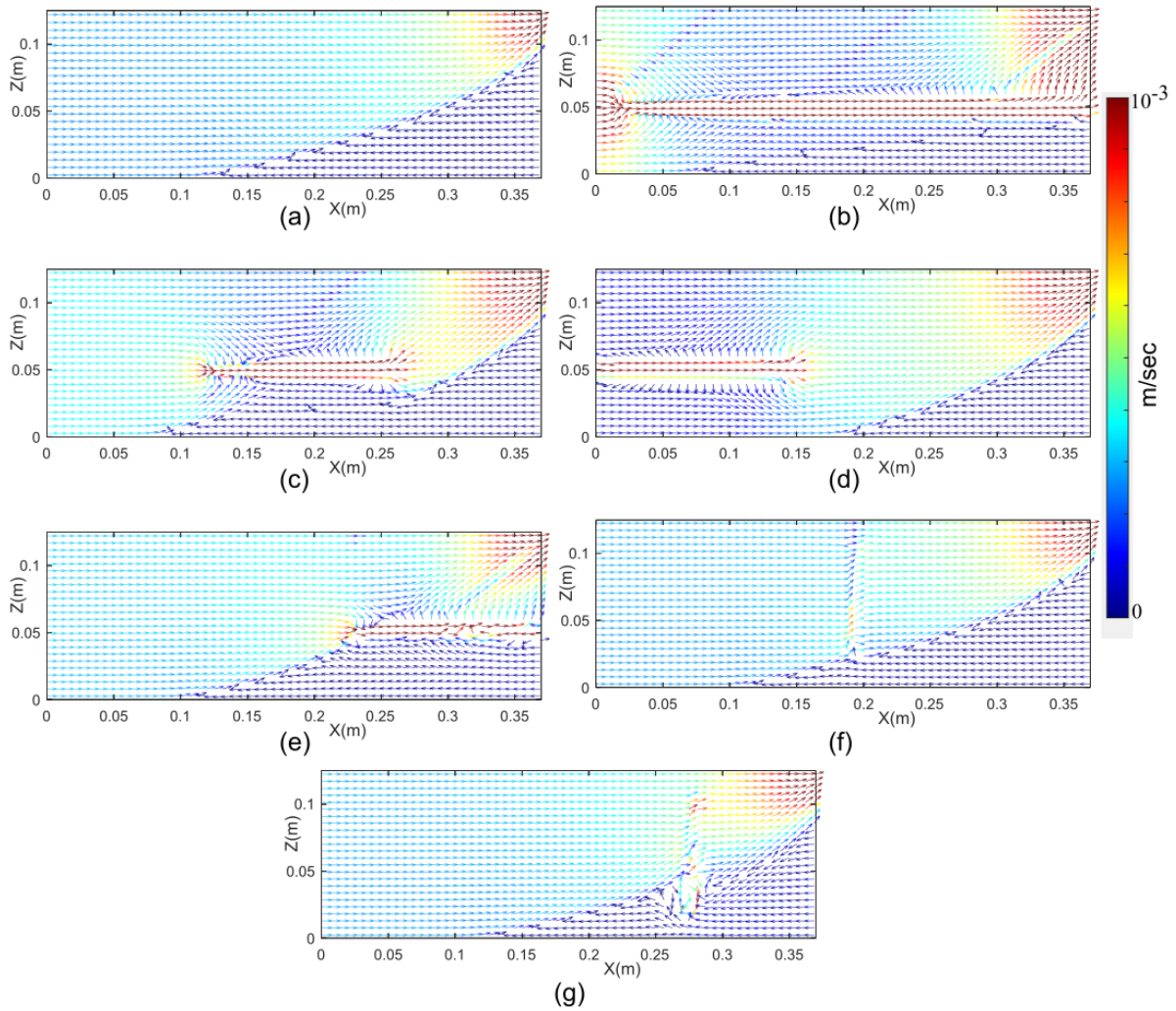
317 Table 3: Comparison between the experimental values of toe length, width of the mixing zone and
 318 percentage of saline volume of the six investigated fractured aquifers and the ones of the equivalent
 319 homogeneous aquifer

Fractured aquifer	Fractured / Homogeneous (%)		
	TL	WMZ	Saline Volume
horizontal-long	113.82	220.74	49.11
horizontal-middle	118.76	285.01	105.98
horizontal-left	80.15	120.92	78.27
horizontal-right	93.12	151.93	93.47
vertical-middle	97.22	150.66	97.89
vertical-right	99.51	244.95	100.88

320

321 To assess the interpretation of the acquired experimental data and to outline the basic physical
 322 mechanisms behind the impact of fractures on SWI dynamics, the velocity vector fields of the seven
 323 investigated hydraulic systems were recreated using SUTRA (Figure 5). In Figure 5, the magnitude of

324 flow velocities is expressed with varying colour instead of varying vector lengths, as it is commonly
325 demonstrated (e.g. Abdoulhalik and Ahmed 2017a). This was deliberately chosen by the authors to
326 assess velocity visualization, since freshwater, both inside the fractures and the porous medium, flows
327 with a velocity that is at least an order of magnitude greater than that of the saltwater. Figure 5a
328 depicts the typical flow velocity distribution occurring during SWI on a homogeneous aquifer. This
329 constitutes a well-documented mechanism; the less dense freshwater overtops saltwater and
330 outflows from the upper right area of the quasi 2-dimensional system. Flow velocity is considerably
331 higher for the freshwater than the saltwater and its magnitude peaks around the outflow zone. In the
332 simulated aquifer depicted in Figure 5b (horizontal-long), two distinct zones of higher velocity were
333 observed at the entrance and exit of the long fracture. The presence of the fracture caused the
334 majority of freshwater mass transport to occur through it, while at the same time limiting the flow of
335 freshwater in the rest of the system, as depicted by the smaller velocity (dark blue) vectors around
336 the central mesh cylinder. This absence of freshwater at the lower part of the aquifer resulted in longer
337 saline intrusion near the aquifer's bottom. Saltwater did not intrude over the fracture, accounting for
338 the relatively small portion of the aquifer being occupied by the saline wedge (Table 3). The horizontal
339 - middle aquifer (Figure 5c) constituted a hydraulic system similar to the horizontal – long one, with
340 the exception of a shorter fracture. Yet again, two distinct zones of higher freshwater velocity were
341 documented at the fracture's edges. The upward movement of freshwater towards the fracture once
342 more augmented SWI. Freshwater outflow from the fracture's right edge contributed to the saline
343 wedge's distinct shape. Lu et al. (2013) indicated that according to conservation of mass, separation
344 of flow streamlines along the freshwater - saltwater interface results in the widening of the mixing
345 zone. This separation, observed in the areas directly underneath the horizontal discontinuities (Figures
346 3b and 3c), alongside the small difference between the flow velocities of the two liquids at the same
347 zone, attributed for the distinctively wider mixing zone in these two systems. The saline volume
348 fraction of the horizontal - middle system was equal to 105.9 % of the saline volume in the benchmark
349 homogeneous case.



350

351 Figure 5: Flow velocity vector fields, generated by a hydraulic head difference of 4 mm for the a)
 352 horizontal-long, b) horizontal-middle, c) horizontal-left, d) horizontal-right, e) vertical-middle, f)
 353 vertical-right fractured and the g) homogeneous aquifer

354 In the horizontal-left (Figure 5d) and horizontal-right (Figure 5e) laboratory aquifers the same fracture
 355 was placed at diametrically opposite positions, resulting in two distinct saltwater concentration
 356 distributions. In the horizontal-left system the majority of freshwater inflow into the porous medium
 357 occurred through the fracture. Two distinct velocity zones were observed in the aquifer: a low velocity
 358 one directly above and below the fracture, and a relatively large area of high velocity vector fields at
 359 the fracture's right edge. The presence of faster moving freshwater at the centre of the system
 360 resulted in the seaward suppression of saltwater, while the saline wedge's shape was comparable to
 361 that of the homogeneous case albeit with a smaller TL (Table 2). In the horizontal-right aquifer, where
 362 the fracture was in direct proximity with the saltwater boundary, the majority of the water outflow
 363 from this discontinuity as depicted by the relatively lower velocities at the upper right edge of the
 364 system. These two zones of water outflow caused by the position of the cylindrical mesh resulted in a
 365 uniquely shaped saltwater wedge. Yet again, the wider mixing zone directly above the fracture should
 366 be attributed to the separation of streamlines. The total saline volume in this system was equivalent
 367 to 93.5 % of the saltwater volume in the homogeneous case.

368 In the experimental aquifers with vertically oriented fractures, the observed steady-state saltwater
369 wedges were similar in shape and size to the one recreated in the homogeneous aquifer. In particular,
370 for the vertical-middle and vertical-right systems, TL equalled to 97.2 % and 99.5 % of the benchmark
371 case, while saline volume corresponded to 97.9 % and 100.9 % of the saltwater in the homogeneous
372 aquifer. The vector fields presented in Figures 5f and 5g indicate that the impact of these fractures on
373 the total distribution of flow was limited. In the vertical-middle case, where the saline wedge never
374 reached the hollow mesh cylinder, the fracture caused a slight variation in the direction and
375 magnitude of the velocities in its interior that had limited effect on the rest of the system. On the
376 other hand, in the vertical-right aquifer where the saltwater wedge extended beyond the fracture's
377 position, the velocity vector fields inside the mesh structure indicate the intensification of the water
378 recirculation that normally occurs inside the saline wedges. This increased water recirculation justifies
379 the widening of the mixing zone around the fracture, documented in both Figures 3g and Table 3.
380 Although this phenomenon was previously reported in purely numerical investigations (Sebben et al.,
381 2015), it was identified on a laboratory scale for the first time in this study.

382 Overall, the experimental data derived from the six fractured aquifers allowed identification of some
383 preliminary trends concerning the impact that individual fractures have on SWI characteristics. It was
384 established that depending on its length and position, horizontally oriented fractures can either
385 augment or suppress saltwater intrusion, while significantly affecting WMZ and the total volume of
386 intruding saltwater. On the other hand, vertical fractures contributed to the widening of the mixing
387 zone, but had a limited impact on the actual length and shape of the intruding wedge. These findings
388 were further expanded in the following section, using a series of sensitivity analysis scenarios.

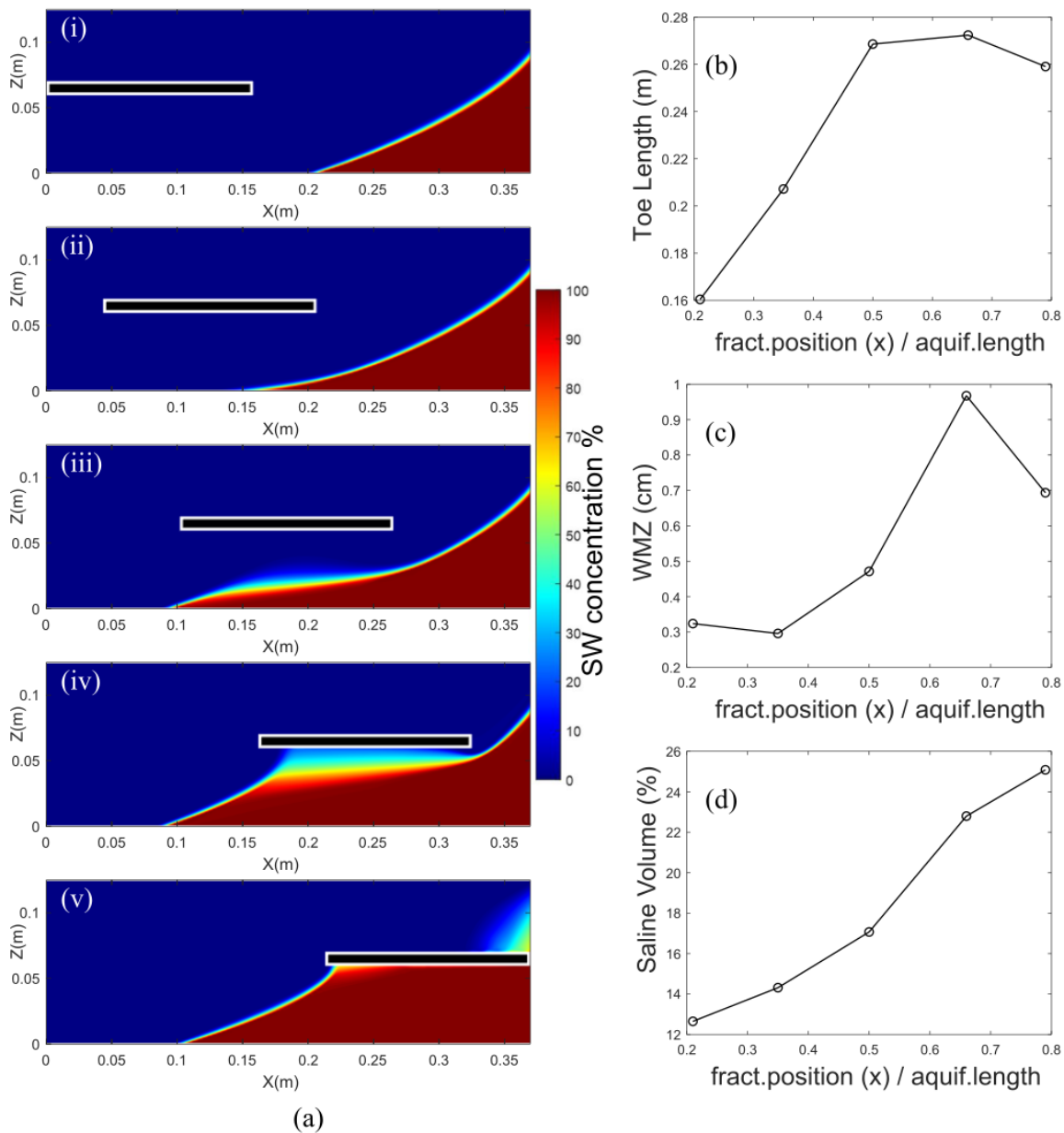
389 **4.2 Sensitivity analysis**

390 The sensitivity analysis presented here comprises four distinct scenarios, each one examining SWI in
391 five numerical aquifers. The utilized model parameters were the same as those used in simulating the
392 experimental saltwater concentration fields (Table 1). Saline intrusion was initiated by applying a
393 hydraulic head difference of 4 mm, while a total run time of 80 minutes ensured that all systems
394 reached quasi steady-state. A fifth set of numerical simulations, where different values of hydraulic
395 gradients were applied to the system, supplemented the findings of the first sensitivity analysis
396 scenario. Unlike the numerical models presented in the previous section in which the fractures' shape
397 and position were determined using image analysis, all fissures present in this investigation had an
398 idealized rectangular shape. In order to assess comparison between experimental data and those
399 derived from the sensitivity analysis, the saltwater concentration fields presented in this section
400 include only the part of the aquifer that it is visible in the laboratory sandbox setup. When determining
401 the depth of discontinuities, the whole height of the numerical aquifer was taken into account, instead
402 of just the part included in the viewing chamber. As a result, discontinuities in this section was placed
403 approximately two centimetres higher than the fractures in the laboratory test cases of the previous
404 section.

405 **4.2.1 Horizontal position of horizontally oriented fractures**

406 In the first sensitivity analysis setup, the impact of position (x) of horizontal fractures on SWI dynamics
407 was investigated. To do so, five aquifers containing a single horizontal fracture with dimensions of
408 15.4 cm × 0.85 cm (126 × 7 elements) were generated. The fractures were placed at the middle of the
409 aquifers' depth. On the horizontal direction, they were either placed adjacent to the system's vertical
410 boundaries (Figures 6a.i and 6a.v), at a distance of 5 cm from them (Figures 6a.ii and 6a.iv) or exactly
411 at the aquifer's centre (Figure 6a.iii). As seen in Figure 6, the closer the fracture was to the freshwater
412 boundary, the more it contributed to the suppression of saline intrusion, while the closer its proximity

413 to the seaward edge, the longer the length of the intruding wedge was (Figure 6b). This is in agreement
414 with the experimental observations presented in section 4.1. The positive correlation between TL and
415 fracture position (x) was disrupted in the case where the fracture was in contact to the sea (right
416 boundary). The adjacency of the high permeability fracture with the saline boundary created two
417 zones of water outflow from the porous medium, significantly altering the underlying flow and mass
418 transport mechanisms. Moreover, the slightly smaller fracture depth in the numerical aquifers of this
419 section is responsible for the difference between the saltwater concentration field depicted in Figure
420 6a.v and that in Figure 3e. Similarly, the mixing zone is widened when the velocity difference between
421 the two fluids is smaller (Figure 6c). This occurred in the low velocity zones underlying the fractures
422 (Figure 6a). The impact of these zones was greater with increasing proximity to the freshwater –
423 saltwater interface. This resulted in a mixing zone, that was three times wider in the aquifer depicted
424 in Figure 6a.iv than in the system illustrated in Figure 6a.i. Finally, a positive correlation was
425 established between the fracture's proximity with the sea boundary and the total volume of the
426 intruding saltwater (Figure 6d), leading to a maximum difference of up to 13 % between the saline
427 volume fraction values of the five test cases.

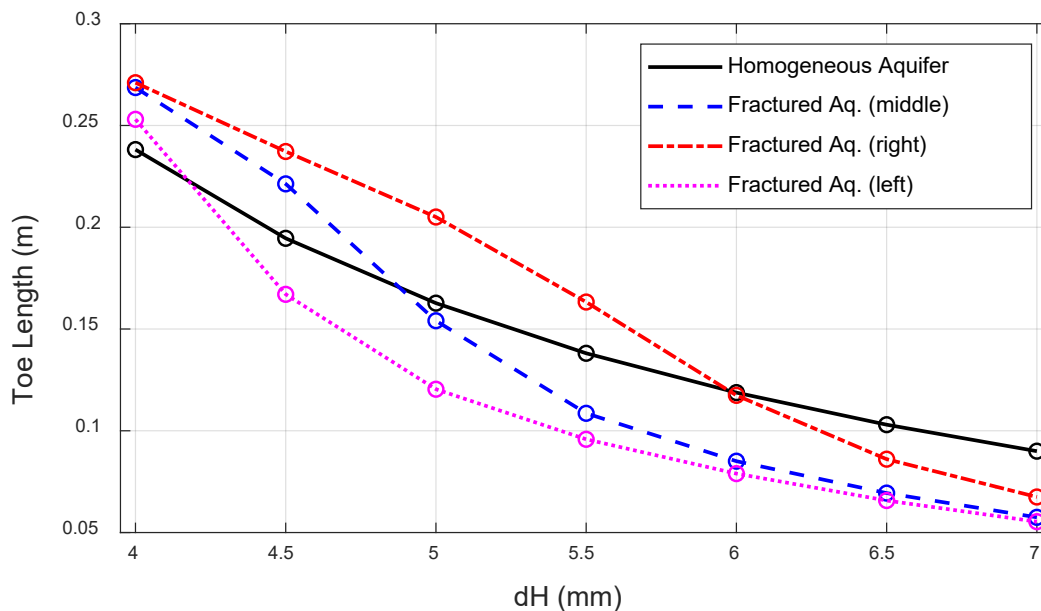


428

429 Figure 6: Simulated a) SW concentration fields alongside their corresponding values of b) toe length,
 430 c) width of mixing zone and d) percentage of saline volume, demonstrating the impact of position (x)
 431 of a horizontal fracture on aquifer SWI

432 In order to identify the main factors determining whether the presence of a discontinuity results in
 433 less or more extended saline intrusion, a supplemental set of numerical simulations were conducted.
 434 In total four aquifers were investigated: the centrally fractured case, presented in Figure 6a.iii; two
 435 fractured aquifers, with a discontinuity at 7.5 cm from the aquifer's left and right boundary; and a
 436 homogeneous system with similar porous medium characteristics. SWI was initiated in these aquifers
 437 by applying seven distinct values of hydraulic head difference ($dH = 4 \text{ mm}, 4.5 \text{ mm}, 5 \text{ mm}, 5.5 \text{ mm}, 6$
 438 $\text{mm}, 6.5 \text{ mm}$ and 7 mm). The generated steady-state TL values are presented in Figure 7. The
 439 numerical results indicated that, depending on the applied hydraulic gradient, the same fracture can
 440 lead to an intruding wedge that is either longer or shorter than the one in the equivalent
 441 homogeneous case. The critical head difference, at which TL was the same for both the fractured and

442 the homogeneous aquifer, was different for each case. Its value was equal to $dH = 4.17$ mm, $dH = 5.97$
 443 mm and $dH = 4.88$ mm for the systems fractured on their left side, right side, and in the middle
 444 respectively. The equivalent TL values for these hydraulic gradients were equal to 22.3 cm, 12 cm and
 445 17.1 cm. When compared to the horizontal position (fracture's centre from the seaward boundary) of
 446 the discontinuity in each system, $x_{\text{left}} = 23.1$ m, $x_{\text{right}} = 13.1$ m and $x_{\text{middle}} = 18.1$ m, a clear linear
 447 relationship between the critical TL value and the discontinuity's position arises. Summarising, the
 448 findings of the numerical investigation proved that the impact of horizontal fractures on the extent of
 449 intrusion depends on both their position and the applied hydraulic gradient. In cases where the
 450 fracture's distance from the seaward boundary was greater than the saline toe length in the equivalent
 451 homogeneous system, for the same hydraulic gradient, the presence of the discontinuity limited the
 452 extent of intrusion, while in cases where this distance was smaller it augmented it. This conclusion has
 453 potentially significant implications for the effective management of real-world fractured aquifers,
 454 since the projected sea-level rise could alter their hydraulic gradient beyond its critical value, thus
 455 leading to a significant and unexpected increase in the extent of saltwater intrusion.

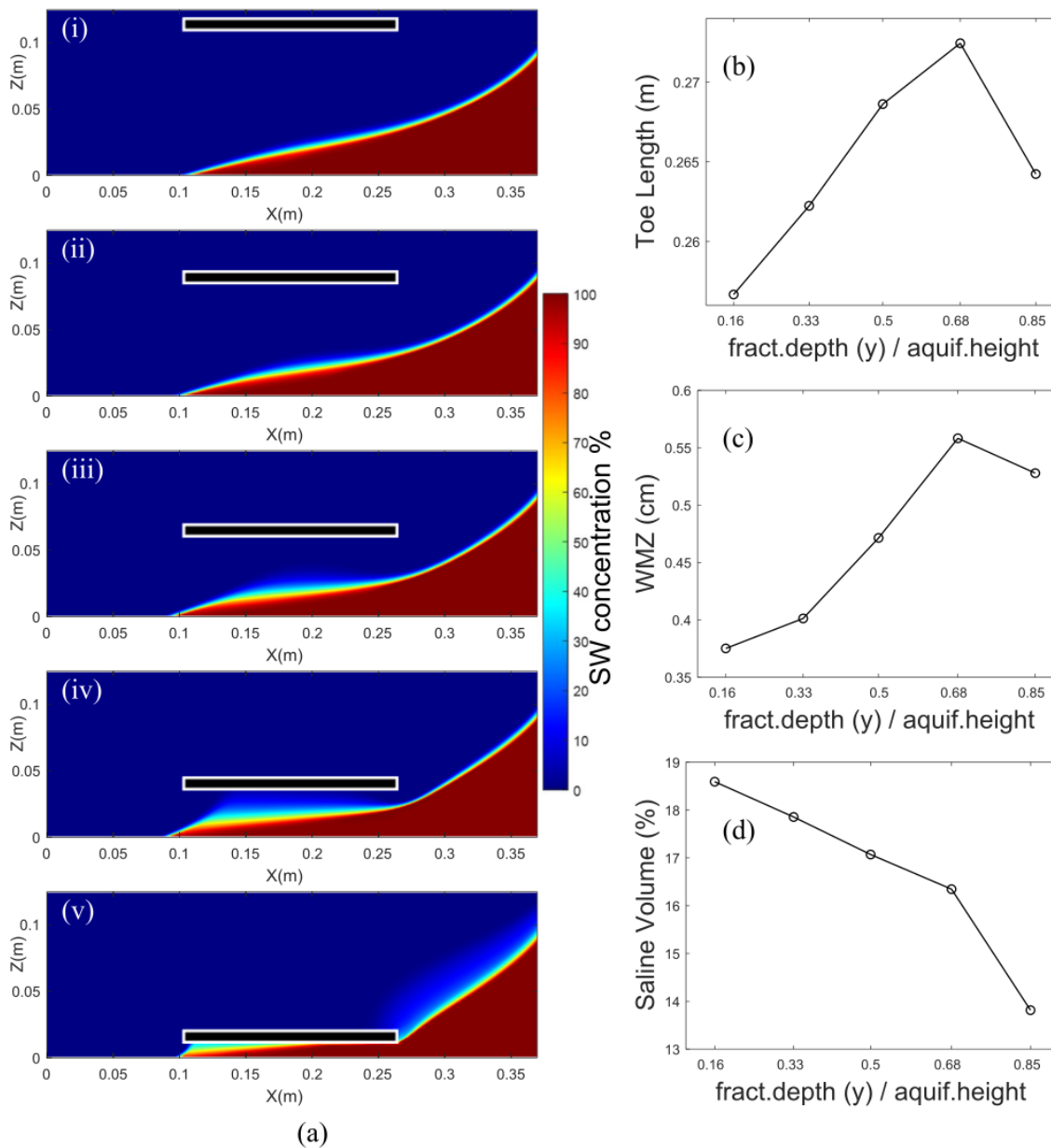


456
 457 Figure 7: Values of saltwater toe length generated by the application of seven distinct hydraulic head
 458 difference values ($dH = 4$ mm – 7 mm) in one homogeneous (black) and three heterogeneous aquifers,
 459 containing a single horizontal fracture at their right (red) and left (purple) side, and their middle (blue)
 460 respectively

461 4.2.2 Depth location of horizontally oriented fractures

462 Five numerical fractured aquifers were utilized to quantify the effect that the depth below the surface
 463 (z) of horizontal discontinuities has on saltwater intrusion. A fracture with the same dimensions to
 464 those described in the previous section was located at a depth from the system's surface equal to $1/6$,
 465 $1/3$, $1/2$, $2/3$ and $5/6$ of the total aquifer width. As seen in Figure 8, the impact of fractures on the
 466 saltwater – freshwater interface became greater the closer the discontinuity was to the impermeable
 467 lower boundary. For the majority of the test cases, the TL was positively correlated to the fracture's
 468 distance from the aquifer's free surface (Figure 8b). Nevertheless, TL variation between the different
 469 aquifers of this setup was much more limited. The shortest TL (Figure 8a.i) equalled to 94.2 % of the
 470 longest TL case (Figure 8a.i.v), translating into a difference of less than 2 cm. In comparison, maximum
 471 toe length difference in the previous sensitivity analysis scenario was approximately 11 cm, or a

472 minimum TL value corresponding to just 58.9 % of the maximum TL. Similar to the previous set of
473 numerical investigations, this trend did not apply to the last test aquifer (Figure 8a.v), where the length
474 of the intruding wedge was shorter than the previous cases. The physical mechanism described for
475 the experimental horizontal – middle aquifer (section 4.1) also applied to these numerical aquifers as
476 well. However, the fracture’s impact proved greater with decreasing distance from the interface of
477 the two liquids. The correlation between fracture depth (z) and WMZ was similar to that reported for
478 the toe length (Figure 8c), while the total volume of intruding saltwater was negatively correlated to
479 the fracture’s distance from the system’s upper boundary (Figure 8d) resulting in an absolute
480 difference of up to 5 % between the investigated aquifers. For all the scenario’s test cases, saline
481 volume constituted a fraction (72.3 % - 97.3 %) of the volume in the equivalent homogeneous case.
482 The current analysis demonstrated that even though the vertical position of a discontinuity has a
483 limited effect on the final length of the intruding wedge, it can significantly influence the extent of
484 intrusion at the upper parts of an aquifer. This could constitute a significant finding in the successful
485 management of real-life fractured systems, where SWI in the upper aquifer region could significantly
486 damage human activities such as farming (Alam et al., 2017).



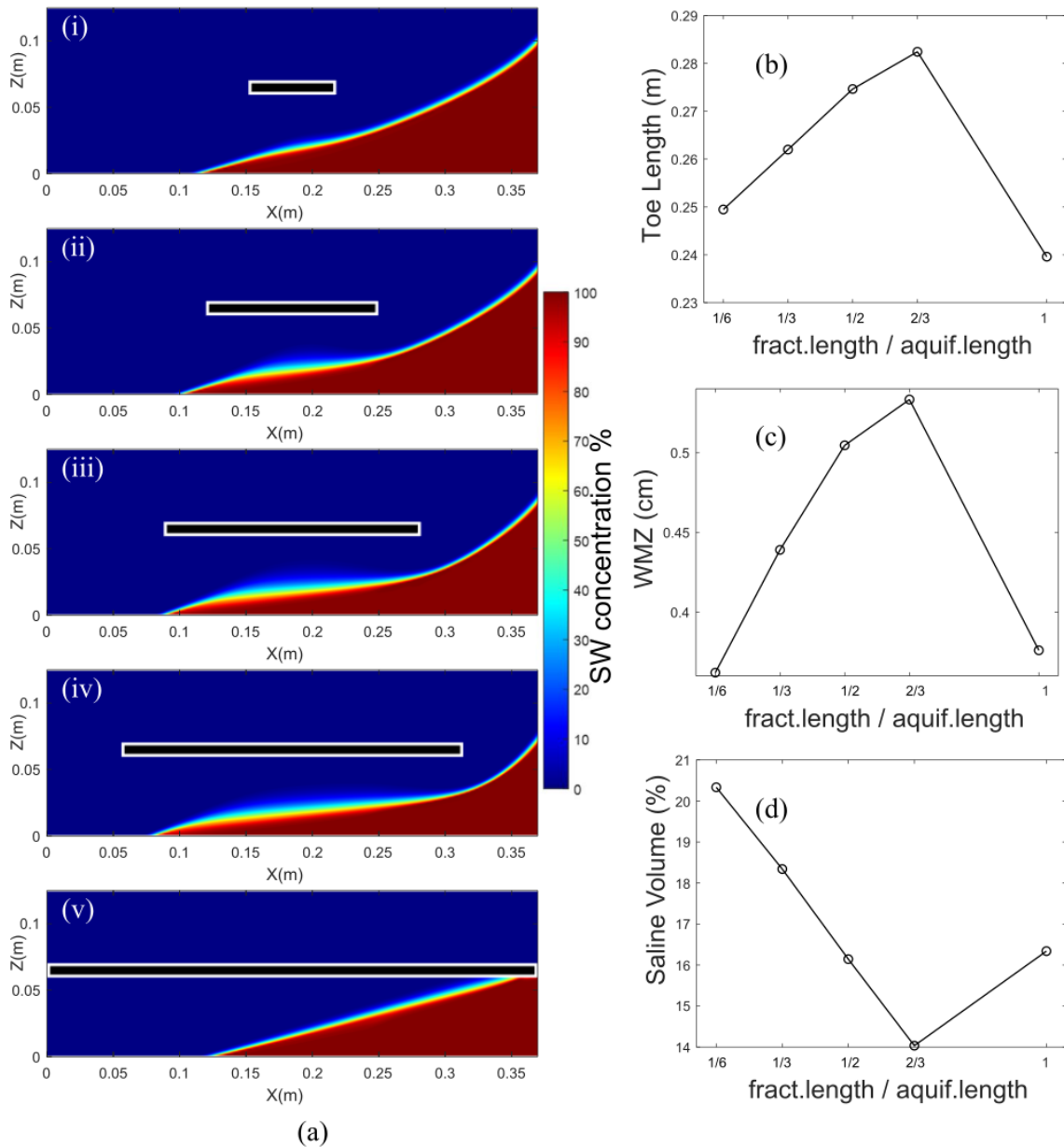
487

488 Figure 8: Simulated a) SW concentration fields alongside their corresponding values of b) toe length,
 489 c) width of mixing zone and d) percentage of saline volume, demonstrating the impact of position (z)
 490 of a horizontal fracture on aquifer SWI

491 4.2.3 Length of horizontally oriented fractures

492 The effect of fracture length on saline intrusion dynamics was studied on the third sensitivity analysis
 493 setup. Aquifers with a single horizontal fracture of a width of 0.85 cm and a varying length equal to
 494 1/6, 1/3, 1/2 and 2/3 of the total aquifer length, as well as an aquifer with a fracture spanning from
 495 one aquifer edge to the other, were tested. With the exception of the last case, the numerical results
 496 indicated a positive relationship between fracture length and length of intrusion (Figure 9b). The
 497 longer the discontinuities, the larger the zones of lower freshwater velocity underneath them, leading
 498 to more space occupied by saltwater. The maximum difference between the TL values generated in
 499 this scenario was less than 5 cm. Both WMZ (Figure 9c) and the total volume of the saline water (Figure
 500 9d) were affected by the fracture length in a similar way. The unique shape of the intruding wedge in

501 the last investigated aquifer (Figure 9a.v), comparable to the interface shape of the horizontal-long
 502 (Figure 3b) experimental system, indicated a fundamentally different flow mechanism. This could be
 503 attributed to the fact that the bulk of the mass transport inside the system occurred exclusively
 504 through its discontinuity, contributing to the limited impact of dispersion effects.

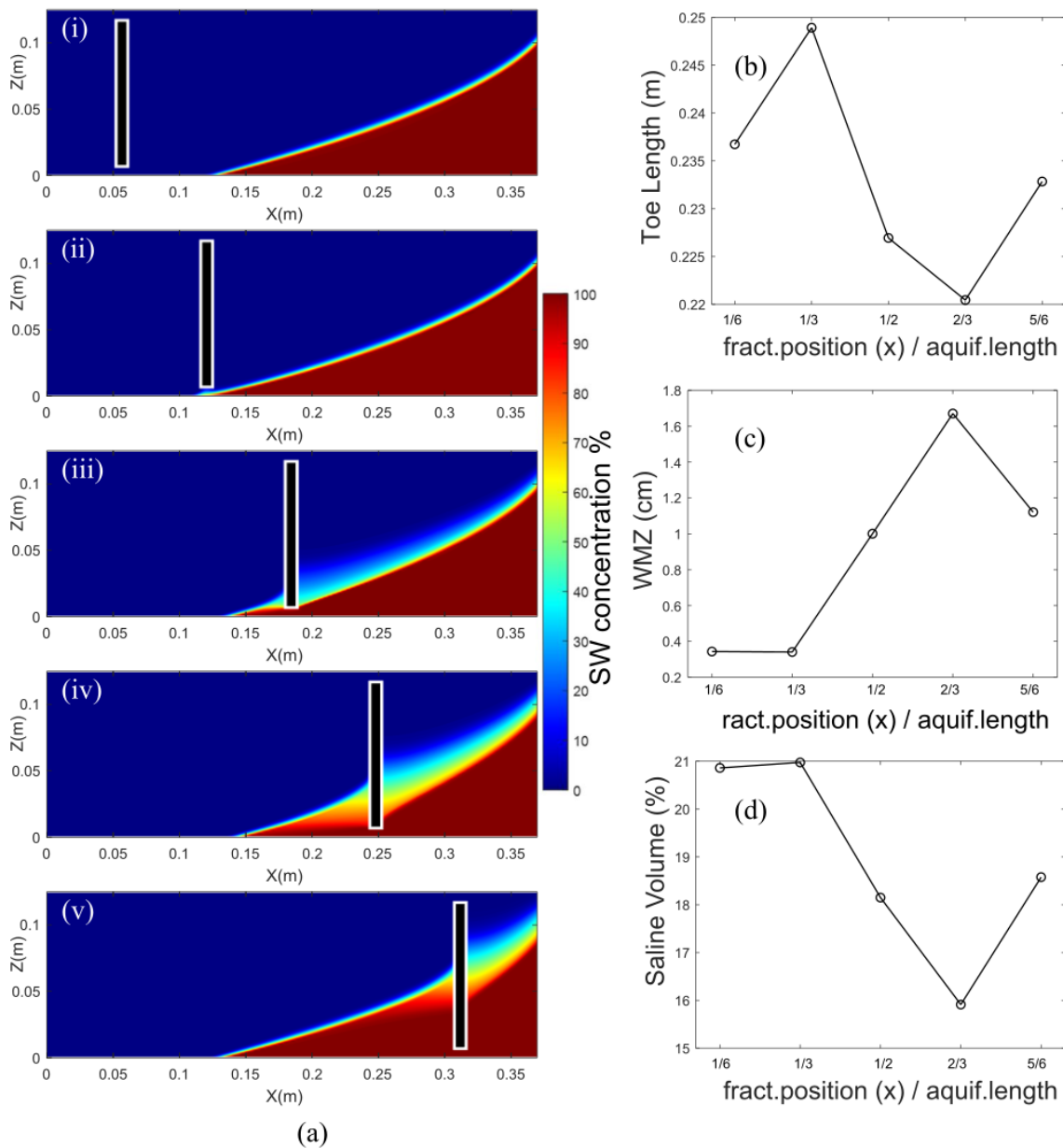


505
 506 Figure 9: Simulated a) SW concentration fields alongside their corresponding values of b) toe length,
 507 c) width of mixing zone and d) percentage of saline volume, demonstrating the impact of a horizontal
 508 fracture's length on aquifer SWI

509 **4.2.4 Horizontal position of vertically oriented fractures**

510 The final sensitivity analysis setup investigated the influence that the position (x) of vertical
 511 discontinuities has on aquifer saline intrusion. Five numerical aquifers with fractures of dimensions
 512 0.85 cm × 11.8 cm placed at different distances from the system's side boundaries were the basis of
 513 this study (Figure 10a). It was shown that the impact of the discontinuities' position on the steady-

514 state length of intrusion was relatively limited, causing a TL variation of less than 12 % between the
515 five tested cases (Figure 10b). On the other hand, the width of the mixing zone was significantly
516 affected by the fractures' location. In aquifers where the intruding saline wedge incorporated the
517 discontinuities, WMZ widened by up to 12 cm in comparison to the cases where the fractures were
518 fully covered by freshwater (Figure 10c). This constituted an increase of approximately 400 %. The
519 correlation between the position of vertical discontinuities and the width of mixing zone is in
520 agreement with the relationship derived by the experimental measurements (section 4.1) as well as
521 the numerical findings of Sebben et al. (2015). The impact of flow recirculation effects inside the
522 discontinuity was more severe on its downstream than its upstream area, i.e. right of the fracture in
523 the investigated numerical setup. This justifies why the increase in WMZ was larger in the numerical
524 aquifer depicted on Figure 10a.iv, where the fracture was approximately at the middle of the intruding
525 wedge. Finally, since no significant alteration of the traditional wedge shape was observed in any
526 aquifer, the trend in the total volume of intruding saltwater was comparable to that of the TL values
527 (Figure 10d). The deviation of saline volume fraction was 5% or less between the investigated
528 fractured systems.



529

530 Figure 10: Simulated a) SW concentration fields alongside their corresponding values of b) toe length,
 531 c) width of mixing zone and d) percentage of saline volume, demonstrating the impact of position (x)
 532 of a vertical fracture on aquifer SWI

533 The sensitivity analysis results presented here further expanded the findings derived in section 4.1.
 534 Horizontally oriented fractures were proven to have a significant impact in both TL, WMZ and the total
 535 volume of intruding saltwater. In cases where horizontal discontinuities were present further toward
 536 the freshwater reservoir, they tended to limit the extent of saline intrusion, while the opposite
 537 occurred as the fractures were closer to the sea. Even though TL values were relatively unaffected by
 538 the depth in which such fractures were located, their vertical position determined the height up to
 539 which saltwater would intrude. To that effect, the fractures acted as high permeability barriers
 540 confining SWI to the lower aquifer levels. This indicated that the impact of horizontal fractures on
 541 saline intrusion dynamics was greater with increasing length. Finally, vertical fractures widened the

542 freshwater - saltwater mixing zone, with their effect proving greater when positioned in the middle of
543 the intruding wedges.

544 **5. Conclusions**

545 Saltwater intrusion in fractured unconfined coastal aquifers was studied on a laboratory scale for the
546 first time in the current paper. Saline intrusion was initiated in a thin sandbox setup by applying three
547 distinct hydraulic gradients. A total of six fractured systems, with discontinuities of different size and
548 orientation, were tested alongside one homogeneous aquifer with similar porous medium attributes.
549 Automated image analysis enabled the recreation of saltwater concentration fields as well as the
550 precise calculation of three SWI characteristics: the length of intruding wedge, the width of the mixing
551 zone and the aquifer volume fraction occupied by saltwater. The experimental results were
552 successfully simulated using a dual porosity model in SUTRA. This model was subsequently employed
553 in conducting a detailed sensitivity analysis to validate and expand the findings derived from the
554 laboratory data.

555 It was established that the presence of fractures significantly impacts all three saltwater intrusion
556 variables. The impact of four specific fracture attributes was investigated, the fracture's horizontal
557 and vertical position, its length and its orientation. The following conclusions can be drawn:

- 558 • In cases where horizontal discontinuities were present closer to the aquifer's saline boundary,
559 the extent of saltwater intrusion was generally bigger, while in systems where the fractures
560 were further away from the sea, the saline wedge was reduced. Rigorous sensitivity analysis
561 revealed that the extent to which a discontinuity affects the toe length of the intruding wedge
562 depends on two variables: its horizontal position and the hydraulic gradient applied in the
563 system. A critical hydraulic head difference was identified as the defining factor behind the
564 discontinuity's impact on saline intrusion for each specific fractured system. For hydraulic
565 gradients steeper than this head difference, the fracture's presence limits the intrusion, while
566 for milder gradients it intensifies it.
- 567 • The vertical distance of horizontal fractures from the aquifer's free surface was positively
568 correlated to the generated saline wedge length and width of mixing zone, while negatively
569 correlated to the percentage of saltwater volume. Discontinuities acted similar to high
570 permeability barriers, confining the saline wedge in the aquifer's lower level.
- 571 • The longer the discontinuities were, the more pronounced their effect was on all three
572 saltwater intrusion characteristics.
- 573 • Vertical fractures had a limited effect on the length of intrusion and the total aquifer fraction
574 occupied by saltwater. On the other hand, they had a significant impact on the width of the
575 mixing zone, resulting in its increase by a factor of up to four, under the scenarios investigated.
576 This effect was more intense in cases where the fractures were located in the middle of the
577 saline wedges.
- 578 • Whenever the discontinuities were in contact to the aquifer's freshwater or saltwater side
579 boundaries, the generated concentration fields deviated significantly from the rest of the
580 cases, indicating a critical alteration of the underlying flow mechanisms.

581 This study outlined the effects of individual fractures on the dynamics of saltwater intrusion in coastal
582 aquifers. The two-dimensional nature of the laboratory setup, alongside the simplified geometry of
583 the investigated fractures and the homogeneous nature of the surrounding porous medium, enabled
584 the identification of the fundamental physical flow mechanisms generated by the presence of high
585 permeability - high porosity discontinuities without the uncertainty that usually accompanies the
586 study of SWI on field scale investigations. Even though derived for idealized conditions, the findings

587 of this study could be of significant importance for the effective management of real-world fractured
588 aquifers. Moreover, the precise laboratory data presented in this study may be employed to
589 benchmark numerical models studying saline intrusion in fractured media. Future work could expand
590 the conclusions of this study by investigating more complex configurations of fractures with varying
591 size and orientation, as well as different boundary conditions and porous medium characteristics.

592 **Acknowledgements**

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